

# **<sup>1</sup> Weakening of Jupiter's main auroral emission during <sup>2</sup> January 2014**

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**Key Points.**

- Jupiter's auroral power decreased by 70% over 2 weeks of observations by the Hubble Space Telescope
- Could be caused by expansion of the magnetosphere or increase in hot plasma transport
- Aurora is variable without enhanced Io volcanism or solar wind pressure – implications for Juno

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3 In January 2014 Jupiter's FUV main auro-  
4 ral oval decreased its emitted power by 70%  
5 and shifted equatorward by  $\sim 1^\circ$ . Intense,  
6 low latitude features were also detected. The  
7 decrease in emitted power is attributed to a  
8 decrease in auroral current density rather than  
9 electron energy. This could be caused by a de-  
10 crease in the source electron density, an or-  
11 der of magnitude increase in the source elec-  
12 tron thermal energy, or a combination of these.  
13 Both can be explained either by expansion of  
14 the magnetosphere, or by an increase in the  
15 inward transport of hot plasma through the  
16 middle magnetosphere and its interchange with  
17 cold flux tubes moving outward. In the lat-  
18 ter case the hot plasma could have increased  
19 the electron temperature in the source region  
20 and produced the intense, low latitude features,  
21 while the increased cold plasma transport rate  
22 produced the shift of the main oval.

## 1. Introduction

Auroral images provide a valuable way to remotely observe magnetospheric dynamics. At the gas giant Jupiter there are distinct regions of auroral emissions corresponding to different regions of the magnetosphere. At the lowest latitudes are the auroral footprint spots of the moons Io, Europa, and Ganymede, which are caused by the perturbation of the planet's magnetic field as it rotates past these conducting bodies [Connerney *et al.*, 1993; Clarke *et al.*, 2002; Bonfond, 2012]. The main emission encircling the magnetic poles is associated with magnetosphere-ionosphere coupling currents acting to transfer angular momentum from the planet to the sub-corotating iogenic plasma in the middle magnetosphere at  $\sim 20 - 30 R_J$  [Cowley and Bunce, 2001; Hill, 2001; Grodent *et al.*, 2003a]. The dynamic, patchy 'polar' region inside the main emission may partly map to field lines in the outer magnetosphere or connected to the interplanetary magnetic field in the solar wind [Pallier and Prangé, 2001; Gladstone *et al.*, 2002; Grodent *et al.*, 2003b; Vogt *et al.*, 2011]. A diffuse equatorward arc is sometimes apparent and corresponds to a transition at 10–17  $R_J$  in the magnetosphere from field-perpendicular (smaller radial distances) to field-aligned (larger radial distances) electron distributions, where the radial distance of the transition varies from orbit to orbit. At radial distances outside the transition, electrons are thought to be scattered to the field-aligned distribution (and thus into the loss cone) by whistler waves [Tomás *et al.*, 2004; Radioti *et al.*, 2009]. Longitudinally confined, diffuse 'low latitude' emissions are often observed in a similar region between the main emission and the Io footprint, and are possibly associated with

injections of energetic electrons detected by Galileo at radial distances of 9–27  $R_J$  [Mauk *et al.*, 1999, 2002; Bonfond *et al.*, 2012; Dumont *et al.*, 2014].

The components of the aurora display variability on timescales of seconds to weeks, which can be interpreted as a response to solar wind influence superposed on internal magnetospheric dynamics [e.g. Nichols *et al.*, 2009]. A compression of the magnetosphere by the solar wind is expected to cause the main emission to dim as the mass-loaded field lines conserve angular momentum as they are displaced radially inward [Southwood and Kivelson, 2001; Cowley *et al.*, 2007]. However, the timescales on which the compression propagates through the magnetosphere and the neutral atmosphere responds are not well constrained so that brief increases in the main oval field-aligned currents may also occur [Cowley *et al.*, 2007; Yates *et al.*, 2014]. Cassini observations demonstrated auroral brightenings related to solar wind compressions at Jupiter but the auroral observations did not have sufficient spatial resolution to identify which auroral region(s) became brighter [Gurnett *et al.*, 2002; Pryor *et al.*, 2005]. Overall, ambiguity in the timing of solar wind conditions arriving at Jupiter and the limited cadence of auroral imaging have not yet allowed the full auroral response to solar wind compressions or rarefactions to be conclusively identified [Nichols *et al.*, 2007, 2009; Clarke *et al.*, 2009].

The auroral emissions also demonstrate a response to changes in the inner magnetosphere related to the mass-loading and field-stretching. Grodent *et al.* [2008] and Bonfond *et al.* [2012] suggested that movement of the main oval to lower latitudes, observed in images separated by months or years, could be caused by a change in the magnetic field stretching or an inward shift in the corotation breakdown boundary. These effects were

related to an increase in mass-loading from Io [e.g. *Yoneda et al.*, 2010]. An increase in the outflow rate of iogenic plasma is expected to affect the intensity of the main aurora but whether it increases or decreases depends on the model employed [*Nichols and Cowley*, 2003; *Nichols*, 2011; *Ray et al.*, 2012].

In this study a two-week sequence of auroral observations is used to investigate the variation in both the intensity and location of Jupiter's aurora in relation to magnetospheric conditions.

## 2. Auroral observations

### 2.1. Data

Jupiter's northern aurora was observed using the Hubble Space Telescope (HST) Space Telescope Imaging Spectrograph (STIS) during 14 'visits' (i.e. observation sequences) over 16 days in January 2014. Images were acquired using the SrF2 longpass filter, which excludes H Lyman-alpha emission at 121.6 nm but covers the H<sub>2</sub> Lyman and Werner bands in the range 125–190 nm. The data were processed using a pipeline developed at Boston University, including dark count subtraction, flat-fielding, geometric distortion correction, scaling to a standard opposition distance between HST and Jupiter of 4.2 AU, and subtraction of an empirical disk background [*Clarke et al.*, 2009; *Nichols et al.*, 2009]. The images were projected onto a planetocentric latitude and System III longitude grid at an emission altitude of 240 km above the 1-bar pressure level [*Vasavada et al.*, 1999]. The spatial uncertainties in the projected images come mainly from determining the centre of the planet and the 'stretching' of pixels close to the planet's limb; these uncertainties

are fully described by *Grodent et al.* [2003a], who show that the inaccuracies are typically  $\sim 1^\circ$  for the main auroral oval observation geometry.

Observations were made in two sets of duration  $\sim 700$  s on each HST visit. For images shown in this study the photon counts were integrated over intervals of 100 s to achieve both good temporal resolution and signal-to-noise. The counts were converted to a brightness in kR using the conversion factor given by *Gustin et al.* [2012] of  $1 \text{ kR} = 2.211 \times 10^{-4}$  counts. This assumes a colour ratio of 2.5 across the auroral region, as inferred from STIS spectral observations made during the same campaign [*Tao et al.*, 2016], where the colour ratio is the ratio of intensity in a UV wavelength band unabsorbed by atmospheric hydrocarbons (155–162 nm) to the intensity in an absorbed band (123–130 nm), i.e. a measure of auroral electron penetration depth and hence electron energy. The auroral powers quoted below correspond to the auroral  $\text{H}_2$  emission across a wavelength range of 70–180 nm [*Gustin et al.*, 2012].

