

22 **Abstract**

23 Atmospheric nitrogen (N) deposition in forests can affect soil microbial growth and turnover
24 directly through increasing N availability and indirectly through altering plant-derived carbon
25 (C) availability for microbes. This impacts microbial residues (i.e., necromass), a major
26 component of soil organic C. Previous studies in forest ecosystems have so far focused on the
27 impact of understory N addition on microbes and microbial residues, but the effect of N
28 deposition through plant canopy, the major pathway of N deposition in nature, has not been
29 explicitly explored. We investigated whether and how the quantity and modes (canopy vs.
30 understory) of N addition affect soil microbial residues in a temperate broadleaf forest under
31 10-yr N additions. Our results showed that N addition enhanced soil amino sugars and
32 microbial residual C concentrations, especially in the topsoil under high N addition. Canopy N
33 addition had stronger positive effects on soil amino sugars and microbial residual C than
34 understory N addition in the subsoil, implying that the indirect pathway via plants plays a more
35 important role. Also, neither canopy nor understory N addition significantly affected soil
36 microbial biomass and microbial community structure, suggesting that enhanced microbial
37 residues under N deposition stemmed from increased microbial biomass turnover. These
38 findings indicate that understory N addition underestimates the impact of N deposition on soil
39 microbial residues, suggesting that canopy related processes should also be considered in
40 temperate forest ecosystems.

41 **Keywords:** microbial residues, amino sugars, nitrogen deposition, canopy interception, soil
42 microbial community, microbial biomass turnover, nitrogen addition modes

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45 1. INTRODUCTION

46 The soil organic carbon (SOC) pool is twice larger than the amount of current atmospheric C,
47 and a small change in SOC may significantly affect the atmospheric CO₂ concentration (Smith,
48 2012). It is assumed that high atmospheric CO₂ may increase plant growth and then soil C
49 storage, especially if it is accompanied with an increasing reactive N input (Ackerman et al.,
50 2019; Stevens, 2019). Human activities, N fertilization in particular, have increased global
51 reactive N input and N deposition (Zaehle et al., 2011). Elevated N deposition alters soil N
52 availability for plants and microbes (Stevens et al., 2018). Because N is the primary limiting
53 factor of plant growth in terrestrial ecosystems, N promotion of plant growth likely enhances
54 microbial growth and allows microorganisms to transform more plant-derived C into microbial
55 residues (Ataka et al., 2020; Wang et al., 2022). Also, N input can modify the composition of
56 the soil microbial community, altering soil organic matter decomposition (Morrison et al.,
57 2016). Altered microbial growth and biomass turnover would affect microbial residues such as
58 amino sugars (Freppaz et al., 2014; Miltner et al., 2011; Ni et al., 2020). Microbial residual C
59 is an important component of SOC, accounting for up to half of the stable SOC pool in
60 terrestrial ecosystems (Liang et al., 2019; Wang et al. 2021a). Therefore, the knowledge of N
61 deposition effects on soil microbial residues is vital to comprehensively understand the soil C
62 dynamics under N deposition (Gilliam et al., 2019).

63 Many studies have examined the responses of soil microbial residues to atmospheric N
64 deposition (Averill et al., 2018; Griepentrog et al., 2014). Amino sugars, as specific biomarkers
65 of microbial residues in soils (Joergensen et al., 2018), have been extensively used to assess
66 the contribution of microbial-derived C to SOC and their responses to climate change (Liang

67 et al., 2019; Malik et al., 2020). Generally, changes in soil microbial biomass, microbial
68 community structure, and biomass turnover all can alter microbial residues (Wang et al., 2021b).
69 In forest ecosystems, N deposition has been shown to enhance soil microbial residual C in
70 some studies (Liao et al., 2022; Zhang et al., 2023), but not in others (Ma et al., 2021; Yuan et
71 al., 2021; Zhang et al., 2016). Also, N addition can reduce fungal residues in soil as high
72 mineral N often suppresses fungi (Ma et al., 2020; Treseder et al., 2008). These inconsistent
73 results may stem from the contrasting responses of soil microbial groups (bacteria and fungi)
74 to N addition. Alternatively, they may be due to various responses of microbial turnover and
75 soil organic matter decomposition to changes in soil physicochemical properties (e.g., soil pH)
76 and plant associated processes (e.g., quantity and/or quality of rhizodeposition and/or litter)
77 caused by N addition (Khan et al., 2016; Wang et al., 2023). Also, the rate of N additions impact
78 soil microbial biomass and community structure (Frey et al., 2004; Tian et al., 2022). For
79 instance, low N addition usually increases microbial biomass in forests (Waldrop et al., 2004),
80 but high N addition may reduce microbial biomass and suppress microbial activity and
81 microbial decomposition of organic matter (Jing et al., 2021; Meunier et al., 2016). However,
82 the exact mechanisms underpinning N rate effects on microbial residues are still unclear in
83 temperate forests.

84 The current understanding of atmospheric N deposition effects on forest SOC and
85 microbial residues is mainly built on the studies of understory N additions (Chang et al., 2019;
86 Gurmesa et al., 2022). Understory N addition affects soil microbial biomass and community
87 structure mainly via direct N effects on microbes and soil physicochemical properties (Chen et
88 al., 2023; Jia et al., 2020). Understory N addition may overestimate the N effect on soil

89 microbes because of the absence of canopy N interception (Zhang et al., 2015). Forest canopy
90 may retain a remarkable proportion of incoming atmospheric N and alter the quality and
91 quantity of N that enters into soils (Guerrieri et al., 2021). For example, canopy leaves, twigs
92 and branches could absorb the N from canopy addition (Houle et al., 2015). These canopy-
93 associated processes will indirectly affect soil microbial biomass, microbial community
94 structure, and soil organic matter decomposition through altering litter properties, soil-derived
95 N absorption and C availability to microbes. It was reported that canopy N interception
96 mitigated the direct impact of N on soil biota (Liu et al., 2020). Moreover, N addition modes
97 also affect the magnitude of N-induced changes in soil physicochemical properties and plant
98 chemical properties. Therefore, the mode of N addition may exert a critical control over the N
99 effects on soil ecological and plant physiological processes that affect C cycling in forest
100 ecosystems (Lu et al., 2021; Yu et al., 2019).

