

1 **Numerical simulations of pressure buildup and salt precipitation**
2 **during carbon dioxide storage in saline aquifers**

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10

11 **Abstract**

12 The storage of large amounts of carbon dioxide (CO₂) captured from fossil fuel fired power
13 plants in deep saline aquifers can be an effective and promising measure for reducing the
14 emissions of greenhouse gases. Massive CO₂ injection into saline aquifers may cause
15 multi-scale phenomena such as pressure buildup in a large scale, CO₂ plume evolution in a
16 medium scale and salt precipitation in a small scale. In this study, three-dimensional
17 simulations are performed to investigate the propagation of pressure and the impact of salt
18 precipitation on the process of large scale CO₂ injection into the saline aquifers. Apart from the
19 different scales of the processes, the numerical results show clearly different behaviours of the
20 pressure changes in saline aquifers with different boundaries. Different types of salt

21 precipitation occur adjacent to the injection well, presenting distinct impacts on the fluid flow.

22 Affected by salt precipitation, the porosity and permeability are reduced, leading to declined

23 transportation and degraded injectivity with different boundary conditions. The interplay

24 between pressure buildup and solid saturation is compared in saline aquifers with different

25 boundary conditions.

26

27 **Keywords:** CO₂ storage, Pressure buildup, Salt precipitation, CO₂ plume, Numerical

28 simulations

29

30 **NOMENCLATURE**

31

Symbols

d	diffusivity
D	distance between meshes m and n
g	gravitational acceleration
k	permeability tensor
k_{rg}	the relative permeability of CO_2
k_{rl}	the relative permeability of brine
n	normal vector
P	pressure
q	Darcy flux
S	saturation
t	time
T	temperature
V	volume
X	mass fraction
x, y, z	Cartesian coordinates

Greek symbols

Γ area

μ dynamic viscosity

ρ density

Σ summation

τ tortuosity

ϕ porosity

∇ gradient operator

Subscripts/superscripts

c capillary, critical

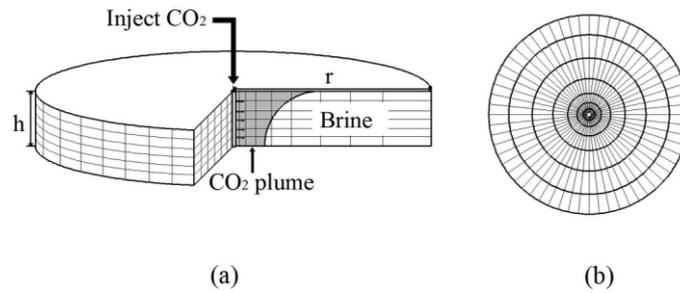
i, j, m, n index

s solid

α, β fluid phase

33 **1. Introduction**

34 Carbon dioxide storage in deep saline aquifers is potentially the most promising method for
35 massively reducing the ever increasing amount of CO₂ in the global atmospheric environment
36 because of combustion utilization of fossil fuels [1-3]. Massive CO₂ injection into the saline
37 aquifers may cause multi-scale spatial phenomena, including pressure buildup occurred in a
38 large scale [4-6], CO₂ plume in a medium size [4, 5] and the distribution of precipitation in a
39 small dimension [7]. When large volumes of CO₂ are injected into saline aquifers, pressure
40 buildup may be produced which can quickly propagate in a large space. At the temperature and
41 pressure conditions for CO₂ storage, the injected CO₂ will tend to accumulate at the top of
42 reservoir and spread out along the top caprock, as schematically shown in Fig. 1(a). Meanwhile,
43 the injection of dry supercritical CO₂ will displace the resident brine immiscibly, combined
44 with the evaporation of water, which may eventually cause the aqueous phase dry-out and salt
45 precipitation near the injection well [7-14]. The spatial size of precipitation region is just a
46 small fraction of the plume. These phenomena are of great importance to the safety of CO₂
47 storage. On the one hand, excessive pressurization may cause a series of problems, involving
48 the caprock fracture, the pollution of shallow groundwater resources, and the seismicity [15-18].
49 On the other hand, salt precipitation may lead to salt blockage near the injection well, which
50 would obstruct the transportation of CO₂ and the propagation of pressure to the far field [7, 8].
51 Therefore, predicting the propagation of pressure and the impact of salt precipitation on
52 injectivity is crucial to the security of CO₂ storage in saline aquifers.



53

54 **Fig. 1.** Schematic representation of (a) CO₂ injection into an aquifer via a vertical well and (b)
 55 top view.

56 The pressure buildup during CO₂ injection into saline aquifers has been the focus of research
 57 by a number of theoretical analyses and numerical simulations. In terms of the theoretical
 58 analyses, several simple semi-analytical methods using Buckley-Leverett equation are used to
 59 study the distribution of pressure, which describe the one-dimensional immiscible flow in the
 60 absence of compression of rock pores and brine and capillary pressure [19-21]. Mathias et al. [4]
 61 improved the Buckley-Leverett method by incorporating the compressibility of rock and brine
 62 to study the pressure buildup during CO₂ injection into a closed saline aquifer. Zhou et al. [6]
 63 developed a quick assessment method of CO₂ storage capacity due to the formation and fluid
 64 compressibility, with assumptions that pressure buildup is spatially uniform and independent of
 65 formation permeability. Although these theoretical analyses may efficiently predict the pressure
 66 changes in some cases, detailed numerical simulations of carbon storage to calculate the
 67 pressure buildup including the spatial and temporal distributions are needed. For numerical
 68 studies, the important physical phenomena of pressure buildup are observed. Nonlinear
 69 behaviours of pressure change near wellbore during CO₂ injection into saline aquifers are
 70 observed [22]. Large-scale CO₂ injection could cause groundwater pressure perturbation and
 71 hydrological impact on groundwater resources [5, 17, 23, 24]. If the pressure buildup is above a

72 threshold value, fracturing may occur. There is a stipulation by the U.S. Environment
73 Protection Agency, stating that the maximum pressure must not exceed 90% of the fracture
74 pressure in the injection zone [25]. Coupled reservoir geomechanical analyses are performed to
75 check the fracture pressures by numerical simulations [26, 27]. Numerical simulations and
76 optimization schemes are increasingly used to investigate this phenomenon, e.g. [28].
77 Optimization and parallel algorithms are also available to improve computation performance,
78 e.g. [29-32]. The previous studies indicate that the pressure buildup in the injection zone is
79 crucial to the security of CO₂ storage.

80 The process of salt precipitation has also been investigated by several theoretical analyses,
81 experimental studies and numerical simulations. For theoretical analyses, Zeidouni et al. [10]
82 developed a graphical method to determine the location of the front of solid salt. However, their
83 results neglect the effects of the capillary pressure and the gravitational force. In addition their
84 results are only applicable to a very simplified one-dimensional situation. For experimental
85 studies, the reduction of permeability induced by drying of brine in porous media is studied for
86 different rocks and salt contents [33]. A lab-on-a-chip approach is developed to study the
87 pore-scale salt precipitation dynamics during CO₂ injection into saline aquifers [34]. Although
88 experimental studies can provide first-hand results, detailed measurements are always difficult
89 especially when information on flow quantities over a broad range of time and length scales is
90 needed. In numerical studies, several researchers have shown that salt precipitates
91 preferentially near the injection well as resident saline water is evaporated by injected CO₂ [7, 8,
92 14, 35-37]. For example, Hurter et al. [35] investigated the drying out and salting out
93 phenomena using a commercial code. However, their results ignore the precipitation impact on

94 permeability. Pruess and Müller [7] carried out one- and two-dimensional studies to predict salt
95 precipitation and to understand the influencing factors for this process. Kim et al. [8] pointed
96 out that there are two types of precipitation at different injection rates using two-dimensional
97 simulations, which are characterized by different level of salt precipitation near the well. Their
98 results suggest that great pressure buildup would occur near the lower portion of the injection
99 well in some cases. These previous studies indicate that salt precipitation could cause reduction
100 of aquifer porosity and permeability near the well and thus deterioration of injectivity.

101 Although some understandings on the impacts of pressure buildup and salt precipitation of
102 CO₂ injection into the saline aquifers have been obtained, more studies are needed to
103 understand the interplay between pressure buildup and salt precipitation. In previous numerical
104 studies of salt precipitation in saline aquifers, the injection period is short and the injection rate
105 was low, which does not meet the requirements of long-term and large-scale CO₂ storage. In the
106 meantime, comparisons of the two phenomena in storage systems with different boundary
107 conditions, namely the closed, open and semi-closed systems, are important but have not been
108 investigated systematically.

109 In this study, the distributions of pressure buildup and salt precipitation, the specific
110 processes and the impacts of solid precipitation on the long-term injection in the three storage
111 systems are investigated by three-dimensional (3D) simulations. In the following, the
112 governing equations together with the initial and boundary conditions used in the simulations
113 are presented first, followed by numerical results and discussions of the results for the three
114 systems investigated. Finally, some conclusions are drawn.

115 2. Modelling and mathematical formulation

116 2.1 Physical problem and computational domain

117 The physical problem is CO₂ injection and propagation, via a vertical well, into saline
118 aquifers, as indicated in Fig. 1(a). The storage formation, located at a depth of approximate
119 1200 m below the ground surface, is 100 m thick with a radius of 40 km for the closed and
120 semi-closed systems. The lateral extent of computation model for the open system is 100 km,
121 which ensures that the lateral boundary could have a minimal effect on the simulation results.

