1	Changepoint Detection: An Analysis of the Central England Temperature
2	Series
3	Xueheng Shi
4	Department of Statistics, University of California, Santa Cruz
5	Claudie Beaulieu*
6	Department of Ocean Sciences, University of California, Santa Cruz
7	Rebecca Killick
8	Department of Statistics, Lancaster University
9	Robert Lund
10	Department of Statistics, University of California, Santa Cruz

¹¹ **Corresponding author*: Claudie Beaulieu, beaulieu@ucsc.edu

ABSTRACT

This paper presents a statistical analysis of structural changes in the Central England tempera-12 ture series, one of the longest surface temperature records available. A changepoint analysis is 13 performed to detect abrupt changes, which can be regarded as a preliminary step before further 14 analysis is conducted to identify the causes of the changes (e.g., artificial, human-induced or natural 15 variability). Regression models with structural breaks, including mean and trend shifts, are fitted 16 to the series and compared via two commonly used multiple changepoint penalized likelihood cri-17 teria that balance model fit quality (as measured by likelihood) against parsimony considerations. 18 Our changepoint model fits, with independent and short-memory errors, are also compared with 19 a different class of models termed long-memory models that have been previously used by other 20 authors to describe persistence features in temperature series. In the end, the optimal model is 21 judged to be one containing a changepoint in the late 1980s, with a transition to an intensified 22 warming regime. This timing and warming conclusion is consistent across changepoint models 23 compared in this analysis. The variability of the series is not found to be significantly changing, 24 and shift features are judged to be more plausible than either short- or long-memory autocorrela-25 tions. The final proposed model is one including trend-shifts (both intercept and slope parameters) 26 with independent errors. The analysis serves as a walk-through tutorial of different changepoint 27 techniques, illustrating what can be statistically inferred. 28

2

29 1. Introduction

Climate time series often contain abrupt changes and other nonlinearities in their behavior. 30 Changepoints are times of abrupt shifts in a series' characteristics, including means, trends, vari-31 ances, and autocorrelations. For examples, a sudden change from a cooling period (i.e., decreasing 32 trend) to a warming period can be characterised by a changepoint in the trend; a sudden increase due 33 to the relocation of a station may be characterised as a changepoint in the mean. Abrupt changes 34 may be caused by changes in climate forcings, related to climate variability in the ocean and 35 atmosphere, or induced by artificial changes in measurement procedures such as station relocations 36 or instrumentation changes. 37

It is crucial to know changepoint times in climate series, especially when assessing long-term 38 trends, as their presence may grossly alter trend estimates, which impedes our understanding of 39 external forcings and climate variability over the instrumental record (Lund et al. 2007; Beaulieu 40 et al. 2012; Cahill et al. 2015; Beaulieu and Killick 2018). Series with artificial changes merit 41 adjustment via homogenization methods, as trends and extreme quantiles are more accurately 42 estimated from homogenized data (Hewaarachchi et al. 2017; Trewin et al. 2020; Vincent et al. 43 2020). On average, approximately six station relocations or instrumentation changes occur over 44 a century in a randomly selected US climate station (Mitchell Jr. 1953; Menne and Williams Jr. 45 2009). As such, a changepoint analysis of a climate series is often a worthy initial exploratory 46 endeavor. 47

Statistical methods to detect changepoints have rapidly evolved over the last few decades. These
include methods to detect a single shift in the series' mean (Chernoff and Zacks 1964), in its
variance (Hsu 1977), or in a general linear regression model (Quandt 1958; Robbins et al. 2016).
In the climate literature, changepoint detection has most often been used to detect mean shifts.

3

However, this may result in misinterpreting a long-term climate trend as a sequence of mean shifts
 that follows (approximates) the trend (Beaulieu and Killick 2018).