## 2.2. Auroral Morphology

One image from each of the 14 HST visits is shown in Figure 1. At the start of the campaign, on days 1–3, the main oval was bright and composed of narrow arcs at most longitudes (Figure 1a–c).

On day 4 (Figure 1d) the auroral morphology was noticeably different: the main oval was dimmer at all longitudes than in the previous images, and the brightest emission came from an extended region of diffuse emission at longitudes  $140 - 190^\circ$ . This region is highlighted by the red line on the image.

The diffuse emission was fixed in SIII longitude, i.e. was corotating with the planet, and the Ganymede footprint could be observed moving out of this western edge of the diffuse structure over the sequence although it is not distinct in the snapshot shown. The diffuse emission extended across  $\sim 3 - 4^\circ$  latitude, from the main oval to  $\sim 1^\circ$  poleward of the Io footprint contour (the Io footprint itself was not captured in these images).

Approximately 25 h later, on day 5, the diffuse equatorward feature had disappeared and the main oval was slightly higher intensity again (Figure 1e). Similar morphologies were observed in the subsequent images taken on days 6, 7, and 10, shown in Figure 1f–h.

The first of two sets of images on day 11 (Figure 1g) shows another very different auroral morphology. The main oval region was formed of bright patches. Large regions of equatorward emission were observed, extending from one of the main oval patches at longitudes  $185 - 190^\circ$ , and as a distinct equatorward feature at longitudes  $135 - 170^\circ$  (highlighted by red lines). The Ganymede footprint was observed to move between these two structures over the interval but again is not visible in the snapshot shown. Some bright polar features were observed. The second set of images on day 11 began  $\sim 18$  h later and reveal that all regions of the aurora had become fainter over this interval.

The main oval remained relatively dim and accompanied by the faint secondary arc for the rest of the observations on days 13 and 16. The brightest arcs along the main oval were in the longitude sector  $100 - 160^\circ$ . Some equatorward patches were also observed (e.g. early on day 13) but they were not as large or bright as those observed on days 4 and 11.



It is clear from the images and above discussion that many intriguing features were observed in different regions of the aurora, representing different magnetospheric dynamics, over the duration of the campaign in January 2014. In the subsequent sections we focus on one aspect of the auroral variability: the power emitted from different regions as a function of time.

### 2.3. Auroral Power

To quantify the variability of the auroral power the auroral region was sub-divided into different latitudinal regions, corresponding to different source regions in the magnetosphere, following *Nichols et al.* [2009]. The main oval region was defined as a strip  $2^\circ$  wide in latitude, centred on the average main oval determined from all images. The polar region was defined as the region poleward of this, and the low latitude region was the region equatorward of the main oval region, up to a contour  $1.5^\circ$  poleward of the Io footprint contour defined by *Bonfond et al.* [2009]. The average emission intensity over the campaign is shown in Figure 1o with these boundaries overlaid.

The fraction of Jupiter's auroral region visible to HST varies as the planet rotates because of the offset of the magnetic axis from the spin axis. This variability needs to be accounted for so that powers from different images can be compared. To achieve this the observed powers were scaled by a function representing the observable fraction of the auroral region for all CML, following the method described by *Nichols et al.* [2009]. This assumes that the auroral emission is roughly homogenous over each region. The corrected powers are shown as a function of time in Figure 2. Panels (b)–(d) present the power emitted in the main oval, low latitude, and polar regions, respectively. The dotted lines

show the mean value across the observations, and the grey shading indicates the standard deviation from this value. The variation of the total power summed over these regions is shown by the crosses in Figure 2a for each 100 s integration. The black dotted line and grey shading in the top panel represent the mean total power and the standard deviation of the values over the campaign.

The power emitted from the main oval declined gradually over the campaign, with the exception of visit 9 on day 11, during which a localised bright patch extended across the main oval latitudes (see Figure 1i). This feature was the brightest of the campaign and affected the power in both the main oval and low latitude regions. The average main oval emitted power on days 1–2 was  $\sim 480$  GW, decreasing to  $\sim 170$  GW on days 13–16.

The polar region also emitted low powers at the end of day 13 and on day 16, however, a general decrease in the polar power over the whole campaign is not apparent. The overall standard deviation of the polar emitted power was lower than that of the main oval power but individual days show much greater variation, i.e. days 7–13. This indicates that the intensity of the polar region is highly variable on minute timescales.

The low latitude region showed little variation in emitted power over the campaign (average 395 GW) with the exception of two large increases on days 4 and 11. The total power also shows a net decrease in emitted power over the campaign, in line with the reduced contribution from the main oval. It falls from an average power of 1380 GW on days 1 and 2 to an average of 900 GW on days 13 and 16.

The decrease in total auroral power captured by the HST observations was also detected by the Hisaki/EXCEED mission [Yoshikawa *et al.*, 2014; Yamazaki *et al.*, 2014], which

monitored Jupiter's EUV auroral emission quasi-continuously during December 2013–  
 March 2014 [Kimura *et al.*, 2015]. The total EUV auroral power over 90–148 nm detected  
 by Hisaki is represented in Figure 2a by the solid line, where the values have been averaged  
 using a running median with window 39.7 h, i.e. four jovian rotations, to remove the quasi-  
 sinusoidal variation imposed by the planetary rotation, and scaled by a factor of 4 for ease  
 of viewing on this scale. (Full details of the Hisaki auroral power estimation are given  
 by Kimura *et al.* [2015].). The smoothed EUV power decreased from  $\sim 320$  GW on days  
 1–2 to  $\sim 270$  GW on days 13–16. A decrease in auroral power over these timescales  
 was previously identified from International Ultraviolet Explorer observations [Prangé  
*et al.*, 2001], which also lacked spatial resolution. The HST observations provide spatially  
 resolved images from which we can determine that the overall decrease in power over this  
 two-week interval was mainly driven by a decrease in the emission from the main oval.

## 2.4. Auroral Location

Figure 1p shows the location of the peak brightness at certain longitudes, tracing out  
 the main oval, for selected HST visits at the start (1, 3 Jan) and end (13, 16 Jan) of  
 the interval. The position of the peak brightness had shifted slightly equatorward, by  
 an average of  $1^\circ$ , at the end of the campaign compared to at the start. For comparison,  
 the latitude of the main oval can vary over a full visit ( $2 \times 700$  s) by  $0\text{--}0.5^\circ$  on average,  
 while the maximum displacement along a given line of longitude across all visits is  $2.5^\circ$   
 (excluding regions where the main oval could not be precisely located because of e.g.  
 proximity to the edge of the field-of-view or where the auroral oval was particularly faint  
 or diffuse). We take the  $1^\circ$  shift between days 1 and 16 as representative of the long-term

equatorward shift, while acknowledging that this neglects variability on shorter timescales. The magnitude of the observed shift is comparable to the expansion of the main oval previously identified over longer intervals [*Grodent et al.*, 2008; *Bonfond et al.*, 2012]. Although the magnitude of the shift is comparable to the spatial uncertainties described above, the fact that it represents a long-term trend rather than random fluctuations leads us to consider this shift as real.