122 2.2 Governing Equations

123 The governing equations for the fluid flows of multiphase and multicomponent fluid
124 mixtures in porous media are used to describe CO₂ geological storage in saline aquifers [3],
125 which are similar to those for oil, water, and gas flows through porous media. For isothermal
126 problems, only the mass conservation equations for CO₂, water and salt are considered. The
127 integral form of the mass equations for an individual *i*th species or component is given as [38]:

$$128 \quad \frac{\partial}{\partial t} \int_{V_n} \phi \sum_{\alpha} (\rho_{\alpha} S_{\alpha} X_i^{\alpha}) dV_n + \int_{\Gamma_n} \sum_{\alpha} (\rho_{\alpha} X_i^{\alpha} \mathbf{q}_{\alpha}) \cdot \mathbf{n} d\Gamma_n - \int_{\Gamma_n} \sum_{\alpha} (\phi S_{\alpha} \tau_{\alpha} d_i^{\alpha} \rho_{\alpha} \nabla X_i^{\alpha}) \cdot \mathbf{n} d\Gamma_n = \int_{V_n} Q_i dV_n \quad (1)$$

129 where \mathbf{n} is the normal vector on the surface element $d\Gamma_n$ (assumed pointing inward into the
130 mesh n). Eq. (1) is constructed by the balance of four terms representing all the possible
131 mechanisms for mass transfer, which are the time rate of change of mass at a fixed point (or the
132 local derivative or storage term), convective and diffusive transports, and source/sink term of
133 mass respectively. \mathbf{q}_{α} can be defined by Darcy's law [39]:

$$134 \quad \mathbf{q}_{\alpha} = - \frac{\mathbf{k} k_{r\alpha}}{\mu_{\alpha}} (\nabla P_{\alpha} + \rho_{\alpha} \mathbf{g} \nabla z) \quad (2)$$

135 Eq. (2) is a multi-phase extension of Darcy's equation. Darcy's law is an approximate form of
 136 the fluid momentum balance in creeping flow through porous media. The law is only valid for
 137 steady, slow viscous flow, which can be derived from the Navier–Stokes momentum equations.

138 Eqs. (1)–(2) constitute the fundamental governing equations for the numerical simulations
 139 studied here. They are a coupled nonlinear system involving the geo-mechanical effects such as
 140 permeability and porosity of the solid rock matrix, multi-phase fluid properties like density and
 141 viscosity, which all affect the flow and transport behaviours. In order to close this mathematical
 142 problem, constitutive relationships and supplementary constraints for saturations, component
 143 compositions and pressures are needed [3].

144 The relative permeability $k_{r\alpha}$ is the ratio of the α phase permeability to the permeability of the
 145 porous medium. Under all-gas condition, the relative permeability of CO₂ is equal to 1.0. In
 146 order to close Eqs. (1)–(2), relationships for the relative permeability and capillary pressure are
 147 needed. In general, the two-phase characteristic curves are a function of the pore structure,
 148 phase saturation, surface tension, contact angle, and hysteresis [38].

149 The relative permeabilities of brine and CO₂ are calculated as follows:

$$150 \quad k_{rl} = \sqrt{S^*} \left\{ 1 - \left(1 - [S^*]^{1/\lambda} \right)^\lambda \right\}^2 \quad (3)$$

$$151 \quad k_{rg} = (1 - \hat{S})^2 (1 - \hat{S}^2) \quad (4)$$

$$152 \quad S^* = (S_l - S_{lr}) / (1 - S_{lr}) \quad (5)$$

$$153 \quad \hat{S} = (S_l - S_{lr}) / (1 - S_{lr} - S_{gr}) \quad (6)$$

154 where k_{rl} and k_{rg} are the liquid and gas relative permeabilities, respectively. S_l is the liquid

155 saturation, while S_{lr} and S_{gr} are the irreducible liquid and gas saturations, respectively. Eq. (3)

156 for liquid is developed by van Genuchten [40]; eq. (4) for gas is due to Corey [41].

157 The formation for capillary pressure is given by van Genuchten [40]:

158
$$P_{c,\alpha\beta} = -P_0 \left([S^*]^{-1/\lambda} - 1 \right)^{1-\lambda} \quad (7)$$

159 where P_0 is the strength coefficient, and λ is a parameter depending on pore geometry.

160 The difference of pressures between the two phases satisfies the following relation [39]:

161
$$P_\beta = P_\alpha + P_{c,\alpha\beta} \quad (8)$$

162 Eq. (8) shows that the fluid pressure in phase β is the sum of the gas phase pressure P_α and the
163 capillary pressure $P_{c,\alpha\beta}$.

164 The evaporation model for H₂O partitioning into CO₂-rich phase is given by Spycher and
165 Pruess model [42], which gives the mutual solubilities of CO₂ and H₂O in a non-iterative
166 manner.

167 The salt precipitation due to the evaporation of injected CO₂ affects the fluid flows of gas and
168 aqueous phases by changing the porosity and permeability of the formations. The solid salt
169 occupies a fraction of the volume of the pores, which will lead to the decrease of space available
170 for gas and aqueous phases. In this study, the solid salt is assumed to be immobile. Similar to
171 the saturations of gas and aqueous phases, solid saturation is defined to describe the fraction of
172 pore space occupied by salt precipitation.

173 In modelling the interplay between the two-phase flow and salt precipitation, it is important
174 to specify the relationship between porosity and permeability. The underground formations

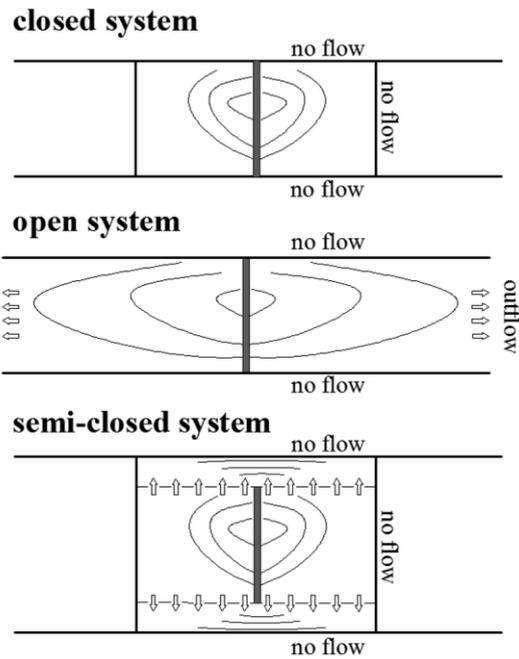
175 contain different sizes of pores. Some precipitation can occur in the large pores, in which the
176 permeability may not change much; others are found in the small pores, in which the
177 permeability may decrease dramatically. The porosity-permeability relationship has been
178 discussed by many investigators [14, 43, 44], whose results differ considerably from each other
179 due to the complexity of the problem.

180 A tubes-in-series model is used to describe the permeability change due to the solid
181 precipitation [44]. The model is composed of a series of parallel tubes with larger and smaller
182 radii. The axes of the tubes are parallel to the fluid flows. The flow channels contain a great
183 number of pore throats, hence even small changes in porosity may lead to dramatic
184 permeability change due to the blockage of the pore throats. This permeability may be reduced
185 to zero at a finite porosity, which can be defined as the “critical porosity”. In this study, the
186 permeability decreases to zero when the porosity is reduced to 80% of its original value, i.e.,
187 when the solid saturation reaches 0.20.

188 **2.3 Boundary and initial conditions**

189 In terms of physical boundaries, the storage systems can be theoretically divided into three
190 categories: (i) a closed system in which all the boundaries are impervious; (ii) an open system
191 whose lateral boundaries are open so that the native brine can flow out; and (iii) a semi-closed
192 system in which the lateral boundaries are impervious, while the storage formation is vertically
193 bounded by sealing units with low permeability [6, 15]. For a closed system, the storage
194 depends on the compressibility of the formation fluids and rock material as well as the
195 dissolution rate of CO₂, which can provide expanded volumes available for storing the injected
196 CO₂ [4, 16]. For an open system, the injected CO₂ displaces the brine laterally and is stored in

197 the space that filled with aqueous phase [4-6]. For a semi-closed system, some fraction of the
 198 brine in the storage formation can migrate into the sealing units, which will increase the storage
 199 capacity for the injected CO₂ [6, 17, 18].



200

201 **Fig. 2.** Schematic representation of boundary conditions for the three storage systems: (a) open
 202 system, (b) closed system, and (c) semi-closed system.

203 All the boundaries for the closed and semi-closed systems are assumed to be impermeable to
 204 both supercritical CO₂ and brine except the wellbore boundary. For the open system, the
 205 volumes of grid blocks at the lateral boundary are assigned with an extremely large numerical
 206 value of 10⁵⁰ m³, thereby imposing a constant pressure condition at the far field. The top and
 207 bottom boundaries are also impervious. For the semi-closed system, two sealing formations
 208 with 60 m thick each are located at the top and the bottom of the storage system. The boundary
 209 conditions for the three storage systems are shown in Fig. 2.

210 **Tab. 1.** Hydrogeological properties of the storage formation.

Initial conditions

Temperature	$T = 45^{\circ}\text{C}$
Salinity	$X_s = 0.15$
Pressure	$P_{ini} \approx 120\text{-}131 \text{ bars}$
Dissolved CO ₂ concentration	$X_1 = 0.$

Formation properties

Horizontal permeability	$k_h = 10^{-12} \text{ m}^2$
Vertical permeability	$k_v = 10^{-12} \text{ m}^2$
Porosity	$\phi = 0.12$
Pore compressibility	$D = 4.5 \times 10^{-10} \text{ Pa}^{-1}$

211 Tab. 1 lists the assigned values of parameters used in this study, which are the typical
212 conditions suitable for CO₂ storage. The formations are initially fully brine-saturated with the
213 hydrostatic pressure distributing over the depths of the formations. The injection rate and
214 injection period in the three systems are the same, which are 100 kg/s and 30 years, respectively.
215 In order to examine the effect of the injection rate, a rate of 50 kg/s with injection period of 60
216 years is also considered for the closed system. Temperature is fixed at 45 °C throughout the
217 simulations, representing an isothermal condition for the simulations considered here.