Much of the changepoint literature assumes independent and identically distributed model errors 54 (termed white noise here). However, climate time series are often autocorrelated, inducing memory 55 at time scales longer than the measurement frequency (Hasselmann 1976). This memory is often 56 modeled as a first-order autoregressive (AR(1)) process in climate studies (Lund et al. 2007; Robbins 57 et al. 2011; Hartmann et al. 2013). In an AR(1) model, autocorrelation geometrically decays to 58 zero with increasing time, representing one type of short-term memory. In the climate setting, it 59 is important to allow autocorrelation and mean shift model features in tandem as both can inject 60 similar run patterns into a climate series. An alternative is to use pre-whitening techniques that 61 mitigate the effects of autocorrelation (Robbins et al. 2011; Serinaldi and Kilsby 2016). Beaulieu 62 and Killick (2018), Shi et al. (2022), and Gallagher et al. (2021) show that changepoint inferences 63 can be drastically wrong if autocorrelation in a series is ignored. The memory in climate series has 64 also been modeled as a long-memory process, where autocorrelation decays as a power law (Yuan 65 et al. 2015). Long-memory processes and changepoint models can be confused as they both have 66 similar spectrums. Unfortunately, this ambiguity may lead to mislead inferences. Beaulieu et al. 67 (2020) discuss how to distinguish changepoints and long-memory in surface temperatures. 68

⁶⁹ Multiple changepoints may be present in climate series. Methods designed to detect a single ⁷⁰ changepoint have been applied iteratively to estimate multiple changepoint configurations through ⁷¹ a process known as binary segmentation (Scott and Knott 1974; Rodionov 2004). Binary seg-⁷² mentation is now known to perform poorly in multiple changepoint problems (Shi et al. 2022) ⁷³ (see Fryzlewicz (2014) for an interesting attempt to fix binary segmentation). Penalized likelihood ⁷⁴ methods, the approach taken here, were developed in Davis et al. (2006); Lu et al. (2010); Killick ⁷⁵ et al. (2012); Li and Lund (2012) and tend to perform better (Shi et al. 2022). Here, a likelihood,

which measures the goodness of the statistical model fit, is balanced against a penalty that pre-76 vents fitting too many changepoints. Penalized likelihood methods can allow for autocorrelation. 77 Bayesian approaches to the multiple changepoint problem also exist. Most of these place some 78 sort of prior distribution on the changepoint times, for instance a spike and slab prior (see Barry 79 and Hartigan (1993); Chib (1998); Fearnhead (2006); and Cappello et al. (2021) and the references 80 within). Li et al. (2019) construct an informative prior on the changepoint times from the sta-81 tion's metadata record. The references above are by no means exhaustive; indeed, the changepoint 82 literature is vastly expanding. 83

As most methodological statistics papers are not written with user comprehension in mind, the technical changepoint literature can seem impenetrable to non-statisticians, making it challenging to select an appropriate approach for the climate scientist. Compounding difficulties, Lund and Reeves (2002) and Beaulieu and Killick (2018) show that spurious changepoint inferences easily occur when prominent data features (e.g. autocorrelation, long-term trend) are ignored — the choice of model and method is critical in changepoint analyses. Indeed, changepoint techniques can produce different results when the models and assumptions are only slightly changed.

The aim of this paper is to present, through an example, a comprehensive changepoint analysis 91 of a climate series. To this end, we analyze the Central England temperature (CET) series 92 by fitting different changepoint models capable of detecting shifts in trends. We also compare 93 our changepoint fits with long-memory models. Our focus is on penalized likelihood multiple 94 changepoint techniques, enabling us to compare several models while preventing overestimation of 95 the number of changepoints. We also discuss mean shift models and how they fit data containing a 96 long-term trend such as the CET series. Emphasis is placed on implementation and interpretation 97 over the theoretical foundations of penalized likelihoods. Nonetheless, references to the formal 98 statistical literature are provided. 99

The rest of this paper proceeds as follows. The CET series used here is introduced in the next section. Section 3 then provides some rudimentary background on changepoint models, describing the penalized likelihood methods used here. The next three sections present fits of various multiple changepoint models. Results for each type of model motivate the subsequent fits. Remarks about the optimal model are made in the final section along with concluding comments.