101 To explore the effects of N deposition modes and quantities on forest ecosystem structure
102 and functions, we initialized a long-term field manipulation study with two modes (canopy and
103 understory additions) and three N-addition rates in a temperate forest in 2013. As a component
104 of this extensive field study, we examined the effects of N deposition on soil microbes and
105 microbial C residues in soil. Because forest litters have high C:N ratios and microbes in forest
106 soils are also N-limited, we predict that N addition alleviates N limitation on microbes and
107 increases microbial biomass and microbial residues in soil (Hypothesis 1). Given that canopy
108 N interception reduces the quantity of N that directly enters soil, we also expect that understory
109 N addition has stronger effects on microbial residues than canopy N addition (Hypothesis 2).

110 **2. MATERIALS AND METHODS**

111 **2.1. Site description**

112 This study was conducted at the Dabieshan National Field Observation and Research Station
113 of Forest Ecosystem (31°46'-31°52'N, 114°01'-114°06'E), located in the Jigongshan National
114 Nature Reserve, in Henan Province of China. The climate is characterized by a subtropical-
115 warm temperate climate. In this studied region, the mean annual precipitation was
116 approximately 1119 mm and the mean annual temperature was 15.2 °C (Zhang et al., 2015).
117 The temperate deciduous broadleaf forest is a typical vegetation type, and the dominant tree
118 species include *Liquidambar formosana* Hance, *Quercus acutissima* Carruth, and *Quercus*
119 *variabilis* Bl. The forest soil is classified as yellow-brown sandy-loam soil (Liu et al., 2020).
120 The background wet N deposition rate was approximately 20 kg N ha⁻¹ yr⁻¹ at this study site
121 (Zhang et al., 2015).

122 **2.2. Experimental design**

123 This experiment included understory N addition (UN) and canopy N addition (CN), and was
124 set as a randomized block design with four blocks. Each block contained five treatments: CT
125 (control, no N addition), CN25 (canopy N addition at 25 kg N ha⁻¹ yr⁻¹, low-N), CN50 (canopy
126 N addition at 50 kg N ha⁻¹ yr⁻¹, high-N), UN25 (understory N addition at 25 kg N ha⁻¹ yr⁻¹,
127 low-N), and UN50 (understory N addition at 50 kg N ha⁻¹ yr⁻¹, high-N). There were 20 plots
128 in total, and one plot was a circle with a diameter of 34 m, which was surrounded by a > 20 m
129 buffer zone. Cement boards were inserted into the soil to a depth of 50 cm to separate each plot
130 from adjacent plots. The N form in the treatments was NH₄NO₃ solution, which was sprayed
131 monthly from April to October (seven times per year). All treatments were initiated in April
132 2013. More details on this experimental design were included in Zhang et al. (2015).

133 **2.3. Soil sampling and physicochemical properties**

134 Soil samples were collected from the depths of 0-10 cm and 10-20 cm in July 2022, respectively,
135 corresponding to the 10th year of N addition. In each plot, five cores with the same depth were
136 sampled by an auger (3.0 cm in diameter) from five randomly selected microsites and bulked
137 into one pooled sample. There were 40 soil samples in total. Visible plant materials and roots
138 were removed by hand. All fresh soil samples were sieved using a 2 mm soil sieve and taken
139 back to the laboratory. Each soil sample was divided into two subsamples: one was air-dried
140 for analysis of soil physicochemical properties, and the other was stored at -20°C for
141 determination of microbial biomass and community structure.

142 Soil pH was measured in a slurry (soil: water = 1: 2.5, w/v) with a pH meter (FiveEasy
143 PlusTM FE28, Mettler Toledo). SOC, soil total nitrogen (TN) and total phosphorus (TP)
144 concentrations were determined by the concentrated sulfuric acid-potassium dichromate
145 external heating method, the concentrated sulfuric acid digestion-phenol blue colorimetric
146 method, and the concentrated sulfuric acid digestion-molybdenum antimony anti-colorimetric
147 method, respectively (Lu, 1999).

148 **2.4. Soil microbial biomass and community structure**

149 Soil microbial community was characterized by the phospholipid fatty acids (PLFAs) method
150 (Bossio & Scow 1998). Concentration of each PLFA was calculated based on 19:0 internal
151 standard concentration. The PLFAs i14:0, i15:0, a15:0, i16:0, a16:0, i17:0, a17:0, a18:0, i18:0,
152 a19:0, 16:1 ω 7c, 16:1 ω 9c, 17:1 ω 8c, 18:1 ω 7, cy17:0, and cy19:0 were used to indicate bacterial
153 biomarkers. The PLFAs 18:1 ω 9c, 18:2 ω 6,9c, and 18:3 ω 6,9,12c were applied to denote fungal
154 biomarkers. The PLFA 16:1 ω 5c was considered as arbuscular mycorrhizal fungal (AMF)

155 biomarker. The PLFAs 10Me 16:0, 10Me 17:0, and 10Me 18:0 were used as actinomycetes
156 biomarkers. Total microbial biomass was represented by the sum of bacterial, fungal, AMF,
157 and actinomycetes biomarkers. Soil microbial community structure was represented by fungal:
158 bacterial biomass ratio (F:B ratio) (Frostegård & Bååth 1996).