218 2.4 Numerical methods

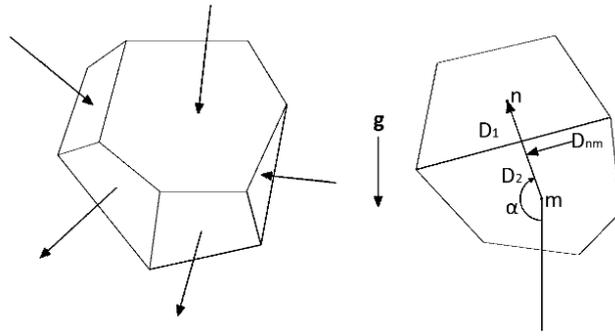
219 The mass equations are discretized temporally using an implicit finite difference scheme and

220 in space using an integral finite difference method as follows:

$$\begin{aligned}
 & \left[\phi_n^{t+\Delta t} \sum_{\alpha} \left(\rho_{n,\alpha}^{t+\Delta t} S_{n,\alpha}^{t+\Delta t} X_{n,i}^{\alpha,t+\Delta t} \right) - \phi_n^t \sum_{\alpha} \left(\rho_{n,\alpha}^t S_{n,\alpha}^t X_{n,i}^{\alpha,t} \right) \right] \cdot \frac{V_n}{\Delta t} + \\
 221 \quad & \sum_m \left\{ \left[\sum_{\alpha} \left(\rho_{\alpha,nm}^{t+\Delta t} X_{i,nm}^{\alpha,t+\Delta t} q_{\alpha,nm}^{t+\Delta t} \right) \right] \cdot \Gamma_{nm} \right\} - \\
 & \sum_m \left\{ \left[\sum_{\alpha} \left[\phi_{nm}^{t+\Delta t} S_{\alpha,nm}^{t+\Delta t} \tau_{\alpha,nm}^{t+\Delta t} d_{i,nm}^{t+\Delta t} \rho_{\alpha,nm}^{t+\Delta t} \left(X_{i,n}^{\alpha,t+\Delta t} - X_{i,m}^{\alpha,t+\Delta t} \right) / D_{nm} \right] \right] \cdot \Gamma_{nm} \right\} \\
 & = Q_{n,i}^t \cdot V_n
 \end{aligned} \tag{9}$$

222 where $t + \Delta t$ represents the new time step, and flux terms are treated as fully implicit, given by

223 the values at the new time step.



224

225 **Fig. 3.** Spatial discretization considered in this study.

226 The Darcy's law is discretized in the following way:

$$227 \quad q_{nm}^{t+\Delta t} = - \left(\frac{k}{\mu} \right)_{nm}^{t+\Delta t} \left(\frac{P_n^{t+\Delta t} - P_m^{t+\Delta t}}{D_{nm}} - \rho_{nm}^{t+\Delta t} g \cos \alpha \right) \tag{10}$$

228 where α is the intersection angle between gravitational acceleration and the line segment from

229 mesh m to n with rotation direction from g to the line segment clockwise as indicated in Fig. 3.

$$230 \quad \cos \alpha = \frac{Z_2 - Z_1}{D_1 + D_2} \tag{11}$$

231 The variables in Eq. (9) - (10) on the interface are treated by distances harmonic averages

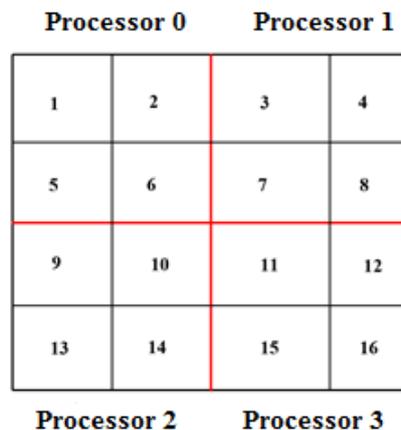
232 method, given by,

233
$$f_{\alpha, nm}^{t+\Delta t} = \frac{D_1 f_{\alpha, m}^{t+\Delta t} + D_2 f_{\alpha, n}^{t+\Delta t}}{D_1 + D_2} \quad (12)$$

234
$$f = \{ \phi, S_\alpha, \tau_\alpha, d_i^\alpha, \rho_\alpha, h_\alpha, \lambda \} \quad (13)$$

235 A set of coupled nonlinear equations are obtained from Eq. (9)-(10). The compressed sparse
 236 row (CSR) format is adopted to store the sparse matrix linearized by the Newton-Raphson
 237 iteration [28, 45]. Nonzero elements of the matrix are stored in CSR format. Afterwards, the
 238 obtained system of linear equations is solved by parallel algorithm. In order to perform parallel
 239 simulations, domain decomposition method is used. The computational domain is decomposed
 240 into a number of subdomains. A global solution is formed through the local solutions on the
 241 subdomains. Solutions for subdomains can be sought simultaneously. In order to achieve better
 242 computational performance, each processor is assigned to the roughly the same number of
 243 meshes.

244 In order to track the process accurately and effectively, the temporal differencing is based on
 245 an automatic scheme, by changing the time steps according to the variations of solutions
 246 between adjacent time steps.



247

248

Fig. 4. A 16-meshes domain partitioning on 4 processors.

249

Fig. 4 shows a scheme for partitioning a sample domain with 16 meshes into four parts. Grids

250

are assigned to four different processors and reordered to a local index ordering at each

251

processor. The partitioned meshes are stored in each processor's update set. The update set is

252

further divided into two subsets: internal and border. The solutions of elements in the internal

253

subset only use the information on the current processor. The border subset includes grids that

254

would require values from the other processors to be updated. An external set stores the meshes

255

that are not in the current processor, which are needed to update the grids in the border set. Tab.

256

2 shows an example of the domain partitioning and local numbering.

257

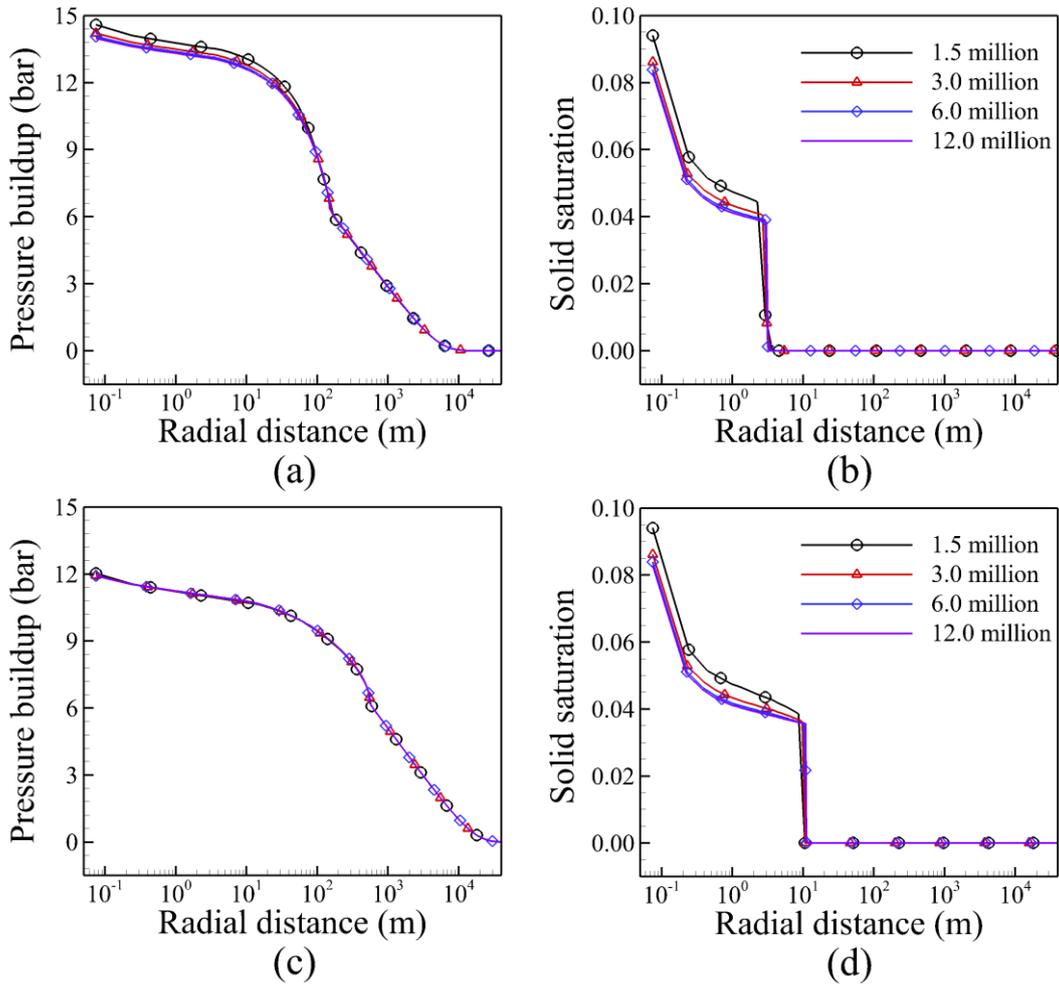
Tab. 2. Example of domain partitioning and local numbering.