105 2. The CET Series

The CET time-series is perhaps the longest instrumental record of surface temperatures in the 106 world, commencing in 1659 and spanning 362 years through 2020. The CET series is a benchmark 107 for European climate studies, as it is sensitive to atmospheric variability in the North Atlantic 108 (Parker et al. 1992). This record has been previously analyzed for long-term changes (Plaut et al. 109 1995; Harvey and Mills 2003; Hillebrand and Proietti 2017); however, to our knowledge, no 110 detailed changepoint analysis of it has been previously conducted. Changepoints are plausible 111 in the CET record for several reasons. First, artificial shifts near the record's onset may exist 112 when data quality was lower (Parker et al. 1992). Furthermore, an increase in the pace of climate 113 warming arising globally during the 1960s-1970s (Beaulieu and Killick 2018; Cahill et al. 2015) 114 may be present. The length of the CET record affords us the opportunity to explore a variety of 115 temperature features. 116

The CET series, available at https://www.metoffice.gov.uk/hadobs/hadcet/, was provided by the UK Met Office. Measurements commenced in 1659 and were mostly compiled by Manley (1953, 1974) until 1973, then continued and updated to 1991 in Parker et al. (1992). The series is now kept by the Hadley Centre, Met Office. The CET time series is an annual composite of 15 stations in the UK, located over a roughly triangular area bounded by Lancashire, London, and Bristol. The series is thus representative of the climate of the English Midlands.

The station locations used to form the composite series are depicted in the top graphic in Figure 123 1. The CET temperatures, presented in the bottom graphic of Figure 1, have been previously 124 adjusted for inhomogeneities due to changes in measurement practices through time (Manley 1953, 125 1974; Parker et al. 1992), and for urban warming since 1960 (Parker and Horton 2005). However, 126 until 1722, available instrumental records used in the CET time series did not overlap. As such, 127 non-instrumental weather diaries and the Utrecht instrumental series were used to adjust the CET 128 series and fill the gaps (Parker et al. 1992). Between 1722 and 1760, there are no gaps in the 129 composite record of all stations, but observations were generally collected in unheated rooms as 130 opposed to outdoors. A few outdoor temperature measurements were collected and used to estab-131 lish relationships between temperatures in unheated rooms and outdoors. These relationships were 132 then used to adjust the CET time series (Parker et al. 1992). The daily CET time series starts in 133 1772, and has been used to update the monthly series (Parker et al., 1992). As such, some authors 134 use only the data post-1772 for their analyses (Hillebrand and Projetti 2017). In this paper, we 135 conduct a changepoint analysis on both the full CET time series (1659-2020) and the truncated 136 series (1772-2020) that excludes the poorer quality data at the beginning of the record. 137

3. Structural Change Models

To explore structural changes in the CET series, a hierarchical changepoint analysis, gradually building on past findings, will be conducted. Let X_t denote the annual temperature observed at time *t* and suppose that data from the years 1,..., *N* are available. In general, a changepoint analysis partitions the series into *m* + 1 distinct regimes, each regime having homogeneous characteristics. The number of changepoints *m* is unknown and needs to be estimated from the series. Let τ_i denote the *i*th changepoint time; boundary conditions take $\tau_0 = 0$ and $\tau_{m+1} = N$. All regression models in this paper have the time series regression form

$$X_t = f(t) + \epsilon_t, \qquad t = 1, 2, \dots, N, \tag{1}$$

where $f(t) = E[X_t]$ is the mean of the series at time *t*. The structural form of *f* will vary, generally containing location and/or trend parameters and their shifts; each model form will be discussed as we proceed. The model errors $\{\epsilon_t\}_{t=1}^N$ have zero mean and may be correlated in time. We work with AR(1) errors for simplicity, but more complex time series models are possible. While it is important to allow for autocorrelation in annual data, the form of the correlation structure is typically not as crucial as its presence.