159 **2.5. Soil amino sugars and microbial residual C**

160 Soil amino sugar (ASs) concentrations, including muramic acid (MurN), galactosamine (GalN),
161 and glucosamine (GluN), were determined as described by Indorf et al. (2011). In brief, amino
162 sugars were hydrolyzed, extracted, and derivatized with ortho-phthaldialdehyde, determined
163 by high-performance liquid chromatography (Dionex Ultimate 3000, Thermo Fisher Scientific).
164 The detailed relevant information was described by Yuan et al. (2021). Microbial residual C
165 (MRC) was calculated by the following formulas:

$$166 \quad \text{F-GluN } (\mu\text{g g}^{-1}) = \text{total GluN } (\mu\text{g g}^{-1}) - 2 \times \text{MurN } (\mu\text{g g}^{-1}) \times (179.2/251.2) \quad (1)$$

$$167 \quad \text{Fungal MRC } (\mu\text{g g}^{-1}) = \text{F-GluN} \times 9 \quad (2)$$

$$168 \quad \text{Bacterial MRC } (\mu\text{g g}^{-1}) = \text{MurN} \times 45 \quad (3)$$

$$169 \quad \text{Total MRC } (\mu\text{g g}^{-1}) = \text{Fungal MRC} + \text{Bacterial MRC} \quad (4)$$

170 Where F-GluN is fungi-derived GluN. Fungal MRC, bacterial MRC, and total MRC are
171 fungi-derived, bacteria-derived, and total microbial residual C, respectively. It was assumed
172 that MurN and GluN occurred at a 1 to 2 molar ratio in bacterial cell walls (Engelking et al.,
173 2007). Where 179.2 and 251.2 are the molecular weights of GluN and MurN, respectively
174 (Shao et al., 2017). Where 9 and 45 are conversion factors (Appuhn & Joergensen, 2006).

175 **2.6. Data analysis**

176 One-way ANOVA was employed to examine the effects of N deposition on concentrations of
177 soil total amino sugar (total ASs), MurN, GalN, GluN, fungal MRC, bacterial MRC, total MRC,
178 and the ratio of fungal to bacterial MRC. The effects of N deposition on the contributions of
179 total ASs, fungal MRC, bacterial MRC, and total MRC to SOC were tested by one-way
180 ANOVA as well. The N deposition impacts on soil physicochemical properties (pH, SOC, TN,
181 TP) and soil microbial parameters (the biomasses of fungi, bacteria, AMF, actinomycetes, and
182 total microbes, and the F:B ratio) were examined by one-way ANOVA. Multiple comparison
183 analyses (LSD) were used after one-way ANOVA. Pearson correlation analysis was performed
184 to assess the relationships of measured soil microbial residues (amino sugars and residual C)
185 with soil physicochemical properties and soil microbial parameters (microbial biomass and
186 community structure). All statistical analyses were carried out with SPSS 18.0 (SPSS, Chicago,
187 Illinois, USA), and results were considered statistically significant at $p < 0.05$.

188 **3. RESULTS**

189 **3.1. Soil amino sugars and microbial residual C**

190 High-N addition ($50 \text{ kg N ha}^{-1} \text{ yr}^{-1}$) significantly enhanced soil amino sugar concentrations in
191 the 0-10 cm layer, regardless of canopy or understory N addition (Figure 1a,c). Higher
192 concentrations of MurN, GluN, and total amino sugars were observed in CN50 and UN50,
193 compared with that in CT. They increased by 43.50-53.36% and 58.09-62.58% for CN50 and
194 UN50, respectively. The concentration of GalN was significantly higher in UN50 not CN50
195 than that in CT ($p = 0.010$ and 0.066 , respectively). Low-N addition ($25 \text{ kg N ha}^{-1} \text{ yr}^{-1}$) did not
196 significantly increase soil amino sugar concentrations including MurN, GalN, GluN, and total

197 amino sugars, no matter canopy or understory N addition. In the 10-20 cm layer, canopy N
198 addition increased soil amino sugar concentrations, especially canopy high-N addition (Figure
199 1b,d). Specifically, the concentrations of MurN, GluN, and total amino sugars in CN50 were
200 significantly higher than that in CT. Also, the MurN concentration in CN25 was higher than
201 that in CT ($p = 0.030$). Neither understory high-N nor low-N addition significantly affected soil
202 amino sugar concentrations in the 10-20 cm layer (Figure 1b,d), and the GalN concentration
203 was not significantly affected by canopy or understory N addition. The mean contributions of
204 total amino sugars to SOC were 3.92-4.63% and 4.37-5.20% in the 0-10 cm and 10-20 cm
205 layers, respectively, these were not altered by canopy or understory N addition (Figure 2).
206 Overall, the contribution of total amino sugars to SOC was greater in the 10-20 cm than that in
207 the 0-10 cm layer ($p = 0.023$, Figure 2).

208 Both the canopy and understory high-N additions significantly increased soil microbial
209 residual C (MRC) concentration in the 0-10 cm layer. Specifically, the fungal MRC, bacterial
210 MRC, and total MRC concentrations were significantly higher in the CN50 and UN50 than
211 that in the CT (Figure 3a). Canopy and understory low-N additions did not significantly affect
212 the microbial residual C as fungal MRC, bacterial MRC, and total MRC. Besides, the ratio of
213 fungal to bacterial MRC was not significantly influenced by canopy and understory N addition
214 (Figure 3a). In the 10-20 cm layer, canopy not understory N addition increased soil MRC
215 concentrations, especially canopy high-N addition (Figure 3b). The concentrations of fungal,
216 bacterial, and total MRC were significantly higher in CN50 than that in CT. The bacterial MRC
217 concentration in CN25 were also significantly higher than that in CT. While no matter low- or
218 high-N, understory N addition did not significantly affect microbial residue C (Figure 3b). The

219 ratios of fungal to bacterial MRC were not also affected by canopy and understory N additions
220 (Figure 3b). The contributions of MRC (fungal, bacterial, and total MRC) to SOC were not
221 significantly altered by canopy and understory N additions in the 0-10 cm and 10-20 cm layers
222 as well (Figure 3c,d).

223 **3.2. Soil microbial biomass and community structure**

224 In the 0-10 cm layer, N addition did not significantly affect soil microbial biomass such as the
225 biomasses of fungi, bacteria, actinomycetes, AMF, and total microbes (Table 1; all $p > 0.05$).
226 The F:B ratio was not significantly affected by N addition, regardless of canopy or understory
227 N addition. In the 10-20 cm layer, there were no significant differences in soil microbial
228 biomass including the biomasses of bacteria, fungi, actinomycetes, AMF, and total microbes,
229 and the F:B ratio (Table 1; all $p > 0.05$). The biomasses of bacteria, fungi, actinomycetes, AMF,
230 and total microbes were also not significantly higher under the N addition treatments relative
231 to CT in the 0-10 and 10-20 cm layers.