Processor		update		
		internal	border	external
Processor 0	Mesh	1	2, 5, 6	3, 7, 9, 10
	Local Numbering	1	2, 3, 4	5, 6, 7, 8
Processor 1	Mesh	4	3, 7, 8	2, 6, 11, 12
	Local Numbering	1	2, 3, 4	5, 6, 7, 8
Processor 2	Mesh	13	9, 10, 14	5, 6, 11, 15
	Local Numbering	1	2, 3, 4	5, 6, 7, 8
Processor 3	Mesh	16	11, 12, 15	7, 8, 10, 14

258 Communication between processors is an essential task of the parallel algorithm. Global
259 communication is used to contribute grid blocks to all processors and check the convergence. In
260 order to solve the linear equation system, communications between adjacent processors and
261 linear solver routine are needed. When the meshes are in the border subset, exchange of data
262 corresponding to the external set is performed.

263 **2.5 Grid dependence tests**

264 In order to obtain a better understanding on how the grid resolution affects numerical
265 solutions, grid dependence is examined for the 3D closed system. Four different sets of grids in
266 the range of 1.5–12 million are used to evaluate the dependence of the results on the grid
267 number and determine the optimum number of grids, as shown in Fig. 5. The plots show the
268 radial distributions of pressure buildup (compared with the initial pressure) and solid saturation
269 at 10 days and 100 days along the top aquifer. It is evident that the optimal number of grids is 6
270 million by considering the computational accuracy and the efficiency.



271

272 **Fig. 5.** The radial profiles of pressure buildup and solid saturation at the aquifer top for 10 days

273 (top) and 100 days (bottom) for the grid dependence tests.

274 In this meshing system, the computation domain is discretized into 1000 grids in the radial

275 direction, 60 grids in the axial direction and 100 grids in the vertical direction. The grid size

276 increases logarithmically from the injection well, with the finest grid located close to the

277 wellbore and the coarsest at the far side boundary in the radial direction. Every circle of the

278 mesh in the axial direction is divided uniformly, and the targeted formation is also divided

279 uniformly in the vertical direction. Similar meshing methods for the open and semi-closed

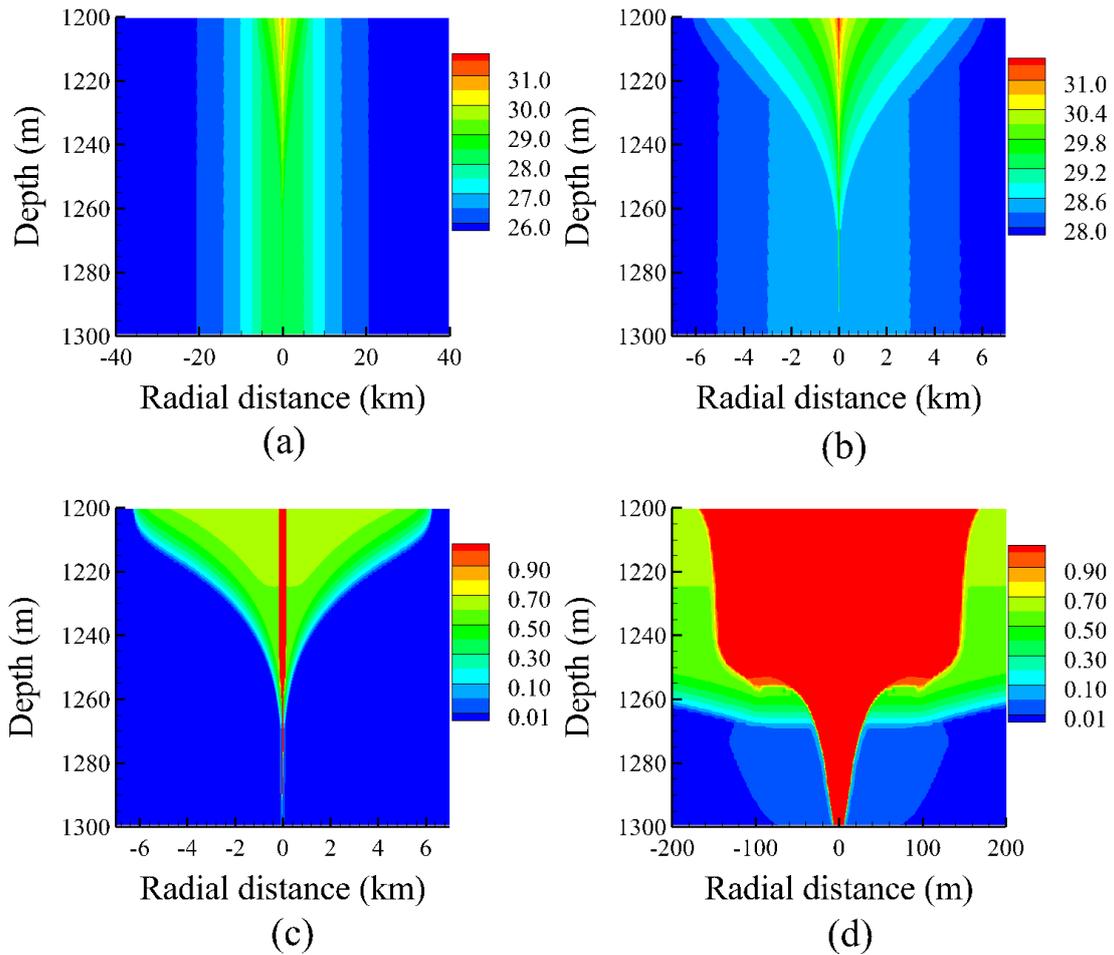
280 systems are adopted, except that the target formation for the semi-closed system is divided into

281 220 grid blocks in the vertical direction.

282 **3. Results and discussion**

283 **3.1 The results of the closed system**

284 The snapshots shown in Fig. 6 correspond to the cross sections of pressure buildup and gas
285 saturation at the end of the 30-year injection period. When large volumes of CO₂ are injected
286 into this system, a significant pressure buildup is produced. The range of pressure perturbation
287 covers the whole domain, with an elevated pressure of 31.5 bars near the injection well and of
288 26.0 bars at the lateral boundary shown in Fig. 6(a). The radius of CO₂ plume region is about 6
289 km and the plume is concentrated at the top portion of the aquifer, as shown in Fig. 6(c). It is
290 clear that the scale of elevated pressure is much larger than the CO₂ plume size. The contour
291 lines of pressure buildup in the CO₂ plume region shown in Fig. 6(b) are inclined, caused by the
292 buoyancy and nonlinearity inherent in the two-phase flow system [3]. Meanwhile the contour
293 lines away from the CO₂ plume region are mostly vertical, indicating a horizontal brine
294 displacement.



295

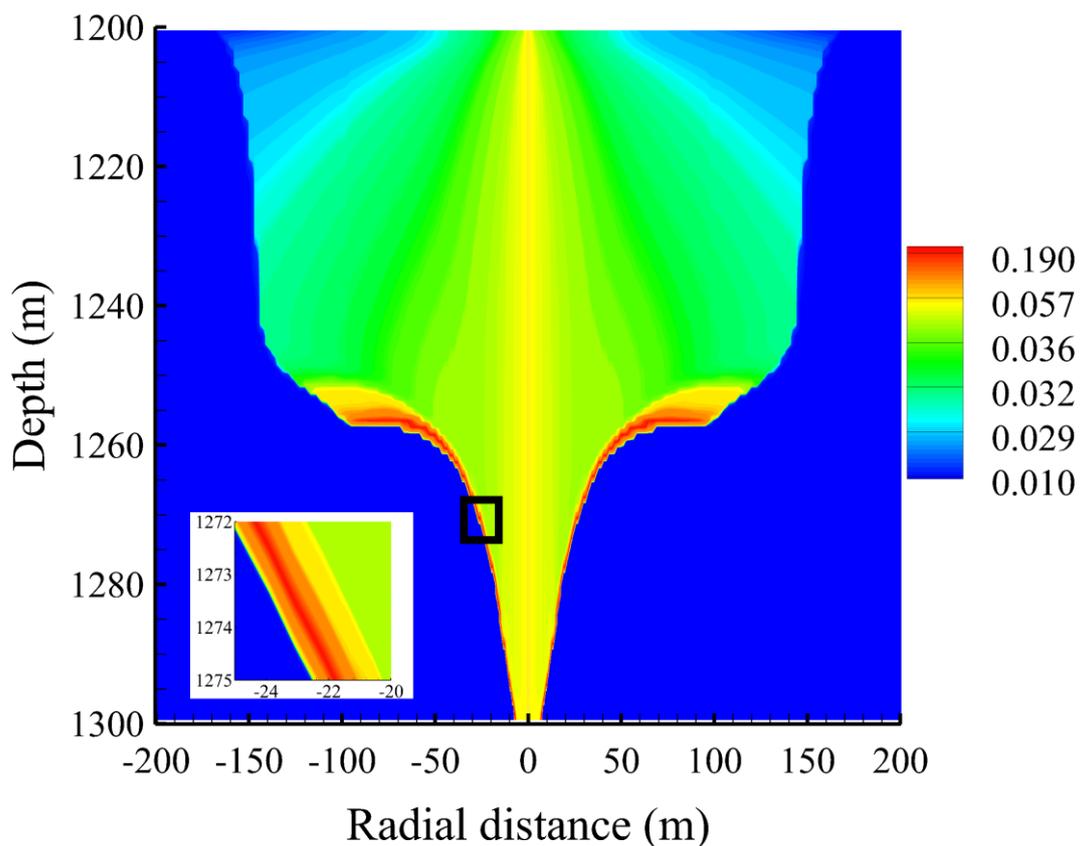
296 **Fig. 6.** Cross sections of pressure buildup (top: (b) is a zoom-in of (a)) and gas saturation

297 (bottom: (d) is the zoom-in of (c)) for the closed system at 30 years of CO₂ injection;

298 pressure unit: bar.