The AR(1) difference equation governing the errors $\{\epsilon_t\}$ is

$$\epsilon_t = \phi \epsilon_{t-1} + Z_t,$$

where $\phi \in (-1, 1)$ and $\{Z_t\}$ is zero mean white noise (WN) with unknown variance σ^2 . Solutions to the AR(1) equation have exponentially decaying correlations: $\operatorname{Corr}(\epsilon_t, \epsilon_{t+h}) = \phi^h$ for $h \ge 0$. Because the data are annually averaged, Gaussian distributed errors $\{\epsilon_t\}$ are statistically realistic. An implication of this is that future model likelihood functions will be Gaussian based.

Methods for handling multiple changepoint analyses without penalized likelihoods exist. One 157 popular technique is termed binary segmentation (Scott and Knott 1974). Binary segmentation 158 works with any single changepoint technique, termed an at most one change (AMOC) method. 159 Many AMOC tests have been developed, including cumulative sums (CUSUM) (Page 1954), 160 likelihood ratios (Jandhyala et al. 2013), Chow tests (Chow 1960), and sum of squared CUSUM 161 tests (Shi et al. 2022). Binary segmentation first analyzes the entire series for a changepoint. 162 If a changepoint is found, the series is split into subsegments about the identified changepoint 163 time and the two subsegments are further scrutinized for additional changepoints. The procedure 164 is repeated iteratively until no subsegments are deemed to have changepoints. While simple 165

and computationally convenient, binary segmentation is one of the poorer performing multiple 166 changepoint techniques (Shi et al. 2022), often being fooled by changepoints that occur close 167 to one another or multiple shifts that move the series in opposite directions. There have been 168 attempts to fix binary segmentation — see the wild binary segmentation and related methods 169 in Fryzlewicz (2014) and Eichinger and Kirch (2018). Unfortunately, these techniques typically 170 assume independent model errors or are restricted to single parameter changes per regime (for 171 example, mean shifts only). Perhaps worse, wild binary segmentation tends to overestimate 172 changepoint numbers when they are in truth infrequent (Lund and Shi 2020). 173

To estimate the changepoint structure and model parameters from the data, penalized likelihood methods will be used. Likelihood methods choose the model parameters that make seeing the observed data most likely; a penalty is imposed on the changepoint configuration to keep the fitted model parsimonious (from having too many changepoints). Our penalized likelihoods have the following form

$$-2\log(L^*(m;\tau_1,...,\tau_m)) + P(m;\tau_1,...,\tau_m).$$
 (2)

The notation here is as follows: $L^*(m; \tau_1, ..., \tau_m)$ is the optimal Gaussian likelihood that can be 179 achieved from a model with m changepoints that occur at the times τ_1, \ldots, τ_m . Here, the data 180 sample X_1, X_2, \ldots, X_N is regarded as fixed. To determine $L^*(m; \tau_1, \ldots, \tau_m)$, one must estimate all 181 parameters in the mean function f and the AR(1) model errors assuming that m changepoints 182 occur at the times τ_1, \ldots, τ_m . This procedure will be discussed further below. The quantity 183 $P(m;\tau_1,\ldots,\tau_m)$ is the penalty for having a model with *m* changepoints at the times τ_1,\ldots,τ_m . As 184 more and more changepoints are added to the model, the overall fit gets better $(-2\log(L^*))$ gets 185 smaller); the penalty, which is positive and increases with the number of changepoints, prevents 186 an overfitted model (one with too many changepoints). 187