### 3. Causes of Auroral Variability

A decrease in main oval intensity would be caused by a reduction in auroral electron energy flux deposited in the upper atmosphere. This is related to the magnitude of the field-aligned current linking the ionosphere and the corotation-breakdown region in the equatorial magnetosphere. A decrease in the mass loading of the field lines or a reduction in their radial stretching could result in a lower auroral field-aligned current [e.g. *Nichols*, 2011]. One possible cause for a reduction in the radial stretch of the magnetic field lines is a global compression of the magnetosphere by the arrival of a high pressure solar wind region. The solar wind conditions at Jupiter can be estimated using a 1-D MHD code [*Tao et al.*, 2005] to propagate the solar wind measured near Earth out to 5 AU. The uncertainty in the arrival times is less than  $\pm 24$  h at this time because of the small ( $< 25^\circ$ ) Earth-Sun-Jupiter angle. The propagated dynamic pressure is presented in Figure 2e, and shows that the HST auroral observations took place during an interval of decreasing solar wind pressure and radial velocity. This would result in an expansion of the magnetosphere and, assuming conservation of angular momentum, associated increase in the auroral currents [*Cowley et al.*, 2007; *Yates et al.*, 2014], opposite to what is inferred from the auroral

210 observations. We therefore examine other possible causes of a decrease in field-aligned  
 211 current and auroral electron energy flux.

212 Using Hisaki/EXCEED spectra, *Tao et al.* [2015, 2016] showed that the mean energy of  
 213 the electrons precipitating into the main oval remained roughly constant over this cam-  
 214 paign. This implies that the observed decrease in precipitating energy flux is associated  
 215 with a decrease in electron number flux (equivalent to the current density) rather than  
 216 electron energy. The variation in magnetospheric parameters which could cause the ob-  
 217 served decrease in the auroral current density can be examined using the *Knight* [1973]  
 218 relation. The maximum upward current density that can be carried by magnetospheric  
 219 electrons without field-aligned acceleration is

$$j_{||0} = eN \left( \frac{W_{th}}{2\pi m_e} \right)^{1/2}, \quad (1)$$

220 where  $e$  and  $m_e$  are the charge and mass of the electron, and  $N$  and  $W_{th}$  are the number  
 221 density and thermal energy of the source electron population in the magnetosphere. This  
 222 relation assumes a full down-going loss cone and empty up-going loss cone. The field-  
 223 aligned energy flux of these electrons precipitating into the ionosphere is

$$E_{f0} = 2NW_{th} \left( \frac{W_{th}}{2\pi m_e} \right)^{1/2}. \quad (2)$$

224 The current density can be enhanced by a field-aligned potential drop to produce the  
 225 current required in the middle magnetosphere coupling system. Using the linear approx-  
 226 imation to the Knight relation, the enhanced current density just above the ionosphere,

$j_{||}$  results in an increased field-aligned energy flux of the precipitating electrons given by [Lundin and Sandahl, 1978]:

$$E_f = \frac{E_{f0}}{2} \left[ \left( \frac{j_{||}}{j_{||0}} \right)^2 + 1 \right]. \quad (3)$$

The energy flux can be estimated from the observed brightness of the main oval, using the relation that 1 mW m<sup>-2</sup> incident energy flux produces 10 kR of auroral intensity [Gustin et al., 2012, and references therein]. The mean intensity in the main oval region and the derived electron energy flux are shown as a time series in Figure 3a and b. The right hand axis of (b) indicates the corresponding current density  $j_{||}$ , determined by assuming the energy flux is deposited by electrons with mean energy  $\langle W \rangle = 150$  keV as indicated by spectral observations [Tao et al., 2015, 2016; Gérard et al., 2014].

From relations 1–3 above, the average incident energy of the electrons  $\langle W \rangle$  can be expressed in terms of the magnetosphere source electron parameters,  $N$  and  $W_{th}$ , by taking the ratio of the electron energy flux and number flux ( $j_{||}/e$ ), and  $E_f \gg E_{f0}$  [e.g. Gustin et al., 2004]:

$$\langle W \rangle \approx \sqrt{2} W_{th} \left( \frac{E_f}{E_{f0}} \right)^{1/2} \propto \frac{W_{th}^{1/4}}{N^{1/2}} E_f^{1/2}, \quad (4)$$

Figure 3b shows that the precipitating energy flux is reduced by a factor of  $\sim 35/10 \sim 3.5$  (or, a 70% decrease) over the observing interval. Holding  $\langle W \rangle$  constant, as demonstrated by Tao et al. [2015], Equation 4 shows that this reduction in  $E_f$  can be attributed to a factor of  $\sim 3.5$  decrease in  $N$  if  $W_{th}$  also remained constant (fewer current carriers available), or a factor  $\sim 12$  increase in  $W_{th}$  if  $N$  remained constant (as  $j_{||}$  depends on the

difference between  $W_{th}$  and  $\langle W \rangle$ ). These variations in  $N$  and  $W_{th}$  are also shown in Figure 3c and d for the cases where  $W_{th}$  is fixed at 2.5 keV (c) and  $N$  is fixed at  $0.0026 \text{ cc}^{-1}$  (d). These fixed values are taken from the range determined from observations [*Gustin et al.*, 2004; *Tao et al.*, 2015]. From the observations, it is not possible to isolate which of these parameters is varying and it could be a combination of the two. *Ray et al.* [2012] evaluated the full Knight relation (not linear approximation) applied to Jupiter's main auroral currents and showed that the observed change in precipitating electron energy flux could be produced by a similar decrease in  $N$  to that found above, while a lesser dependence on  $W_{th}$  is suggested from their results, although a smaller range of  $W_{th} \leq 5 \text{ keV}$  was considered. In general, if  $\langle W \rangle$  is constant and  $E_{f,2} < E_{f,1}$ , where the subscripts 1 and 2 denote the measurement at the start and end of the interval, respectively, Equation 4 becomes:

$$\sqrt{\frac{W_{th,2}}{W_{th,1}}} \frac{N_1}{N_2} > 1. \quad (5)$$

For example, one possible explanation for the observed variations is an expansion of the magnetosphere under the prevailing decrease in solar wind pressure. Under an adiabatic expansion  $PV^\gamma$  is constant, where  $P = NkT_0$  is the pressure,  $V$  is the flux tube volume, and  $\gamma = 5/3$ . Through conservation of mass (i.e.  $NV = \text{constant}$ ) we obtain  $N^{-2/3}T_0 = \text{constant}$ , and, as  $W_{th} \propto T_0$ ,  $W_{th} \propto N^{2/3}$ . Inserting this relation into Equation 5 we see that this condition on the variation in  $W_{th}$  and  $N$  can be satisfied by an adiabatic expansion. Non-adiabatic expansion in which  $N$  decreases while satisfying Equation 5 is also possible. As mentioned above, this treatment of the magnetospheric expansion neglects the effect

that conservation of angular momentum would have on the field-aligned magnetosphere-ionosphere coupling currents [Southwood and Kivelson, 2001; Cowley et al., 2007; Yates et al., 2014]. While the Yates et al. [2014] model reproduced the equatorward shift in the auroral oval under a transient magnetospheric expansion, the shift was accompanied by an overall increase in main oval intensity which was not observed during this campaign.