299 Due to the evaporation of dry gas, salt precipitation occurs near the well. In order to better
 300 capture the dynamic behaviours of precipitation, the horizontal grid size is set to be 0.15 m near
 301 the well, and increases logarithmically from the injection well. For the domain shown in Fig. 7,
 302 in which salt precipitation takes place, there are 257 grids along the horizontal direction. The
 303 precipitation distribution is controlled by the buoyancy driven CO₂ plume, which presents two
 304 kinds of precipitation, i.e., non-localized salt precipitation with smaller values and localized
 305 salt precipitation with larger values (shown in the closed-up view). Compared with the contour

306 map of gas saturation near the well in Fig. 6(d), the non-localized precipitation occurs inside the
307 zone of single gas phase and the localized precipitation is located at the lower portion of the
308 dry-out front. The highest solid saturation in the localized salt precipitation region amounts to
309 0.20, which results in a zero permeability.

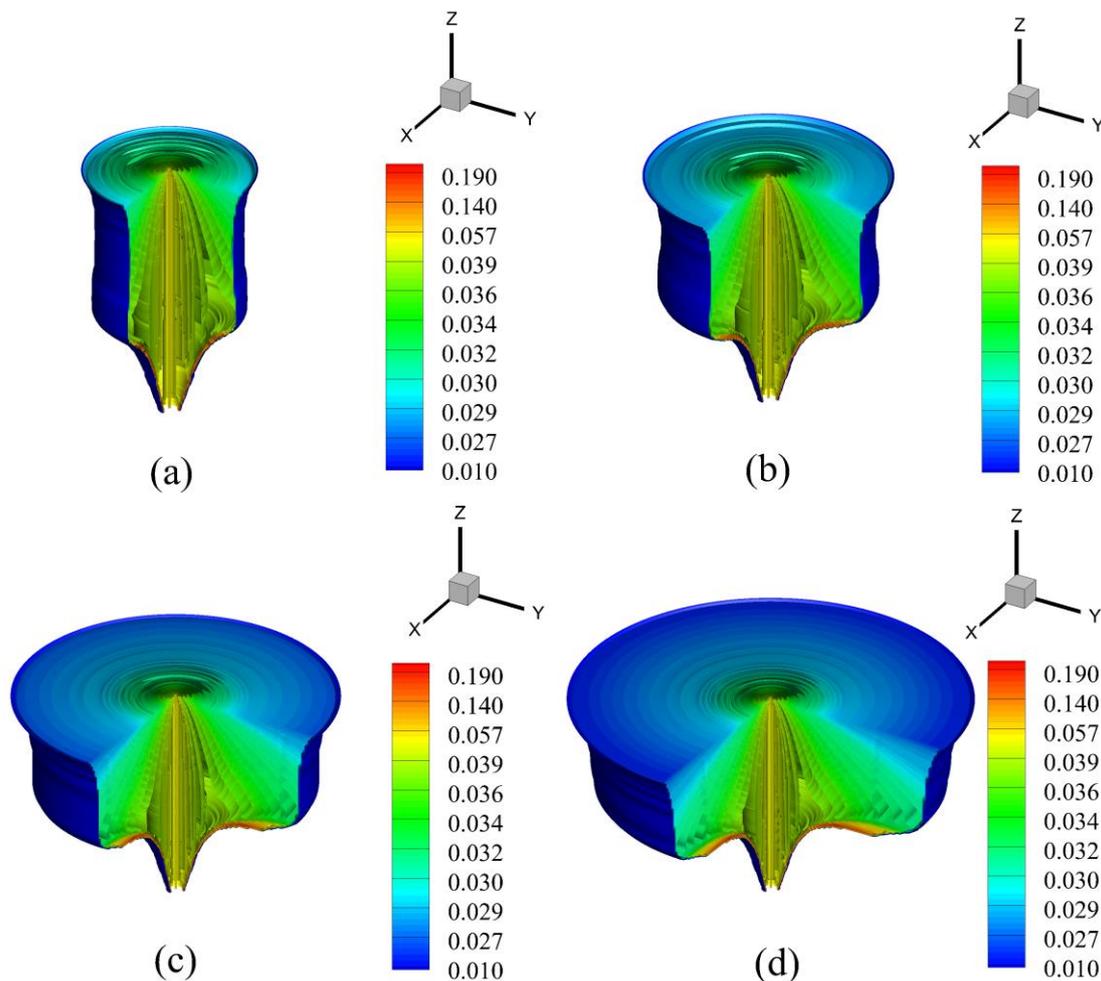


310

311 **Fig. 7.** Cross sections of solid saturation for the closed system at 30 years of CO₂ injection with
312 a close-up view of non-localized precipitation.

313 The solid saturation iso-surfaces at different time instants are shown in Fig. 8, where the 3D
314 results are shown for a three-quarter of the computational domain. The precipitation begins
315 from the injection well and develops with time. Different zones of solid saturation present
316 different behaviours with time, that is to say, the upper zone evolves continuously, while the

317 lower zone tends to be stabilized.

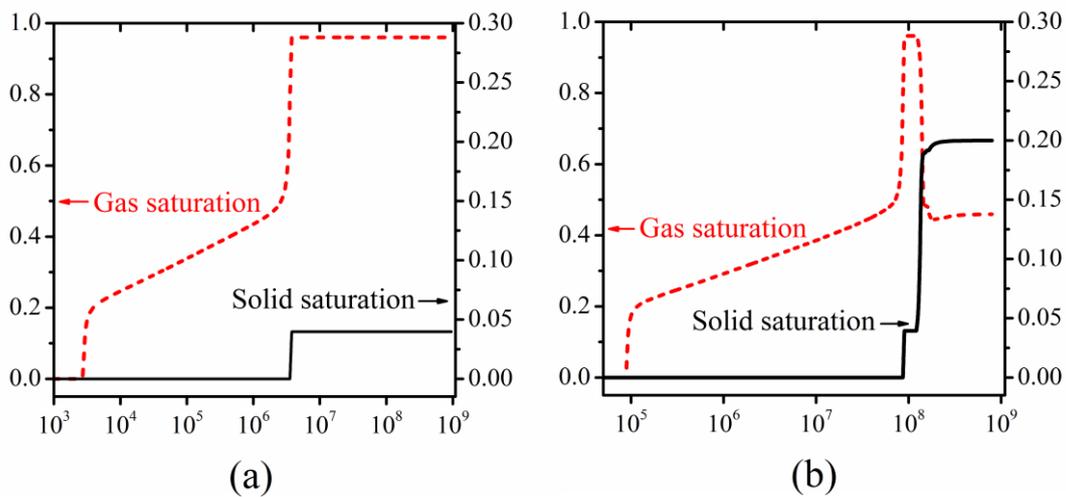


318

319 **Fig. 8.** Instantaneous iso-surfaces of solid saturation for the 3D closed system at different time
320 instants.

321 Fig. 9 shows the temporal evolution of gas saturation and solid saturation, to illustrate the
322 processes of the two types of precipitation. At the early stage, the injected CO_2 mainly displaces
323 the resident brine, accompanied by interphase mass transfer of both CO_2 and brine between the
324 aqueous phase and gas phase. When the brine becomes fully saturated due to the evaporation,
325 the salt can quickly precipitate, corresponding to the quick increase of solid saturation. These
326 trends for the two variables stop for the non-localized precipitation in Fig. 9(a). However, these
327 trends still continue for the localized precipitation in Fig. 9(b). The capillary pressure

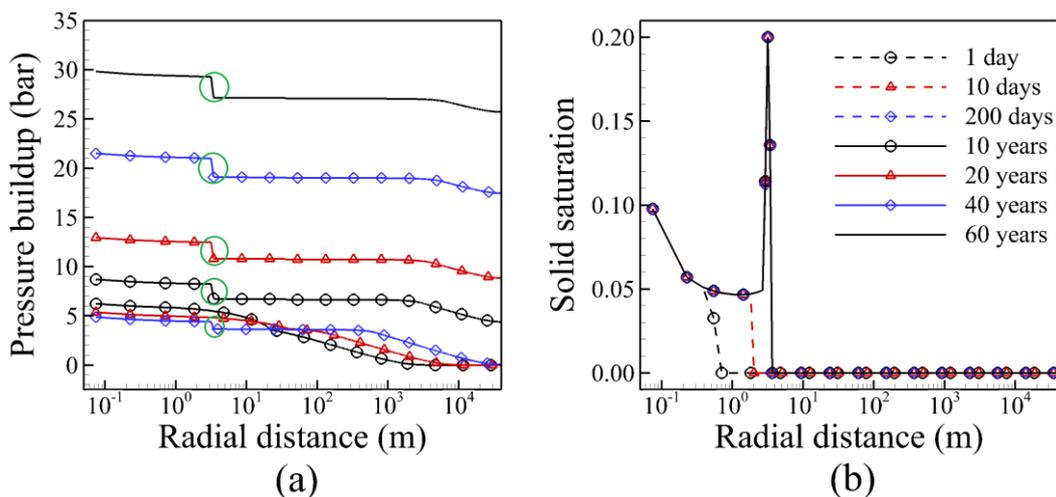
328 overcomes the injection pressure, driving the brine towards the evaporation front. The backflow
 329 of aqueous phase can increase the solid saturation and decrease the gas saturation. Under the
 330 evaporation of gas phase, the precipitation front becomes thicker and more spread out,
 331 representing the increase of solid saturation. Once the solid saturation reaches 0.20, the
 332 composition of phases will not be changed.