Many penalty structures have been proposed in the statistics and climate literature. These include 188 the Akaike information criterion (AIC), the Bayesian information criterion (BIC), the modified 189 Bayesian information criterion (mBIC), and Minimum description lengths (MDL). We will use 190 BIC and MDL here. These two penalties were judged as "winners" in a recent changepoint 191 detection comparison in Shi et al. (2022). AIC penalties are not considered here because they often 192 erroneously estimate an excessive number of changepoints (Shi et al. 2022). The BIC penalty for 193 having m changepoints at the times τ_1, \ldots, τ_m is $m \log(N)$ and is proportional to the number of 194 changepoints; additional parameters are penalized at the rate of log(N) per model parameter. Our 195 penalized likelihood objective functions for structural changes are summarized in Table 1. The 196 individual models will be explained in subsequent sections. The boxed quantities are the model 197 penalties. When m = 0, penalties for any changepoint quantities are taken as zero since changepoint 198 features are absent from the model. 199

²⁰⁰ When comparing models via BIC (or any other model selection criterion), one computes the BIC ²⁰¹ statistic for all fitted models and chooses the one with the smallest BIC score. Differences between ²⁰² BIC values can give a sense of uncertainty between different model fits. The "posterior model ²⁰³ probabilities" of Burnham and Anderson (2004) can further highlight differences. Elaborating, ²⁰⁴ we label the compared models as g_i , (i = 1, ..., R) and let ΔBIC_i denote the difference between ²⁰⁵ the BIC score of model g_i and the model having the smallest BIC score. The posterior model ²⁰⁶ probabilities of Burnham and Anderson (2004) are

$$p_i = \frac{exp(-\Delta BIC_i/2)}{\sum_{r=1}^{R} exp(-\Delta BIC_r/2)}.$$
(3)

Then p_i is the inferred probability that model g_i is the quasi-true model in the model set under a prior where all *R* models are equally likely (prior probabilities are 1/R for each model). These BIC posterior model probabilities highlight uncertainties in our model comparisons. In contrast to the BIC penalty, the MDL penalty is more complex in form, also accounting for the changepoint location times τ_1, \ldots, τ_m . The MDL penalty depends on the form of f and is rooted in information theory, quantifying the computer memory needed to store the model (good fitting models use minimal space). MDL penalties have previously proven useful in changepoint detection (Davis et al. 2006; Li and Lund 2012)). Posterior model probabilities are not available for the MDL information criterion. Other penalties used in the climate literature for changepoint problems include those in Caussinus and Mestre (2004).

A drawback of penalized likelihood methods involves computation time. There are $\binom{N-1}{m}$ distinct 217 changepoint configurations having m changepoints. Summing this over all m shows that there are 218 2^{N-1} distinct changepoint configurations that need to be searched in an exhaustive optimization 219 of a penalized likelihood, a daunting task for long time series. As a solution, genetic algorithms 220 (GA) will be used to optimize our penalized likelihoods. GAs are randomized search algorithms 221 that mimic natural selection processes. In a genetic algorithm, an initial collection (generation) of 222 changepoint configurations is randomly evolved towards ones with improved penalized likelihoods. 223 Better fitting models are allowed priority in passing on their changepoints (genes) to children models 224 of the next generation. Occasionally, mutations (very different changepoint configurations) occur; 225 this keeps the GA from converging to local minimums of the penalized likelihood. Ultimately, the 226 GA converges to a model with a very good penalized likelihood. The natural selection mechanism 227 in GAs make it unlikely to visit suboptimal changepoint configurations. While Li and Lund (2012) 228 illustrate how to devise a GA in climate changepoint applications, generally available GAs have 229 now become savvy enough to capably handle our needs. The GA optimizations performed here 230 use the R package GA (Scrucca 2013). 231

In contrast to GAs, binary segmentation is a greedy algorithm that often becomes trapped at a local penalized likelihood minimum. Killick et al. (2012) and Maidstone et al. (2017b), two rapid ²³⁴ dynamic programming based multiple changepoint configuration optimizers, currently cannot ²³⁵ handle our needs: Maidstone et al. (2017b) assumes independent model errors and Killick et al. ²³⁶ (2012) assumes all parameters change at each changepoint time (including the AR(1) correlation ²³⁷ parameter ϕ and error variance σ^2). GAs are the only optimization method that reasonably handle ²³⁸ all models considered in this paper.