The source auroral electrons with energies  $W_{th}$  of a few keV are considered to be the warm ‘tail’ of the population present in the middle magnetosphere. An alternative scenario to explain the observations is related to an increase in hot plasma transport through this region, which increases the temperature of the warm electrons available to carry the auroral current. The inward transport of hot plasma has been observed as narrow, isolated structures in the Io torus [Kivelson et al., 1997; Thorne et al., 1997] and as larger, energy-dispersed ‘injections’ detected out to 27  $R_J$  [Mauk et al., 1999, 2002]. To conserve magnetic flux, flux tubes loaded with cold plasma must also move outward to replace the inward, hot flux tubes. We explore this scenario because possible signatures of the hot plasma injections are observed in the aurora as the so-called low latitude emissions, and those seen on 4 and 11 Jan 2014 (Figures 1d and i) are among the largest and brightest compared to the main emission [Mauk et al., 2002; Nichols et al., 2009; Bonfond et al., 2012; Dumont et al., 2014].

As the enhanced interchange of outward, cold plasma increases the mass outflow rate, models predict an equatorward shift of the main emission as observed in Figure 1p. In some of the models this is accompanied by an increased [Nichols, 2011] or constant [Ray et al., 2012] auroral current density and brightness, in contrast to the decrease in auroral



intensity observed. *Nichols* [2011] showed that a decrease in the auroral current density can be obtained if the increased mass outflow is driven by an increased rate of outward transport, rather than an increase in the cold plasma density. This is consistent with the interpretation given above: an increase in interchange-driven outflow and in electron temperature ( $W_{th}$ ) while the density ( $N$ ) remains constant. *Nichols* [2011] showed that a decrease in auroral current density of the magnitude shown in Figure 3 can be produced by a relatively modest, e.g.  $\sim 2\times$ , change in the mass outflow rate.

A decrease in the UV main emission intensity, an equatorward shift in the main emission, and increased occurrence of low latitude emissions can also be identified during an earlier set of observations made in 2007 [*Nichols et al.*, 2009; *Bonfond et al.*, 2012]. *Bonfond et al.* [2012] attributed these effects to an increase in Io volcanic activity, demonstrated by an increase in the brightness of the Io sodium nebula [*Yoneda et al.*, 2009]. *Yoneda et al.* [2013] also showed a decrease in the intensity of jovian hectometric auroral radiation following the enhanced Io volcanic activity in 2007. Observations of Io's sodium nebula presented by *Yoneda et al.* [2015] show there was no such increase in the nebula brightness detected in the weeks preceding and encompassing the interval in Jan 2014 discussed here. Similarly, *Tsuchiya et al.* [2015] presented Hisaki observations demonstrating that there was no increase in the EUV intensity emitted from the inner Io plasma torus which would be indicative of enhanced Iogenic mass loading. The Jan 2014 observations suggest that a decrease in auroral current strength and the presence of hot plasma injection events represented by low latitude auroral patches can be triggered without a significant change in Io volcanic activity.

#### 4. Conclusions

Jupiter's main auroral oval was observed to decrease in intensity by 70% and shift slightly ( $\sim 1^\circ$ ) equatorward over a two week interval of observations in Jan 2014. The decrease in auroral intensity represents a decrease in the electron energy flux precipitating into the ionosphere, which can be caused by a variation in the magnetospheric source electron number density and/or thermal energy. To reproduce the observations, a 70% decrease in the source electron density or a factor of 12 increase in their thermal energy is required (if the other parameter is held constant). One possible explanation for the observations is an expansion of the magnetosphere under the prevailing gradual decrease in solar wind dynamic pressure. An alternative explanation for the observations is an increase in the transport rate of hot plasma through the auroral current source region in the middle magnetosphere. Possible signatures of large, hot plasma injections were observed as diffuse, low latitude auroral patches. The corresponding increase in outward transport of cold flux tubes required to conserve magnetic flux could lead to the observed equatorward shift in the auroral oval. We conclude that the observed decrease in the main oval intensity does not require a change in the mass loading rate from Io or compression by the solar wind as previously suggested.

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## References

- Bonfond, B. (2012), When Moons Create Aurora: The Satellite Footprints on Giant Planets, *Washington DC American Geophysical Union Geophysical Monograph Series*, 197, 133–140, doi:10.1029/2011GM001169.
- Bonfond, B., D. Grodent, J.-C. Gérard, A. Radioti, V. Dols, P. A. Delamere, and J. T. Clarke (2009), The Io UV footprint: Location, inter-spot distances and tail vertical extent, *J. Geophys. Res.*, 114, doi:10.1029/2009JA014312.
- Bonfond, B., D. Grodent, J.-C. Gérard, T. Stallard, J. T. Clarke, M. Yoneda, A. Radioti, and J. Gustin (2012), Auroral evidence of Io’s control over the magnetosphere of Jupiter, *Geophys. Res. Lett.*, 39, doi:10.1029/2011GL050253.
- Clarke, J. T., J. Ajello, G. Ballester, L. Ben Jaffel, J. Connerney, J.-C. Gérard, G. R. Gladstone, D. Grodent, W. Pryor, J. Trauger, and J. H. Waite (2002), Ultraviolet emissions from the magnetic footprints of Io, Ganymede and Europa on Jupiter, *Nature*, 415(6875), 997–1000.
- Clarke, J. T., J. Nichols, J. C. Gérard, D. Grodent, K. C. Hansen, W. Kurth, G. R. Gladstone, J. Duval, S. Wannawichian, E. Bunce, S. W. H. Cowley, F. Crary, M. Dougherty,

L. Lamy, D. Mitchell, W. Pryor, K. Retherford, T. Stallard, B. Zieger, P. Zarka, and  
B. Cecconi (2009), Response of Jupiter's and Saturn's auroral activity to the solar wind,  
*J. Geophys. Res.*, *114*, A05210, doi:10.1029/2008JA013694.

Connerney, J. E. P., R. Baron, T. Satoh, and T. Owen (1993), Images of Excited  $H_3^+$   
at the Foot of the Io Flux Tube in Jupiter's Atmosphere, *Science*, *262*, 1035–1038,  
doi:10.1126/science.262.5136.1035.

Cowley, S. W. H., and E. J. Bunce (2001), Origin of the main auroral oval in Jupiter's  
coupled magnetosphere-ionosphere system, *Planet. Space. Sci.*, *49*, 1067–1088, doi:  
10.1016/S0032-0633(00)00167-7.