333 (a) (b)
 334 **Fig. 9.** The temporal evolution of gas saturation and solid saturation in (a) the non-localized
 335 precipitation region, (b) the localized precipitation region.

336 The pressure buildup along the bottom aquifer in Fig. 10(a) shows complicated behaviours.
 337 Simulation results predict an initial jump followed by a quick decline and then a gradual
 338 increase in near wellbore pressure over time. The pressure at these locations away from the
 339 injection well increases monotonously with time. Most notably, the curves demonstrate a
 340 pressure jump in the position of 8 m after an injection period of five years, and the values of
 341 pressure jump increase with the injection time, as marked by the green circles in Fig. 10(a). The
 342 differences of pressure are 0.38 bar for 5 years, 0.50 bar for 10 years, 0.70 bar for 20 years, and
 343 0.89 bar for 30 years, which show an approximately linear behaviour.

344 The zone of solid precipitation in Fig. 10(b) spreads with the injection time. At the early
 345 stage (less than 100 days), the solid saturation zone spreads with time. As the brine is displaced
 346 gradually by the injected CO₂, the amount of precipitable salt declines with the increasing
 347 distance from the injection well, which in turn leads to the decrease of solid saturation. After 1
 348 year injection, the backflow of brine occurs, resulting in a sharp gradient of solid saturation.
 349 Compared with the results in Fig. 10(a), the location of the gradient of solid saturation
 350 corresponds to the location of pressure jump. When the solid saturation amounts to 0.20, the
 351 pores are clogged completely and the horizontal flows of gas and aqueous phase are suppressed.
 352 During the subsequent stages, the profiles of gas and solid saturations remain unchanged.



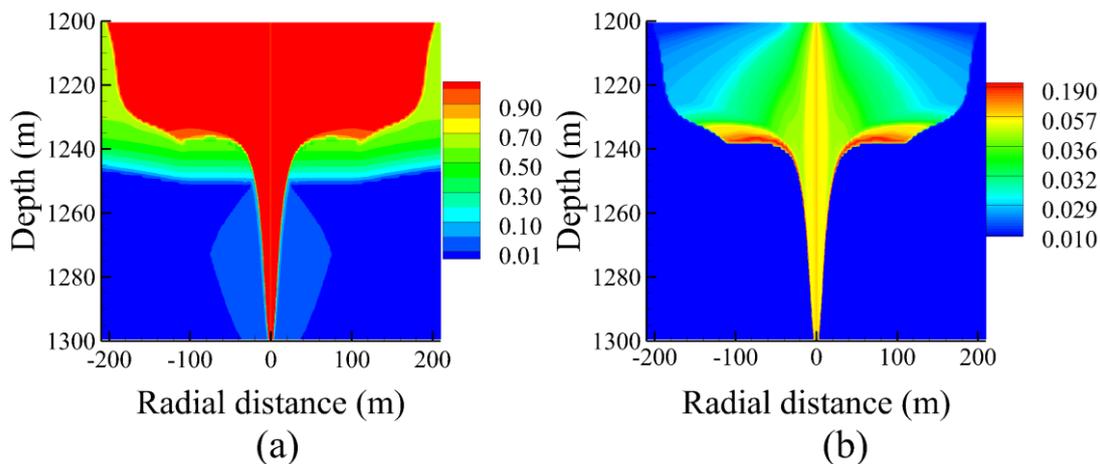
353 (a) (b)

354 **Fig. 10.** Profiles along the bottom of the aquifer for the closed system at different injection time
 355 instants for (a) pressure buildup and (b) solid saturation.

356 The results of lower injection rate of 50 kg/s with the same total amount of CO₂ are given in
 357 Fig. 11-12. The values of hydrogeological parameters used in this model are given in Tab 1.

358 Fig. 11 shows the cross sections of gas and solid saturations with injection rate of 50 kg/s at
 359 the end of the injection period of 60 years. Compared with the larger injection rate case, there

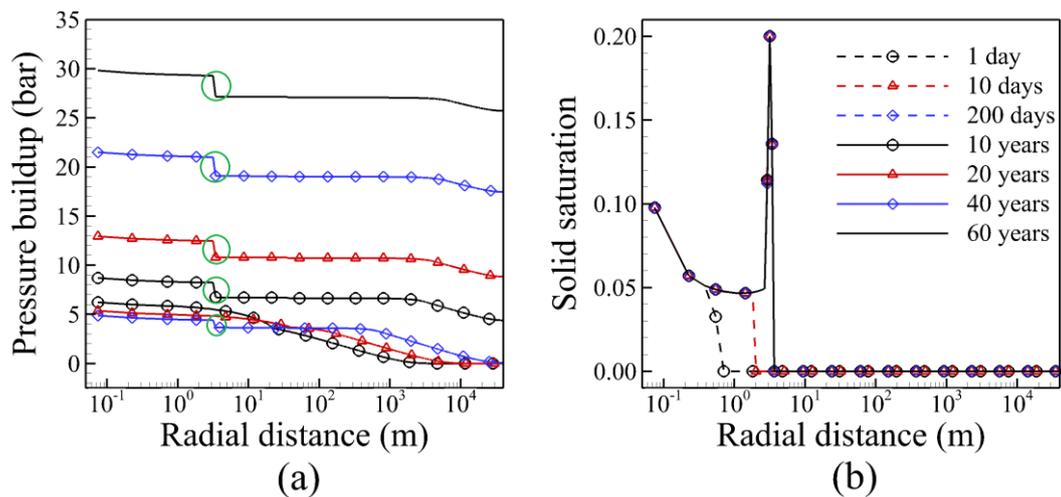
360 are obvious differences in the shapes of these distributions. The horizontal spread of the gas
 361 phase is reduced while the vertical movement is enhanced. Similar to the larger-rate case, the
 362 solid salt appears in the zone of single gas phase. Rather different precipitation behaviours are
 363 observed at the lower rate. The distribution radii of solid precipitation zone are smaller near the
 364 lower portions of the well, while the radii are larger near the upper portions of the well. The
 365 accumulation of solids in this case exacerbates gravity override effect, which means that more
 366 gas phase accumulates at the top aquifer. The distribution of gas phase increases the risk of
 367 leakage and reduces the security of CO₂ storage. The narrower zone at the bottom of the aquifer
 368 attenuates the pressure jump, which can be clearly seen in Fig. 12(a).



369
 370 **Fig. 11.** Spatial distributions for the closed system at 60 years of CO₂ injection of (a) gas
 371 saturation and (b) solid saturation.

372 Fig. 12 shows the cross sections of pressure buildup and solid saturation at different time
 373 instants. Compared with the larger-rate case, the increment of pressure is slightly lower, while
 374 the values of pressure jump are higher at the bottom of the aquifer, as marked by the green
 375 circles in Fig. 12(a). The differences of pressure are 0.73 bar for 200 days, 1.50 bars for 10 years,
 376 1.66 bars for 20 years, 1.94 bars for 40 years, and 2.15 bars for 60 years, respectively. The value

377 of solid saturation at the injection well is 0.0975, which is 16 percent higher than that in Fig.
 378 10(b). The distance between the impervious barrier and the wellbore is shorter, which is only
 379 3.2 m. All of these factors could increase the possibility of fracture near the lower portion of
 380 injection well.

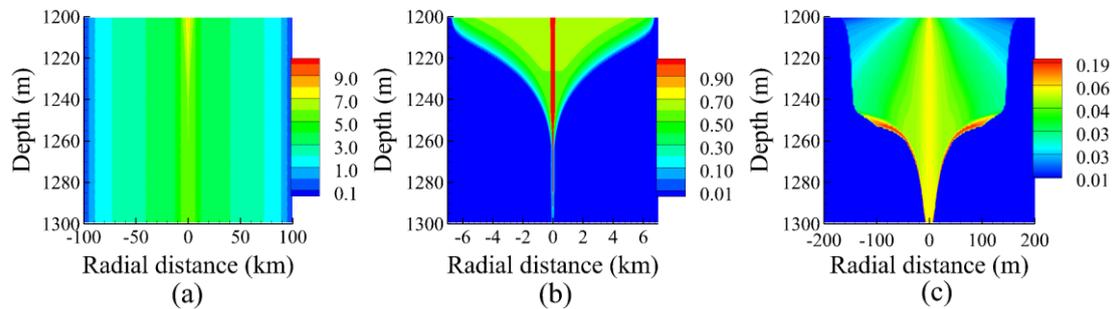


381 (a) 382 **Fig. 12.** Profiles along the bottom of the aquifer for the closed system with injection rate of 50
 383 kg/s at different injection time instants for (a) pressure buildup, and (b) solid saturation.