4. Models fitted

240 *a. Trend shift models*

We start our analysis with models having trends, as a long-term trend in the CET time series has been documented in previous studies (Kendon et al. 2021; Franzke 2012; Karoly and Stott 2006). This model posits $f(\cdot)$ to have the piece-wise linear form

$$f(t) = \begin{cases} \mu_1 + \beta_1 t, & 1 \le t \le \tau_1, \\ \mu_2 + \beta_2 t, & \tau_1 + 1 \le t \le \tau_2, \\ \vdots & & \\ \mu_{m+1} + \beta_{m+1} t, & \tau_m + 1 \le t \le N, \end{cases}$$
(4)

More compactly, one can write $E[X_t] = f(t) = \mu_{r(t)} + \beta_{r(t)}t$, where $r(t) \in \{1, 2, ..., m+1\}$ denotes the regime being used at time *t*; for example, r(t) = 1 for $1 \le t \le \tau_1$.

The changepoint literature has focused primarily on detecting mean shifts; fewer studies have been dedicated to detecting trend shifts. However, Maidstone et al. (2017a) present a dynamic programming approach that estimates trend shift configurations using a penalty based on absolute distances that is neither the MDL nor BIC. Their { ϵ_t } must be white noise (uncorrelated) with a zero mean and constant variance. See Bai and Perron (1998), Bai and Perron (2003), and their related R package strucchange by Zeileis et al. (2015) for more details. ²⁵² The least squares estimators for the *i*th regime's parameters are computed from data in this regime ²⁵³ only:

$$\hat{\beta}_{i} = \frac{\sum_{t=\tau_{i-1}+1}^{\tau_{i}} (X_{t} - \bar{X}_{i})(t - \bar{t}_{i})}{\sum_{t=\tau_{i-1}+1}^{\tau_{i}} (t - \bar{t}_{i})^{2}}, \quad \hat{\mu}_{i} = \bar{X}_{i} - \hat{\beta}_{i}\bar{t}_{i}, \quad i = 1, 2, \dots, m+1,$$
(5)

where $\bar{X}_i = (\sum_{t=\tau_{i-1}+1}^{\tau_i} X_t)/(\tau_i - \tau_{i-1})$ and $\bar{t}_i = (\tau_i + \tau_{i-1} + 1)/2$. While these are not the exact maximum likelihood estimators in correlated settings, they are typically very close to them (Lee and Lund 2012). A detailed discussion of least squares versus maximum likelihood estimator differences for time series is contained in Lee and Lund (2012).

²⁵⁸ One next computes the detrended series via

$$D_t = X_t - \hat{f}(t) = X_t - (\hat{\mu}_{r(t)} + \hat{\beta}_{r(t)}t).$$
(6)

The AR(1) parameter is then estimated via

$$\hat{\phi} = \frac{\sum_{t=1}^{N-1} D_t D_{t+1}}{\sum_{t=1}^{N} D_t^2}.$$
(7)

²⁶⁰ One-step-ahead predictions of the time series are now computed by

$$\hat{D}_t = \hat{\phi} \hat{D}_{t-1}, \quad t \ge 2, \tag{8}$$

with the start-up condition $\hat{D}_1 = 0$. The white noise variance in the AR(1) model is estimated as

$$\hat{\sigma}^2 = \frac{1}{N} \sum_{t=1}^{N} \hat{D}_t^2.$$
(9)

Plugging $\hat{\mu}_k$, $\hat{\phi}$, and $\hat{\sigma}^2$ into the Gaussian likelihood (see Li and Lund (2012) for details) gives a negative Gaussian log-likelihood of

$$-2\log(L^*(m;\tau_1,\ldots,\tau_m)) = N\log(\hat{\sigma}^2) + \underbrace{N + N\log(2\pi)}_{\text{Constant}}.$$
(10)

The underbraced constant term above does not change over distinct changepoint configurations and can be neglected in the changepoint configuration comparisons. The above equations show how to estimate model parameters and evaluate model likelihoods given the changepoint configuration;
the optimal changepoint configuration is found by a GA search. The penalized likelihoods obtained
with two different penalties, MDL and BIC, are presented in Table 1 for the various models used
here. Since regression lines are described by two parameters, all regimes are required to be at least
three years long (so that fits in any single regime are not perfect).