232 **3.3. Soil physiochemical properties**

233 In the 0-10 cm layer, N addition altered soil pH and soil total N concentration. The soil pH was
234 lower in CN25, UN25, and UN50 than that in CT (Table 2). Higher soil N concentration was
235 observed in CN25, CN50, and UN50 than in CT. The SOC concentrations were slightly higher
236 in the treatments of N addition than in CT, although the differences were not statistically
237 significant ($p = 0.472$). There was no significant difference in soil total P concentration among
238 different treatments ($p = 0.934$, Table 2). In the 10-20 cm layer, the patterns of soil pH, SOC,
239 and total P in all treatments were the similar to that in the 0-10 cm layer. Specifically, soil pH

240 was significantly lower in CN25, UN25, and UN50 than that in CT. No significant differences
241 in SOC and total P concentrations were found among different treatments. The treatments did
242 not differ in soil total N concentrations, although they were slightly higher in N addition
243 treatments than that in CT in the 10-20 cm layer (Table 2).

244 **3.4. Linkages between microbial residues and soil properties**

245 Soil physicochemical properties and microbial biomass affected soil amino sugars. The
246 concentrations of amino sugars including MurN, GalN, and GluN, and total ASs were
247 negatively correlated to soil pH, but positively correlated to SOC, TN, and biomasses of
248 bacteria, fungi, actinomycetes, and total microbes in the 0-10 cm layer (Figure 4a). In the 10-
249 20 cm layer, the concentrations of MurN, GalN, GluN, and total ASs were not affected by soil
250 pH, but positively associated with SOC, TN (except GalN), and biomasses of bacteria, AMF,
251 actinomycetes and total microbes (Figure 4b). In addition, the contributions of amino sugars
252 (MurN, GalN, GluN, and total ASs) to SOC were only negatively correlated to SOC in the 0-
253 10 cm and 10-20 cm layers. The AMF biomass negatively affected the contributions of total
254 amino sugars and GalN to SOC in the 0-10 cm layer (Figure 4a). However, other soil
255 parameters (i.e., pH, TN, TP, and biomasses of bacteria, fungi, AMF, actinomycete, and total
256 microbes) did not markedly alter the contributions of amino sugars to SOC in the two studied
257 soil layers (Figure 4a,b).

258 The microbial residual C was significantly affected by soil properties in the two studied
259 soil layers. Fungal MRC, bacterial MRC, and total MRC were negatively correlated to soil pH,
260 but positively correlated to SOC, TN, and biomasses of bacteria, fungi, actinomycetes, and
261 total microbes in the 0-10 cm layer (Figure 4a). In the 10-20 cm layer, there was no significant

262 correlations between microbial residual C (fungal MRC, bacterial MRC, and total MRC) and
263 soil pH. While fungal MRC, bacterial MRC, and total MRC were positively associated with
264 SOC, TN, and biomasses of bacteria, AMF, actinomycetes, and total microbes (Figure 4b).
265 Furthermore, the contributions of microbial residual C to SOC were only negatively affected
266 by SOC in the 0-10 cm and 10-20 cm layers (Figure 4a,b). The ratio of fungal to bacterial MRC
267 was not affected by the studied soil physiochemical properties and microbial biomass in the 0-
268 10 cm and 10-20 cm layers (Figure 4a,b).

269 **4. DISCUSSION**

270 **4.1. Dominant determinants of response of soil microbial residues to N addition**

271 The changes in soil microbial biomass following N addition can reveal the rapid response of
272 microbial growth and metabolism to N addition (Ma et al., 2020). Nitrogen addition had a
273 minor effect on soil microbial biomass in this study, which did not support our first hypothesis
274 that soil microbial biomass would increase under N addition. Additionally, the F:B ratio was
275 not altered by N addition. The lack of response suggests that N was not the most important
276 limiting factor. Microbial community variations could be limited by soil moisture or
277 availability of other nutrients, or other environmental factors (Liu et al., 2022). The effect of N
278 on soil microbial biomass was thus not significant in this study but further studies and long-
279 time monitoring of soil microbial dynamics are needed.

280 In this study, high N not low N deposition significantly enhanced individual and total
281 amino sugar concentrations in the topsoil, which partly supported our hypothesis that N
282 addition would increase microbial residues. The response of microbial residue accumulation to

283 N deposition is related to N addition rate (Tian et al., 2022). Low N addition did not cause
284 changes in soil physiochemical properties and soil microbial biomass, possibly due to the soil
285 capacity to buffer disturbance to some extent (Lo Cascio et al., 2021). Additionally, the
286 significantly positive correlations between soil amino sugars and microbial biomass were
287 detected in this study, although high-N addition did not significantly affect soil microbial
288 biomass. The soil physiochemical properties such as soil pH, SOC, and soil TN displayed
289 significant effects on amino sugars, which could change soil amino sugars via altering soil
290 microbial turnover (Fig. 5) (Brabcová et al., 2018). Since amino sugars provides the
291 information on time-integrated microbial community (Glaser et al., 2004), we propose that the
292 changes in amino sugar concentration could be substantially resulted from the alteration of soil
293 microbial turnover rather than microbial biomass.

294 Soil microbial residual C is regulated by soil microbial residues deposition and
295 decomposition (Fernandez et al., 2019; Freedman et al., 2016; Shao et al., 2021). Soil microbial
296 biomass is a key factor determining soil microbial residues deposition. In this study, although
297 N addition (canopy and understory N addition) did not markedly alter soil microbial biomass
298 and community structure, the positive correlations between microbial residue C and microbial
299 biomass were observed. Microbial residual C is a long-term accumulated product from soil
300 microbes and is not fully equal to living microbial biomass (Camenzind et al., 2023; Zhang et
301 al., 2021). In the long term, a non-significant difference in microbial biomass also may generate
302 an obvious discrepancy in microbial residual C. The previous study reported that N addition
303 did not increase the activity of residue-decomposing enzymes (Yuan et al., 2021). So it is more
304 probably ascribed to the increase in microbial residue deposition rather than reduction in

305 microbial residue decomposition. High N addition increases soil N availability and plant
306 productivity (Lebauer & Tresder, 2008; Li et al., 2021). The enhanced plant productivity needs
307 more nutrient supply (Giardina et al., 2003). To obtain more nutrients, fast turnover rate of
308 microbial biomass thus would be triggered and result in more microbial residue C accumulation.