384 3.2 The results of the open system

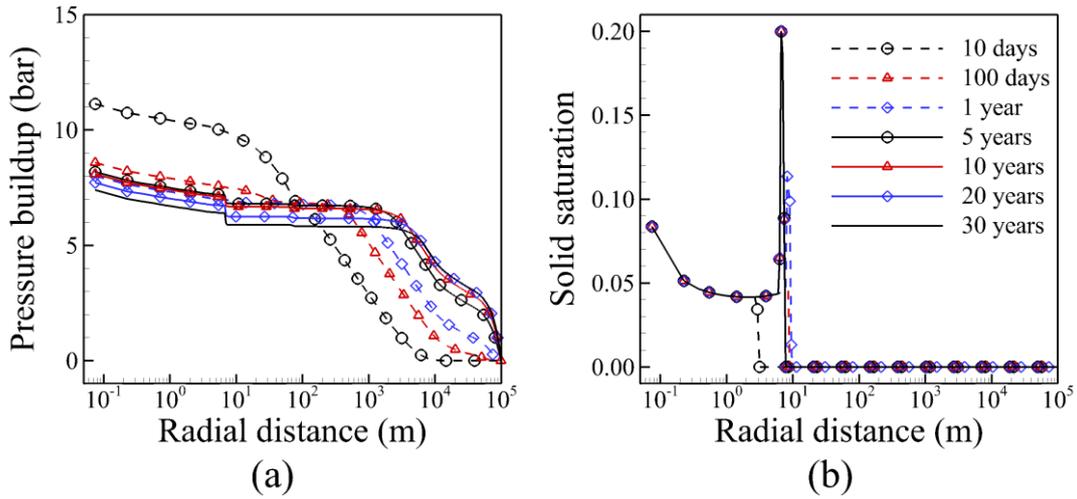
385 The snapshots shown in Fig. 13 correspond to the cross sections of pressure buildup, gas and
 386 solid saturations for the open system at the end of the 30-year injection period. Compared with
 387 the results in the closed system, a significant difference in the contour maps of pressure buildup
 388 is observed. The values of pressure buildup are lower, with maximum value of 9.5 bars at the
 389 top of the injection well. In marked contrast to the difference in the distribution of pressure
 390 buildup, minor differences in the CO₂ plumes and solid saturation distributions are observed.
 391 Comparison of the results in the closed and open systems indicates that the shapes of gas and
 392 solid phase distributions for the two storage systems are generally similar, with a larger distance

393 in the lateral extent of the plume for the open system. The differences in the CO₂ plumes are
394 caused by the differences in pressure buildup.



395 (a) (b) (c)
396 **Fig. 13.** Cross sections of (a) pressure buildup (unit: bar), (b) gas saturation and (c) solid
397 saturation for the open system at 30 years of CO₂ injection.

398 Fig. 14 shows the radial profiles of pressure buildup and solid saturation at the same time
399 instants as the closed system throughout the injection period. The profiles of pressure buildup
400 show different behaviours compared with those in Fig. 10. Simulation results predict an initial
401 jump followed by the continuous decline in the pressure near wellbore over time, while the
402 pressures in the other region increase slightly. As the outflow rates of brine at the lateral
403 boundaries are constant, the pressure changes in the whole domain are not obvious. Meanwhile,
404 the pressure profiles along the bottom aquifer also present a jump near the well due to the
405 localized precipitation. The radial profiles of solid saturation in the closed and open systems are
406 generally similar, with minor differences in the radial distance of solid saturation along the
407 bottom surface of the aquifer.



408

409 **Fig. 14.** Profiles along the bottom of the aquifer for the open system at different injection time

410 instants for (a) pressure buildup (unit: bar), and (b) solid saturation.

411 3.3 The results of the semi-closed system

412 Fig. 15 shows the cross sections of pressure buildup with seal permeabilities of 10^{-17} m^2 , 10^{-18}

413 m^2 and 10^{-19} m^2 at the end of the 30-year injection period. In these cases, a small fraction of the

414 brine in the storage formation is displaced into the overlying and underlying formations during

415 the injection period, which can provide additional storage space for CO_2 . Hence less pressure

416 buildup occurs in the semi-closed system compared with the results in Fig. 6(a). The pressure

417 buildup in the storage formations is very sensitive to the seal permeability. In the lowest seal

418 permeability (10^{-19} m^2) case, the pressure buildup shows similar behaviours to those in the

419 closed system. The propagation of elevated pressure is mainly in the storage formation. The

420 values of pressure buildup in the storage formations are much higher than the values in the seal

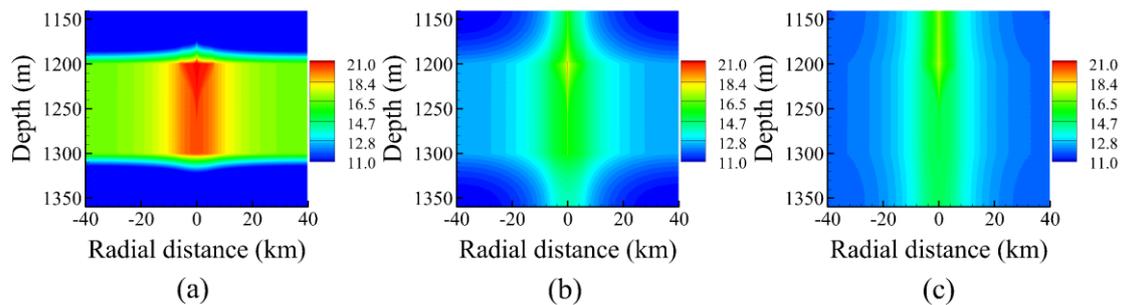
421 formations. In the medium seal permeability (10^{-18} m^2) case, the elevated pressure in the storage

422 saline is lower than that in the lowest seal permeability case. More native brine in the storage

423 formation is discharged into the seal formations. In the largest seal permeability (10^{-17} m^2) case,

424 the propagation of elevated pressure is dominant in the vertical direction. With the increase of

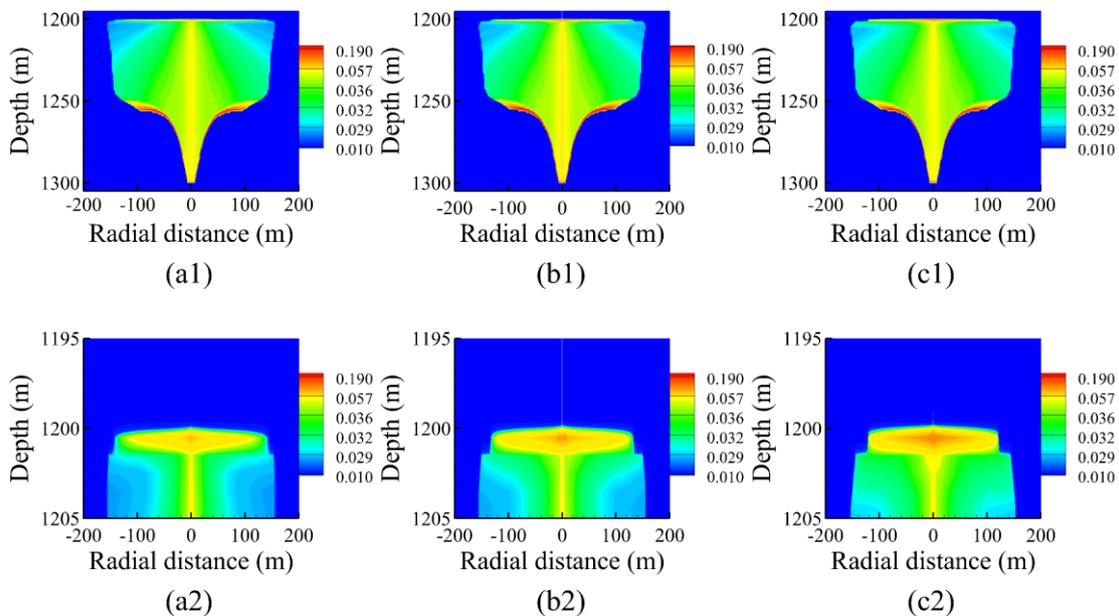
425 the fraction of brine leakage into the seal formation, more space is provided for the injected CO₂
 426 in the storage formation. The zones of higher pressure buildup are all located in the two-phase
 427 regions for the three cases, which are the same as the distributions in the closed and open
 428 systems.



429 (a) (b) (c)
 430 **Fig. 15.** Cross sections of pressure buildup (unit: bar) with seal permeabilities of (a) 10^{-19} m^2 , (b)
 431 10^{-18} m^2 , and (c) 10^{-17} m^2 for the semi-closed system.

432 Fig. 16 shows the cross sections of solid saturation with three different seal permeabilities,
 433 where (a2-c2) are the zoom-in graphs of (a1-c1). In contrast to the distribution of pressure
 434 buildup, the solid saturation is less sensitive to the seal permeability. Comparison of Fig. 16
 435 (a1-c1) indicates that the contour maps of solid precipitation in all the semi-closed cases are
 436 generally similar in shape, with several minor differences at the top of the storage formations.
 437 In addition to the two types of precipitation near the injection well (i.e., non-localized salt
 438 precipitation of smaller values, localized salt precipitation of larger values), the third type of
 439 solid precipitation occurs at the interfaces between the storage formation and the seal
 440 formations. At the interfaces, injected CO₂ in the storage saline hardly enters into the seal
 441 formations, which needs to overcome a considerable capillary entry pressure. Consequently, the
 442 flow of single gas phase at the interfaces can be suppressed and thus more salt can precipitate at
 443 these locations. With the increase of seal permeability, both the thickness of this kind of