Cowley, S. W. H., J. D. Nichols, and D. J. Andrews (2007), Modulation of Jupiter's plasma  
flow, polar currents, and auroral precipitation by solar wind-induced compressions and  
expansions of the magnetosphere: a simple theoretical model, *Ann. Geophys.*, *25*, 1433–  
1463, doi:10.5194/angeo-25-1433-2007.

Dumont, M., D. Grodent, A. Radioti, and J.-C. Gérard (2014), Jupiter's equatorward  
auroral features: Possible signatures of magnetospheric injections, *J. Geophys. Res.*,  
*119*(12), 10,068–10,077, doi:10.1002/2014JA020527.

Gérard, J.-C., B. Bonfond, D. Grodent, A. Radioti, J. T. Clarke, G. R. Gladstone, J. H.  
Waite, D. Bisikalo, and V. I. Shematovich (2014), Mapping the electron energy in  
Jupiter's aurora: Hubble spectral observations, *J. Geophys. Res.*, *119*, 9072–9088, doi:  
10.1002/2014JA020514.

Gladstone, G. R., J. H. Waite, D. Grodent, W. S. Lewis, F. J. Crary, R. F. Elsner, M. C.  
Weisskopf, T. Majeed, J.-M. Jahn, A. Bhardwaj, J. T. Clarke, D. T. Young, M. K.

Dougherty, S. A. Espinosa, and T. E. Cravens (2002), A pulsating auroral X-ray hot spot on Jupiter, *Nature*, *415*, 1000–1003.

Grodent, D., J. T. Clarke, J. Kim, J. H. Waite, and S. W. H. Cowley (2003a), Jupiter's main auroral oval observed with HST-STIS, *J. Geophys. Res.*, *108*, A11, doi:10.1029/2003JA009921.

Grodent, D., J. T. Clarke, J. H. Waite, S. W. H. Cowley, J.-C. Gérard, and J. Kim (2003b), Jupiter's polar auroral emissions, *J. Geophys. Res.*, *108*, 1366, doi:10.1029/2003JA010017.

Grodent, D., J.-C. Gérard, A. Radioti, B. Bonfond, and A. Saglam (2008), Jupiter's changing auroral location, *J. Geophys. Res.*, *113*(A1), doi:10.1029/2007JA012601.

Gurnett, D. A., W. S. Kurth, G. B. Hospodarsky, A. M. Persoon, P. Zarka, A. Lecacheux, S. J. Bolton, M. D. Desch, W. M. Farrell, M. L. Kaiser, H.-P. Ladreiter, H. O. Rucker, P. Galopeau, P. Louarn, D. T. Young, W. R. Pryor, and M. K. Dougherty (2002), Control of Jupiter's radio emission and aurorae by the solar wind, *Nature*, *415*, 985–987.

Gustin, J., J.-C. Gérard, D. Grodent, S. W. H. Cowley, J. T. Clarke, and A. Grard (2004), Energy-flux relationship in the FUV Jovian aurora deduced from HST-STIS spectral observations, *J. Geophys. Res.*, *109*, A10205, doi:10.1029/2003JA010365.

Gustin, J., B. Bonfond, D. Grodent, and J.-C. Gérard (2012), Conversion from HST ACS and STIS auroral counts into brightness, precipitated power, and radiated power for H<sub>2</sub> giant planets, *J. Geophys. Res.*, *117*, A07316, doi:10.1029/2012JA017607.

Hill, T. W. (2001), The Jovian auroral oval, *J. Geophys. Res.*, *106*, 8101–8108, doi:  
10.1029/2000JA000302.

Kimura, T., S. V. Badman, C. Tao, K. Yoshioka, G. Murakami, A. Yamazaki, F. Tsuchiya,  
B. Bonfond, A. J. Steffl, A. Masters, S. Kasahara, H. Hasegawa, I. Yoshikawa, M. Fujimoto,  
and J. T. Clarke (2015), Transient internally-driven aurora at Jupiter discovered by Hisaki and the Hubble Space Telescope, *Geophys. Res. Lett.*, *in press*, doi:  
10.1002/2015GL063272.

Kivelson, M. G., K. K. Khurana, C. T. Russell, and R. J. Walker (1997), Intermittent  
short-duration magnetic field anomalies in the Io torus: Evidence for plasma interchange?,  
*Geophys. Res. Lett.*, *24*, 2127, doi:10.1029/97GL02202.

Knight, S. (1973), Parallel electric fields, *Planet. Space. Sci.*, *21*, 741–750, doi:  
10.1016/0032-0633(73)90093-7.

Lundin, R., and I. Sandahl (1978), Some characteristics of the parallel electric field acceleration of electrons over discrete auroral arcs as observed from two rocket flights,  
*Symposium on European Rocket Research, ESA SP-135*, 125.

Mauk, B. H., D. J. Williams, R. W. McEntire, K. K. Khurana, and J. G. Roederer (1999),  
Storm-like dynamics of Jupiter's inner and middle magnetosphere, *J. Geophys. Res.*,  
*104*, 22,759–22,778, doi:10.1029/1999JA900097.

Mauk, B. H., J. T. Clarke, D. Grodent, J. H. Waite, C. P. Paranicas, and D. J. Williams  
(2002), Transient aurora on Jupiter from injections of magnetospheric electrons, *Nature*,  
*415*, 1003–1005.

- 415 Nichols, J. D. (2011), Magnetosphere-ionosphere coupling in Jupiter's middle magneto-  
416 sphere: Computations including a self-consistent current sheet magnetic field model,  
417 *J. Geophys. Res.*, *116*, A10232, doi:10.1029/2011JA016922.
- 418 Nichols, J. D., and S. W. H. Cowley (2003), Magnetosphere-ionosphere coupling currents  
419 in Jupiter's middle magnetosphere: dependence on the effective ionospheric Pedersen  
420 conductivity and iogenic plasma mass outflow rate, *Ann. Geophys.*, *21*, 1419–1441, doi:  
421 10.5194/angeo-21-1419-2003.
- 422 Nichols, J. D., E. J. Bunce, J. T. Clarke, S. W. H. Cowley, J.-C. Gérard, D. Grodent,  
423 and W. R. Pryor (2007), Response of Jupiter's UV auroras to interplanetary condi-  
424 tions as observed by the Hubble Space Telescope during the Cassini flyby campaign,  
425 *J. Geophys. Res.*, *112*, A02203, doi:10.1029/2006JA012005.
- 426 Nichols, J. D., J. T. Clarke, J. C. Gérard, D. Grodent, and K. C. Hansen (2009), Variation  
427 of different components of Jupiter's auroral emission, *J. Geophys. Res.*, *114*, A06210,  
428 doi:10.1029/2009JA014051.
- 429 Pallier, L., and R. Prangé (2001), More about the structure of the high latitude Jovian  
430 aurorae, *Planet. Space. Sci.*, *49*, 1159–1173, doi:10.1016/S0032-0633(01)00023-X.
- 431 Prangé, R., G. Chagnon, M. G. Kivelson, T. A. Livengood, and W. Kurth (2001), Tempo-  
432 ral monitoring of Jupiter's auroral activity with IUE during the Galileo mission. Implica-  
433 tions for magnetospheric processes, *Planet. Space. Sci.*, *49*, 405–415, doi:10.1016/S0032-  
434 0633(00)00161-6.
- 435 Pryor, W. R., A. I. F. Stewart, L. W. Esposito, W. E. McClintock, J. E. Colwell, A. J.  
436 Jouchoux, A. J. Steffl, D. E. Shemansky, J. M. Ajello, R. A. West, C. J. Hansen, B. T.