309 **4.2. Differentiated responses of soil microbial residues to N addition modes**

310 Canopy and understory N additions exhibited contrast effects on soil microbial residues in this
311 study. Amino sugar and microbial residual C concentrations were enhanced by canopy high-N
312 addition in two soil layers, but increased by understory high-N addition only in the 0-10 cm
313 layer. The result was not consistent with our hypothesis that the effect of understory N addition
314 on microbial residues is greater than that of canopy N addition. It could be attributable to the
315 different impact pathways of canopy and understory N additions on soil microbial residues (Fig.
316 5).

317 The indirect pathway via plants could dominate the N effect on soil microbial residues
318 (Figure 5). Canopy N addition directly influences plants and then indirectly influences soil
319 microbes through altering the traits of leaf, litter, and root of canopy trees, and C availability
320 to soil microbes (Cantarel et al., 2015; Feng et al., 2022; terHorst & Zee, 2016). For instance,
321 the increase in leaf N concentration induced by canopy N addition may improve leaf
322 photosynthesis (Li et al., 2021; Wang et al., 2021c). Whereafter, the amount of photosynthates
323 allocated to belowground plant parts would be enhanced (Farrar & Jones, 2000; Hendricks et
324 al., 2000). The fine root biomass and production were significantly higher with canopy N
325 addition than with understory N addition in the same experimental plots as used in this study
326 (Li et al., 2021). The boosted fine root biomass and production may stimulate soil microbial

327 activities and promote microbial turnover to obtain more nutrients. By contrast, the N from
328 understory addition is sprayed onto the forest floor and go directly into the soil (Figure 5).
329 Understory low-N addition did not cause the significant changes in soil properties and trigger
330 the response of soil microbes. The low N with understory addition may not be enough to lead
331 to a significant and timely response of canopy trees (Forsmark et al., 2021). Alternatively, the
332 negative effect of acidification could be offsetting the positive effect of enhanced resources.

333 **5. CONCLUSIONS**

334 Canopy and understory N additions differed in effects on soil microbial residues. Both canopy
335 and understory high-N additions enhanced soil amino sugar and microbial residual C
336 concentrations in the topsoil. Whereas, in the subsoil, canopy but not understory N addition
337 exhibited a positive effect on soil amino sugars and microbial residue C, especially canopy
338 high-N addition. Moreover, neither canopy nor understory N addition significantly affected soil
339 microbial biomass and community structure. This experimental evidence suggests that
340 microbial turnover may make a notable contribution to microbial residue accumulation relative
341 to microbial biomass, and the N addition impact on soil microbial residues relies on simulation
342 modes of N deposition and N addition rates. These findings indicate understory N addition
343 underestimates the impact of N deposition on soil microbial residues, and N addition effects
344 on plant canopy, plant growth, and microbial turnover exert a major control over the formation
345 of microbial-derived SOC. We thus propose canopy related processes should be considered
346 when assessing and predicting the effect of N deposition on C sink function in temperate forest
347 ecosystems.

348 **AUTHOR CONTRIBUTIONS**

349 Yuanqi Chen, Shirong Liu and Shenglei Fu conceived the study; Yuanqi Chen, Yu Zhang, Xu
350 Zhang, and Xiaowei Li conducted this study; Yuanqi Chen, Teng Feng, and Quan Chen
351 analyzed the data; Yuanqi Chen, Yu Zhang and Xu Zhang wrote the manuscript; Yuanqi Chen,
352 Shirong Liu, Carly Stevens and Shuijin Hu revised the manuscript.

353 **ACKNOWLEDGEMENTS**

354 The authors express their sincere gratitude to Drs. Weixin Zhang and Jie Zhao for their helps
355 in the experiment design. We also thank Dr. Lei Ma and Mrs. Zihao Qiu, Yupeng Yan, and
356 Shengfu Chen for their helps in the fieldwork. This work was financially supported by the
357 National Natural Science Foundation of China (31901194, U21A20189, and 32271729).

358 **CONFLICT OF INTEREST**

359 The authors declare that they have no conflict of interest.

360 **DATA AVAILABILITY STATEMENT**

361 All data are available in the main text or the Supporting Information and raw data are available
362 upon request from the corresponding author.

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Figure legends

Figure 1 Soil amino sugar concentrations in all treatments in the 0-10 cm (1a, 1b) and 10-20 cm (1c, 1d) layers. GalN, GluN, MurN, and Total ASs refer to galactosamine, glucosamine, muramic acid, and total amino sugars. CT, CN25, CN50, UN25, and UN50 stand for the treatments of control, canopy N addition at 25 kg N ha⁻¹ yr⁻¹ and 50 kg N ha⁻¹ yr⁻¹, understory N addition at 25 kg N ha⁻¹ yr⁻¹ and 50 kg N ha⁻¹ yr⁻¹, respectively. Values are means ± SE, n = 4. Different lowercase letters indicate significant differences in the same amino sugar among different treatments in the same soil layer at the $p = 0.05$ level.

Figure 2 The contribution of total amino sugars (ASs) to SOC in all treatments in the 0-10 cm and 10-20 cm layers. Values are means ± SE, n = 4. Different capital letters on the line indicate significant differences between 0-10 cm and 10-20 cm soil layers at the $p = 0.05$ level. See Figure 1 for abbreviations.

Figure 3 Microbial residual C (MRC) and their contribution to SOC in all treatments in the 0-10 cm (3a, 3c) and 10-20 cm (3b, 3d) soil layers. B-MRC, F-MRC, F/B-MRC, and T-MRC stand for the bacteria-derived MRC, fungi-derived MRC, the ratio of F-MRC to B-MRC, and total MRC, respectively. B-MRC/SOC, F-MRC/SOC, and T-MRC/SOC represent the contributions of B-MRC, F-MRC, and T-MRC to SOC, respectively. Values are means ± SE, n = 4. Different lowercase letters indicate significant differences in soil MRC among different treatments in the same soil layer at the $p = 0.05$ level. See Figure 1 for abbreviations.