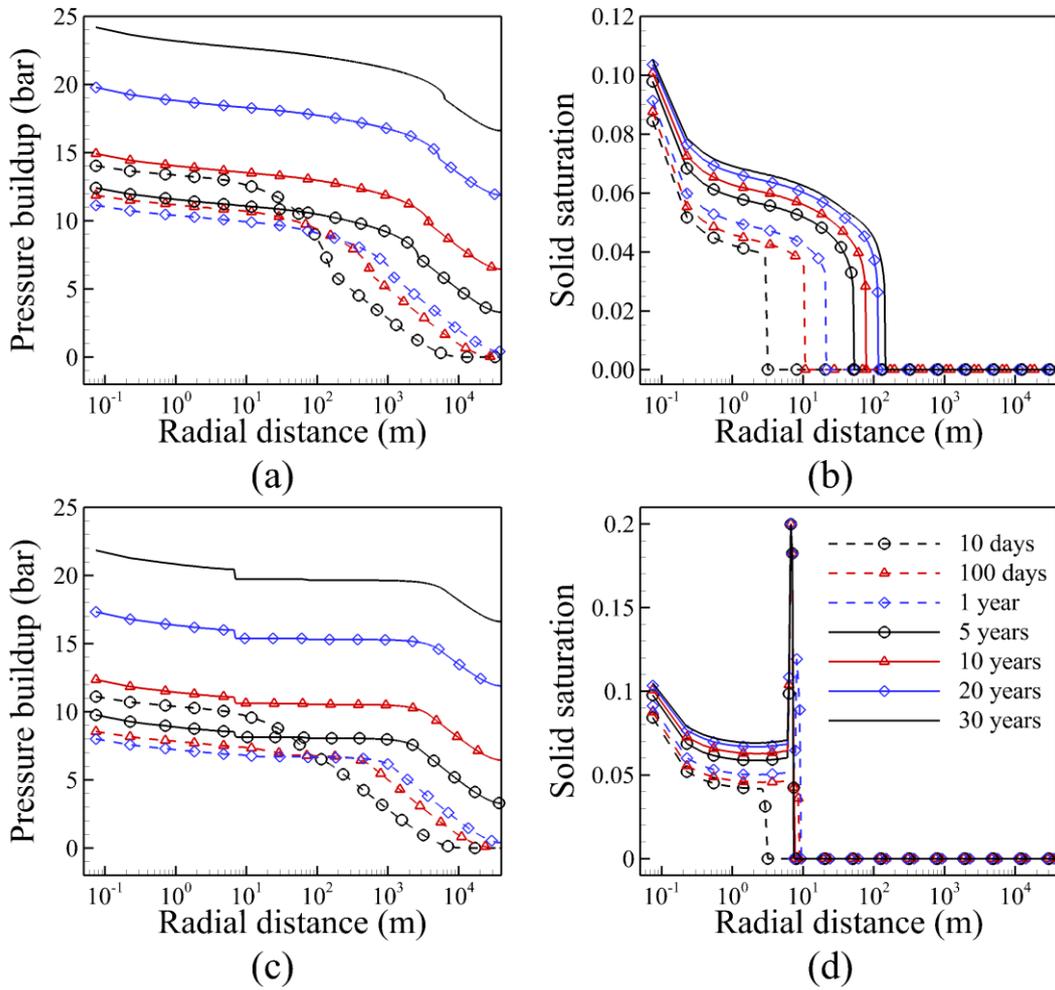
444 precipitation and the maximum value of salt precipitation in this zone increase. The higher
 445 precipitation zones at the interfaces contribute to reducing the leakage rate of gas phase from
 446 the storage saline into the seal formations.



447
 448 **Fig. 16.** Cross sections of solid saturation with seal permeability of (a1, a2) 10^{-19} m^2 , (b1, b2)
 449 10^{-18} m^2 , and (c1, c2) 10^{-17} m^2 for the semi-closed system.

450 Fig. 17 shows the radial profiles of pressure buildup and solid saturation for the semi-closed
 451 system with seal permeability of 10^{-19} m^2 at different time instants throughout the injection
 452 period. The profiles of pressure buildup show similar behaviours to those of the closed system.
 453 Due to the leakage of brine into the seal formations, the values of pressure buildup at the top
 454 and bottom aquifer are lower. The pressure profiles along the bottom aquifer also show a jump
 455 near the wellbore. The values of salt saturation near the well increase during the whole injection
 456 period, which are different from those in the closed and open storage systems. Due to the lower
 457 seal permeability and the capillary pressure, the injected CO_2 hardly enters into the seal
 458 formations. The injected CO_2 will accumulate under the interface and evaporate the water in the

459 brine at the interfaces continuously, which leads to the increase of solid saturation.



460
 461 **Fig. 17.** Profiles of (a, c) pressure buildup (unit: bar) and (b, d) solid saturation for the
 462 semi-closed system along (a-b) the top and (c-d) the bottom of the aquifer with seal
 463 permeability of 10^{-19} m^2 at different injection time instants.

464 4. Conclusions

465 Numerical simulations have been carried out for a better understanding of the phenomena of
 466 pressure buildup and salt precipitation during CO_2 injection period for carbon storage. In order
 467 to understand the effects of boundary conditions on CO_2 storage, three storage systems with
 468 different boundary conditions have been numerically simulated and compared. This study also

469 evaluates the flow of gas phase and the propagation of pressure, taking into account the effect of
470 precipitation. The main conclusions from the numerical simulations are given as follows:

471 (1) It has been shown that the region of elevated pressure is much larger than the CO₂ plume
472 size, while the salt precipitation due to the evaporation of gas phase only occurs in the small
473 zone of single gas phase.

474 (2) The pressure change shows different behaviours for the three systems. However, the
475 contour maps of solid saturation with the same injection rate for the three storage systems are
476 generally similar in shape, with several small differences in precipitation zone observed for the
477 three systems.

478 (3) There are two types of precipitation formed near the well, i.e., non-localized precipitation
479 near the injection well and localized precipitation in the lower portion of the dry-out front. The
480 evaporation of gas phase leads to precipitation near the well and the backflow of brine due to
481 capillary pressure results in the impervious zone near the lower portion of the well. The
482 formation processes of the two types of precipitation are different, which go through different
483 periods. For the semi-closed system, in addition to the two types of precipitation, a third type of
484 solid precipitation forms at the interfaces between the storage and seal formations. The salt
485 precipitation leads to the decrease of porosity and permeability and thus the degradation of
486 injectivity.

487 (4) The precipitation can affect the transportation of the gas phase and the propagation of
488 pressure. The localized precipitation acts as a barrier that suppresses the horizontal flow of gas
489 phase and promotes the upward flow of injected CO₂. The pressure profiles are smooth during

490 the early stage, while the curves reveal distinct gradients when the pores at the bottom aquifer
491 are clogged completely. It can be concluded that the injection rate is important for the salt
492 precipitation process. For the lower injection rate, more backflow of the brine occurs, leading to
493 more gas phase accumulating at the aquifer top, a narrower space for the gas phase flow and a
494 higher pressure jump at the bottom aquifer. The localized precipitation increases the risk of
495 leakage and reduces the security of CO₂ storage.

496 In the present study, the salt precipitation is treated as an immobile phase that clogs the pores.
497 In reality, the transportation of solid salt, from one location to another, can largely follow the
498 movement of fluids such as liquids and gases. However, the flow of precipitation is very
499 complicated and constitutive relations would be needed to specify the motion. In the future, a
500 more sophisticated model for the movement of solid precipitation will be considered. Moreover,
501 in order to effectively capture the dynamic behaviours of pressure buildup and salt precipitation
502 in full-scale carbon storage, sub-grid scale dynamics may be modelled using an upscaling
503 approach of the physical problem in a given time scale, which is being carried out.

504

505

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616

617 **Figure Captions.**

618

619 **Fig. 1.** Schematic representation of (a) CO₂ injection into a closed aquifer via a vertical well
620 and (b) top view.

621 **Fig. 2.** Schematic representation of boundary conditions for the three storage systems: (a) open
622 system, (b) closed system, and (c) semi-closed system.

623 **Fig. 3.** Spatial discretization considered in this study.

624 **Fig. 4.** A 16-meshes domain partitioning on 4 processors.

625 **Fig. 5.** The radial profiles of pressure buildup and solid saturation at the aquifer top for 10 days
626 (top) and 100 days (bottom) for the grid dependence test.

627 **Fig. 6.** Cross sections of pressure buildup (top: (b) is a zoom-in of (a)) and gas saturation
628 (bottom: (d) is the zoom-in of (c)) for the closed system at 30 years of CO₂ injection;
629 pressure unit: bar.

630 **Fig. 7.** Cross sections of solid saturation for the closed system at 30 years of CO₂ injection with
631 a close-up view of non-localized precipitation.

632 **Fig. 8.** Instantaneous iso-surfaces of solid saturation for the 3D closed system at different time
633 instants.

634 **Fig. 9.** The temporal evolution of gas saturation and solid saturation in (a) the non-localized
635 precipitation region, (b) the localized precipitation region.

636 **Fig. 10.** Profiles along the bottom of the aquifer for the closed system at different injection time

637 instants for (a) pressure buildup (unit: bar) and (b) solid saturation.

638 **Fig. 11.** Spatial distributions for the closed system at 60 years of CO₂ injection of (a) gas
639 saturation and (b) solid saturation.

640 **Fig. 12.** Profiles along the bottom of the aquifer for the closed system with injection rate of 50
641 kg/s at different injection time instants for (a) pressure buildup (unit: bar), and (b) solid
642 saturation.

643 **Fig. 13.** Cross sections of (a) pressure buildup (unit: bar), (b) gas saturation and (c) solid
644 saturation for the open system at 30 years of CO₂ injection.

645 **Fig. 14.** Profiles along the bottom of the aquifer for the open system at different injection time
646 instants for (a) pressure buildup, and (b) solid saturation.

647 **Fig. 15.** Cross sections of pressure buildup (unit: bar) with seal permeabilities of (a) 10^{-19} m², (b)
648 10^{-18} m², and (c) 10^{-17} m² for the semi-closed system.

649 **Fig. 16.** Cross sections of solid saturation with seal permeability of (a1, a2) 10^{-19} m², (b1, b2)
650 10^{-18} m², and (c1, c2) 10^{-17} m² for the semi-closed system.

651 **Fig. 17.** Profiles of (a, c) pressure buildup and (b, d) solid saturation for the semi-closed system
652 along (a-b) the top and (c-d) the bottom of the aquifer with seal permeability of 10^{-19} m²
653 at different injection time instants.

654

655 **Table Titles.**

656

657 **Tab. 1.** Hydrogeological properties of the storage formation.

658 **Tab. 2.** Example of domain partitioning and local numbering.