Tsurutani, W. S. Kurth, G. B. Hospodarsky, D. A. Gurnett, K. C. Hansen, J. H. Waite,  
F. J. Crary, D. T. Young, N. Krupp, J. T. Clarke, D. Grodent, and M. K. Dougherty  
(2005), Cassini UVIS observations of Jupiter's auroral variability, *Icarus*, 178, 312–326,  
doi:10.1016/j.icarus.2005.05.021.

Radioti, A., A. T. Tomás, D. Grodent, J.-C. Gérard, J. Gustin, B. Bonford, N. Krupp,  
J. Woch, and J. D. Menietti (2009), Equatorward diffuse auroral emissions at Jupiter:  
Simultaneous HST and Galileo observations, *Geophys. Res. Lett.*, 36, L07101, doi:  
10.1029/2009GL037857.

Ray, L. C., R. E. Ergun, P. A. Delamere, and F. Bagenal (2012), Magnetosphere-  
ionosphere coupling at Jupiter: A parameter space study, *Journal of Geophysical Re-  
search (Space Physics)*, 117, A01205, doi:10.1029/2011JA016899.

Southwood, D. J., and M. G. Kivelson (2001), A new perspective concerning the influence  
of the solar wind on the Jovian magnetosphere, *J. Geophys. Res.*, 106, 6123–6130,  
doi:10.1029/2000JA000236.

Tao, C., R. Kataoka, H. Fukunishi, Y. Takahashi, and T. Yokoyama (2005), Magnetic  
field variations in the Jovian magnetotail induced by solar wind dynamic pressure en-  
hancements, *J. Geophys. Res.*, 110(A11), doi:10.1029/2004JA010959.

Tao, C., T. Kimura, S. V. Badman, N. André, F. Tsuchiya, G. Murakami, K. Yoshioka,  
I. Yoshikawa, A. Yamazaki, and M. Fujimoto (2015), Variation of Jupiter's aurora  
observed by Hisaki/EXCEED: 2. Estimations of auroral parameters and magnetospheric  
dynamics, *J. Geophys. Res.*, *in press*, doi:10.1002/2015JA021272.



- 458 Tao, C., T. Kimura, S. V. Badman, G. Murakami, K. Yoshioka, F. Tsuchiya, N. André,  
459 I. Yoshikawa, A. Yamazaki, D. Shiota, H. Tadokoro, and M. Fujimoto (2016), Variation  
460 of Jupiter's aurora observed by Hisaki/EXCEED: 1. Observed characteristics of auroral  
461 electron energies compared with HST/STIS observations, *J. Geophys. Res.*, *submitted*.
- 462 Thorne, R. M., T. P. Armstrong, S. Stone, D. J. Williams, R. W. McEntire, S. J. Bolton,  
463 D. A. Gurnett, and M. G. Kivelson (1997), Galileo evidence for rapid interchange trans-  
464 port in the Io torus, *Geophys. Res. Lett.*, *24*, 2131, doi:10.1029/97GL01788.
- 465 Tomás, A. T., J. Woch, N. Krupp, A. Lagg, K.-H. Glassmeier, and W. S. Kurth (2004),  
466 Energetic electrons in the inner part of the Jovian magnetosphere and their relation to  
467 auroral emissions, *J. Geophys. Res.*, *109*, A06203, doi:10.1029/2004JA010405.
- 468 Tsuchiya, F., M. Kagitani, K. Yoshioka, T. Kimura, G. Murakami, A. Yamazaki,  
469 H. Nozawa, Y. Kasaba, T. Sakanoi, K. Uemizu, and I. Yoshikawa (2015), Local elec-  
470 tron heating in Io plasma torus associated with Io from HISAKI satellite observation,  
471 *J. Geophys. Res.*, *in press*, doi:10.1002/2015JA021420.
- 472 Vasavada, A. R., A. H. Bouchez, A. P. Ingersoll, B. Little, C. D. Anger, and Galileo SSI  
473 Team (1999), Jupiter's visible aurora and Io footprint, *J. Geophys. Res.*, *104*, 27,133–  
474 27,142, doi:10.1029/1999JE001055.
- 475 Vogt, M. F., M. G. Kivelson, K. K. Khurana, R. J. Walker, B. Bonfond, D. Grodent, and  
476 A. Radioti (2011), Improved mapping of Jupiter's auroral features to magnetospheric  
477 sources, *J. Geophys. Res.*, *116*, A03220, doi:10.1029/2010JA016148.
- 478 Yamazaki, A., F. Tsuchiya, T. Sakanoi, K. Uemizu, K. Yoshioka, G. Murakami, M. Kag-  
479 itani, Y. Kasaba, I. Yoshikawa, N. Terada, T. Kimura, S. Sakai, K. Nakaya, S. Fukuda,

and S. Sawai (2014), Field-of-View Guiding Camera on the HISAKI (SPRINT-A) Satellite, *Space Sci. Rev.*, *184*, 259–274, doi:10.1007/s11214-014-0106-y.

Yates, J. N., N. Achilleos, and P. Guio (2014), Response of the Jovian thermosphere to a transient ‘pulse’ in solar wind pressure, *Planet. Space. Sci.*, *91*, 27–44, doi:10.1016/j.pss.2013.11.009.