Figure 4 The correlations between microbial residues (amino sugars and microbial residual C) and soil physiochemical properties in 0-10 cm (a) and 10-20 cm (b) soil layers. SOC, TN, TP, B, F, AMF, Act, and T-PLFAs stand for soil organic C, soil total N, soil total P, the biomasses of bacteria, fungi, arbuscular mycorrhizal fungi, actinomyces, and total microbes, respectively. See Figure 1 and Figure 3 for abbreviations.

Figure 5 Schematic diagram summarizing the effects of canopy N addition and understory N addition on soil microbial residues. The thin and thick arrows indicate the effects from one factor and many factors, respectively.

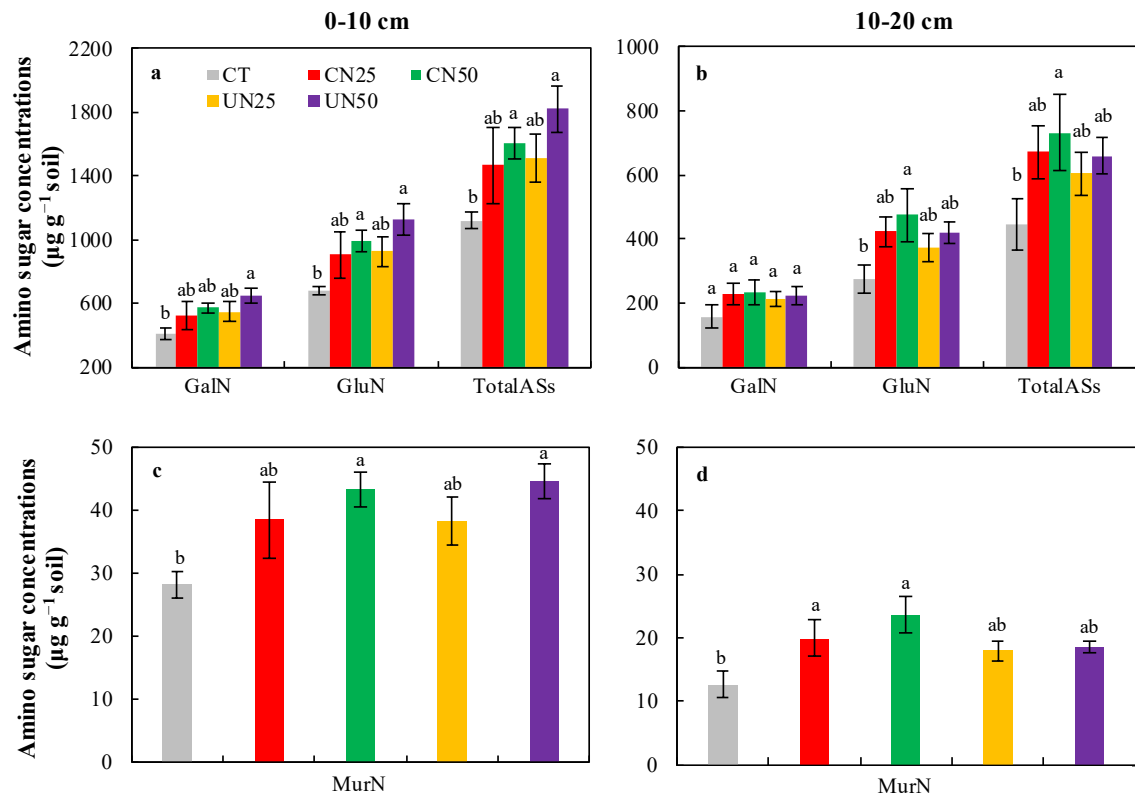


Fig. 1

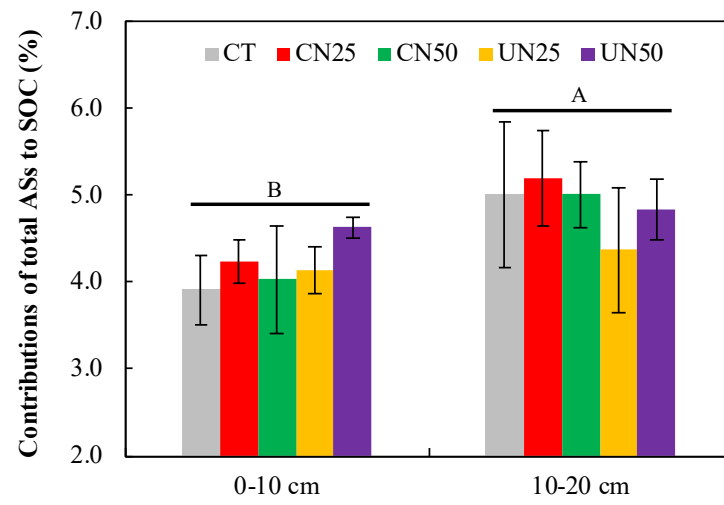


Fig. 2

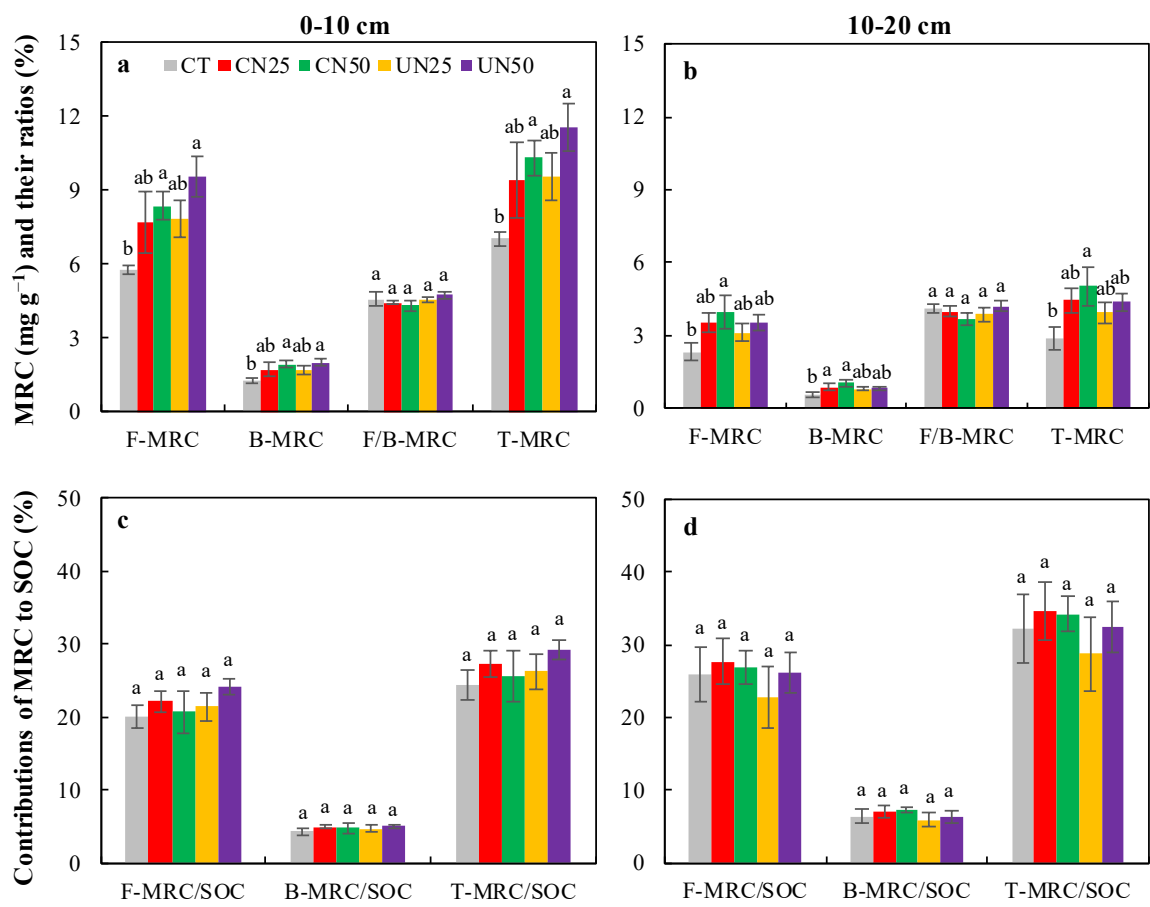


Fig. 3

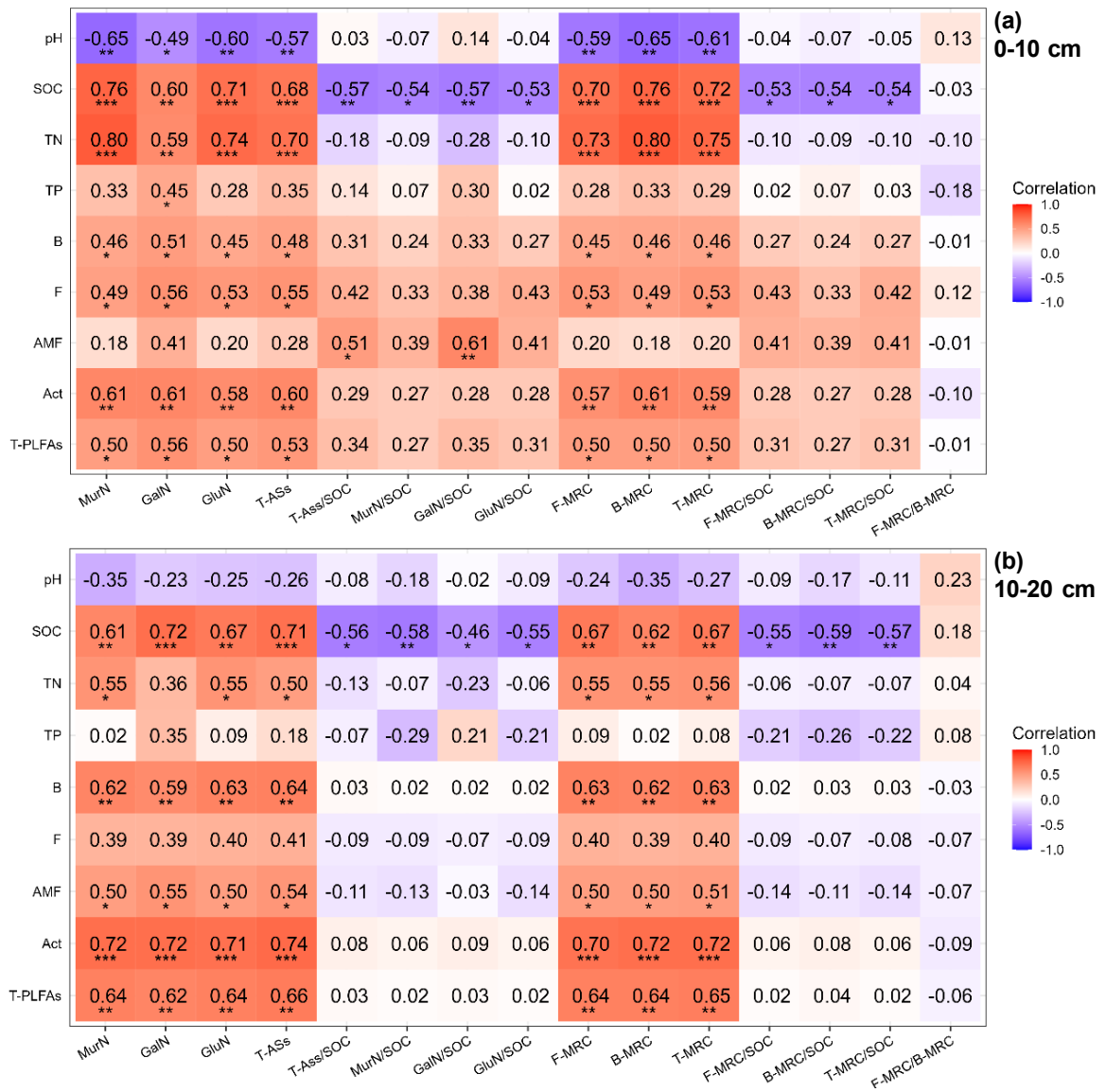


Fig. 4

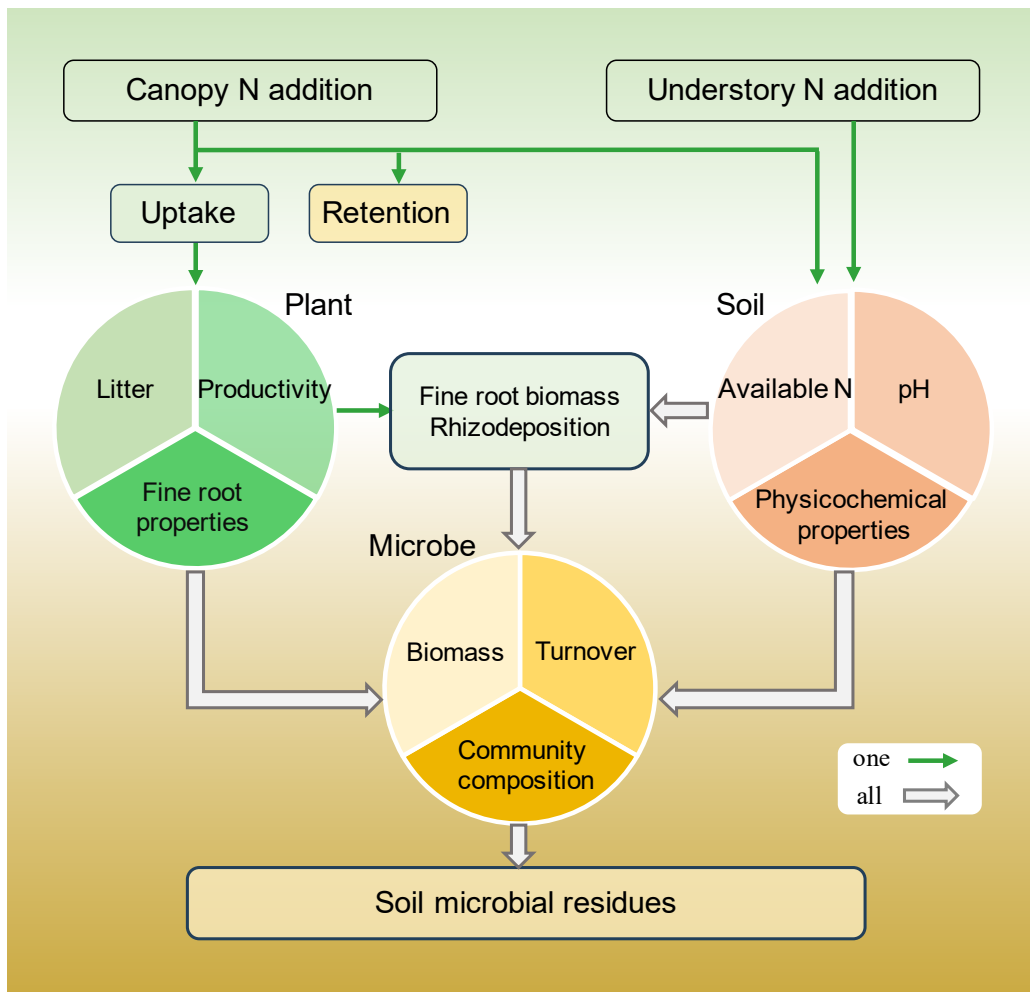