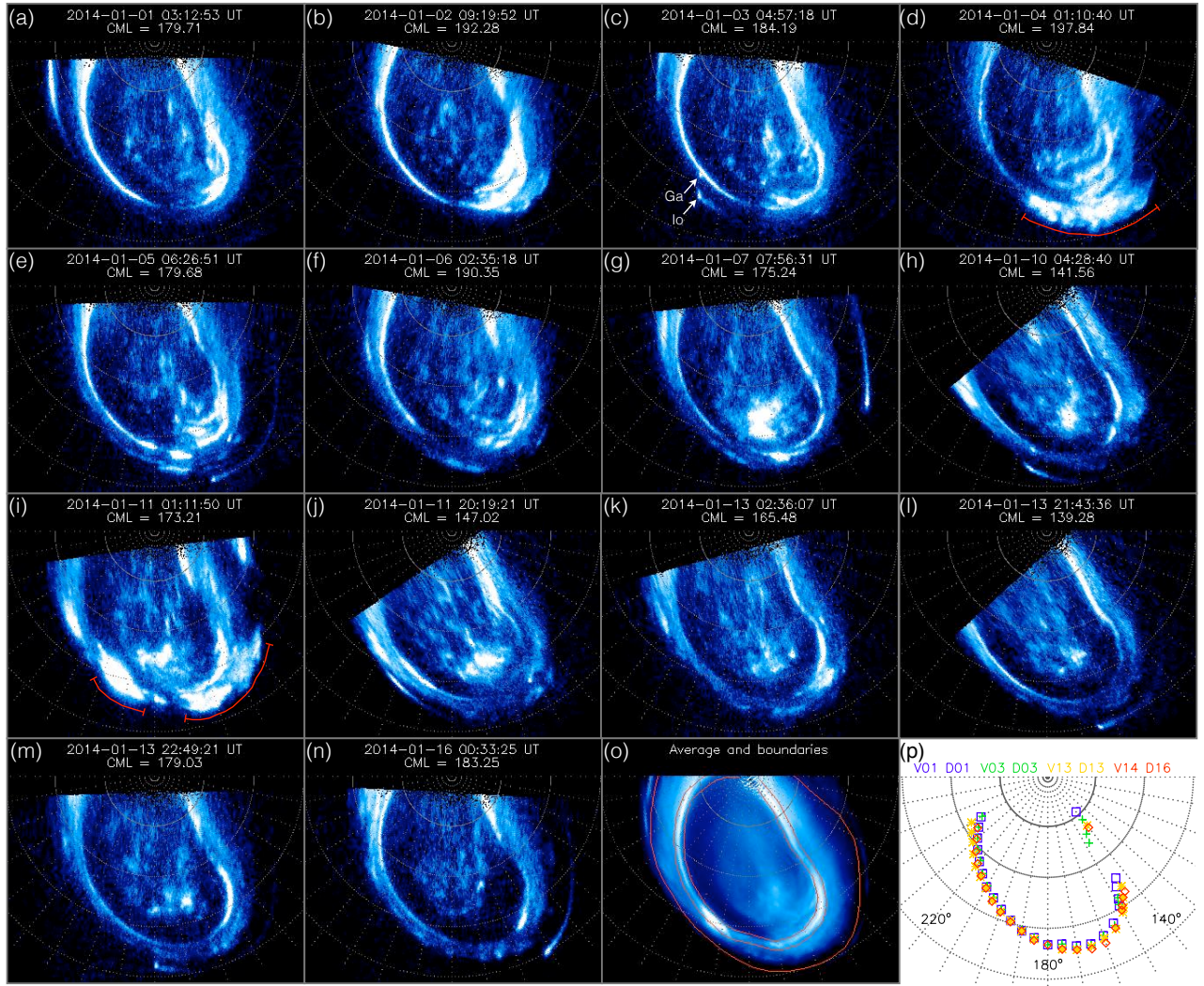
Yoneda, M., M. Kagitani, and S. Okano (2009), Short-term variability of Jupiter’s extended sodium nebula, *Icarus*, *204*(2), 589–596, doi:10.1016/j.icarus.2009.07.023.

Yoneda, M., H. Nozawa, H. Misawa, M. Kagitani, and S. Okano (2010), Jupiter’s magnetospheric change by Io’s volcanoes, *Geophys. Res. Lett.*, *37*, doi:10.1029/2010GL043656.

Yoneda, M., F. Tsuchiya, H. Misawa, B. Bonfond, C. Tao, M. Kagitani, and S. Okano (2013), Io’s volcanism controls Jupiter’s radio emissions, *Geophys. Res. Lett.*, *40*(4), 671–675.

Yoneda, M., M. Kagitani, F. Tsuchiya, T. Sakanoi, and S. Okano (2015), Brightening event seen in observations of Jupiter’s extended sodium nebula, *Icarus*, *261*, 31–33, doi:10.1016/j.icarus.2015.07.037.

Yoshikawa, I., K. Yoshioka, G. Murakami, A. Yamazaki, F. Tsuchiya, M. Kagitani, T. Sakanoi, N. Terada, T. Kimura, M. Kuwabara, K. Fujiwara, T. Hamaguchi, and H. Tadokoro (2014), Extreme Ultraviolet Radiation Measurement for Planetary Atmospheres/Magnetospheres from the Earth-Orbiting Spacecraft (Extreme Ultraviolet Spectroscopy for Exospheric Dynamics: EXCEED), *Space Sci. Rev.*, *184*, 237–258, doi:10.1007/s11214-014-0077-z.