Fig. 5

Table 1 Effects of N addition on soil microbial biomass and community structure by PLFAs methods.

		B	F	AMF	Act	TMB	F: B ratio
0-10 cm	F value	0.34	0.52	0.09	0.50	0.34	1.73
	<i>p</i> value	0.847	0.721	0.985	0.739	0.845	0.197
10-20 cm	F value	0.64	0.26	0.33	1.31	0.71	0.34
	<i>p</i> value	0.642	0.899	0.856	0.312	0.598	0.846

Note: B, F, AMF, Act, TMB, and F:B ratio stand for the biomasses of bacteria, fungi, arbuscular mycorrhizal fungi, actinomycetes, and total microbes, and the ratio of fungal to bacterial biomass respectively. *F* and *p* value is the result of one-way ANOVA. The significant difference was set at the $p = 0.05$ levels.

Table 2 Soil physiochemical properties in all treatments in the 10th years after treatments.

	Treatments	pH	SOC	TN	TP
0-10 cm	CT	4.45a	29.38	1.04b	0.33
	CN25	4.12b	34.88	1.65a	0.27
	CN50	4.18ab	44.09	1.82a	0.34
	UN25	4.10b	37.15	1.37ab	0.30
	UN50	4.11b	39.31	1.69a	0.34
	<i>p</i> value	0.084	0.472	0.096	0.934
10-20 cm	CT	4.53a	8.96	0.30	0.32
	CN25	4.27b	13.77	0.52	0.23
	CN50	4.37ab	14.88	0.56	0.26
	UN25	4.29b	15.13	0.50	0.31
	UN50	4.31b	13.95	0.46	0.28
	<i>p</i> value	0.045	0.386	0.581	0.917

Note: pH, SOC, TN, TP stand for the soil pH value, soil organic carbon, soil total nitrogen, and soil total phosphorus, respectively. Values are means, n = 4. CT, CN25, CN50, UN25, and UN50 stand for the control, canopy N addition at 25 kg N ha⁻¹ yr⁻¹ and 50 kg N ha⁻¹ yr⁻¹, understory N addition at 25 kg N ha⁻¹ yr⁻¹ and 50 kg N ha⁻¹ yr⁻¹, respectively. *P* value is the result of one-way ANOVA in the same soil layer. Different lowercase letters indicate significant differences in the same soil parameter among different treatments in the same soil layer at the *p* = 0.05 levels.