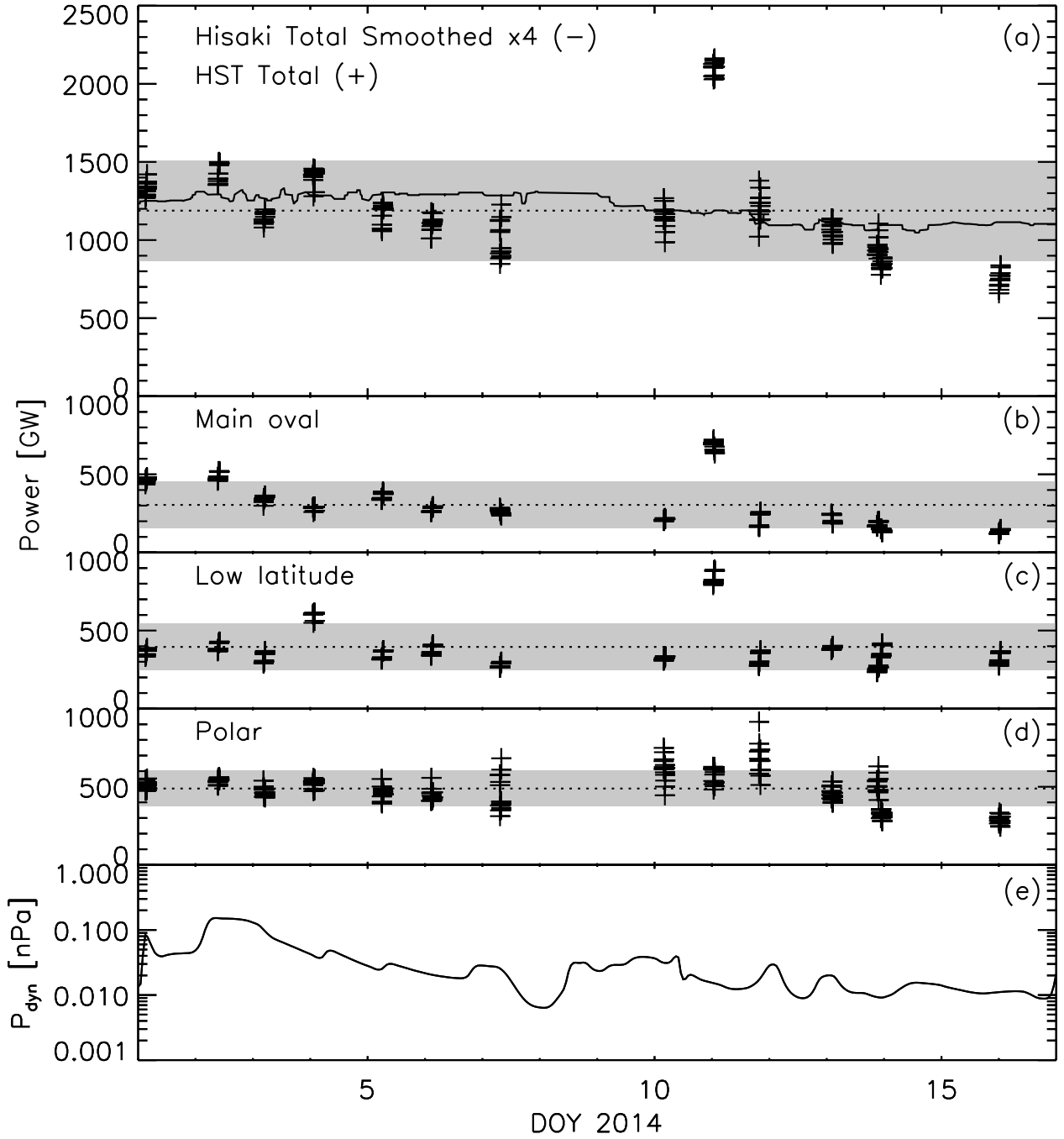


**Figure 1.** Gallery of selected images of Jupiter's northern UV aurora imaged by HST/STIS in January 2014. (a)–(n) The time and CML for each image are labelled. The images have been projected onto a polar grid at an altitude of 240 km above the 1-bar pressure level, and are viewed from above the north pole with SIII longitude  $180^\circ$  at the bottom of each panel. A latitude-longitude grid with spacing of  $10^\circ$  is superposed. The images are plotted using a log colour scale saturated at 500 kR. Red lines on (d) and (i) mark features describe in the text, while the arrowed labels on (c) indicate the Io and Ganymede footprints. (o) The average intensity derived from all images. The red contours show the boundaries of the three auroral regions: polar, main oval, and low latitude, as described in the text. (p) The location of the peak brightness at selected longitudes, tracing out the main oval, for Visits 1, 3, 13, 14 on days 1, 3, 13, 16, as labelled.

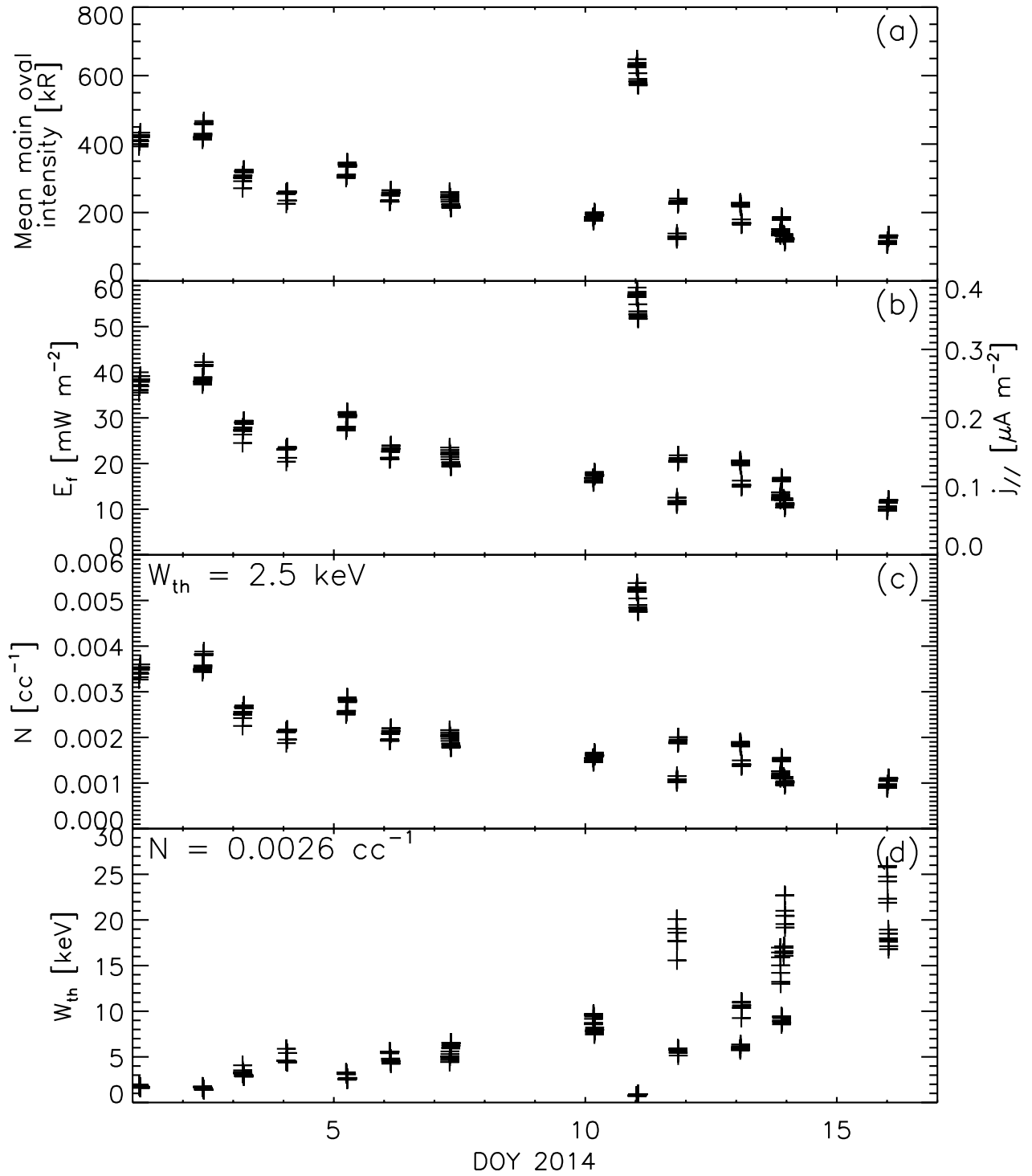
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**Figure 2.** Auroral power and solar wind dynamic pressure during 1–16 Jan 2014. (a) Total emitted FUV auroral power observed by HST/STIS (crosses), their mean (dotted line) and standard deviation about the mean (shading). The solid line shows the total EUV auroral power observed by Hisaki/EXCEED, smoothed by a running median with a window of 39.7 h (4 jovian rotations), and scaled by a factor of 4. (b)–(d) Emitted power from the main oval, low latitude and polar regions, as defined in the text. (e) Solar wind dynamic pressure at Jupiter propagated using a 1-D MHD model.



**Figure 3.** Auroral electron parameters estimated from the observations and using the linear Knight relation during 1–16 Jan 2014. (a) Mean intensity in the main oval region. (b) Incident energy flux (left hand scale) and current density (right hand scale) estimated assuming that 1 mW m<sup>-2</sup> of incident energy flux produces auroral intensity of 10 kR [Gustin *et al.*, 2012], and where the incident electron energy is taken to be 150 keV. (c) Number density of the source electrons in the equatorial middle magnetosphere assuming constant source electron temperature. (d) Temperature of the source electrons assuming constant number density.