On the full CET series, GA optimizations of the BIC and MDL penalized likelihoods estimate 271 identical trend shift configurations, both flagging three breaks at the times 1700, 1739, and 1988 272 (Table 2). This methodological agreement is convenient, but is not typical in changepoint analyses. 273 Figure 2 graphically depicts our model fit. Cooling occurs during the first 39 years, followed by 274 an increasing-trend second regime, with subsequent shifts to two warming trend regimes. The last 275 regime, which starts in 1989, is warming with a trend of 1.1°C per century. When fitting trend 276 shift models to CET series on post 1772 data only, we find a single changepoint in 1987 (Table 3), 277 which is consistent with our analysis on the full series. 278

In both cases, the AR(1) correlation estimate is very small ($\hat{\phi} = 0.058$ for the full CET and 279 $\hat{\phi} = 0.073$ for the truncated), and is not significantly different from zero with standard time series 280 tests (Brockwell and Davis 1991). When $\phi = 0$, an AR(1) model reduces to white noise. This 281 point is worth emphasizing: our model fits prefer the trend shift structure over structures involving 282 autocorrelated errors. This is an important point since positive autocorrelation and shifts can induce 283 similar run patterns in series — likelihood methods can decide which feature (or both) is statistically 284 preferable. Should autocorrelation be neglected, one risks flagging spurious changepoints. And 285 while independent model errors is reasonable here, it may not hold in other applications, especially 286 if monthly or daily data are used. 287

Other assumptions made on the model errors include normality and a constant variance in X_t . To assess normality, we apply a Shapiro-Wilk test to the model residuals. This test does not reject ²⁹⁰ normality (Tables 2- 3) at any common levels of statistical significance. To investigate the constant ²⁹¹ variance assumption, we apply Leneve's test to the residuals. This test does not find evidence ²⁹² of a changing variance in the residuals of the trend shifts models fitted to the CET series at any ²⁹³ appreciable levels of statistical significance. Normality and constant variance assumptions in all ²⁹⁴ future fitted models (Tables 2 and 3 list these) is investigated — these features are not rejected in ²⁹⁵ any of the models compared here.

b. A fixed slope mean shift model

In some cases, it may be appropriate to constrain trends to be identical over all regimes (Wang 2003). This could be the case if artificial changes are expected. For example, a change of instrument may introduce an artificial shift in a time series, but will not necessarily alter the long-term trend in different regimes. A model with a common trend slope in all regimes (Lu and Lund 2007) is

$$f(t) = \begin{cases} \mu_1 + \beta t, & 1 \le t \le \tau_1, \\ \mu_2 + \beta t, & \tau_1 + 1 \le t \le \tau_2, \\ \vdots & & \\ \mu_{m+1} + \beta t, & \tau_m + 1 \le t \le N, \end{cases}$$
(11)

where β is the trend slope, which is the same in all regimes.

³⁰² In compact form, the model can be expressed as

$$X_t = \mu_{r(t)} + \beta t + \epsilon_t, \tag{12}$$

where $\mu_{r(t)}$ is as in (5), and $\{\epsilon_t\}$ is an AR(1) process.

The ordinary least square estimators of β and μ_1, \ldots, μ_{m+1} have the explicit form

$$\hat{\beta} = \frac{\sum_{i=1}^{m+1} \sum_{\tau_{i-1}+1}^{\tau_i} (X_t - \bar{X}_i)(t - \bar{t}_i)}{\sum_{i=1}^{m+1} \sum_{t=\tau_{i-1}+1}^{\tau_i} (t - \bar{t}_i)^2}, \qquad \hat{\mu}_i = \bar{X}_i - \hat{\beta}\bar{t}_i, \qquad i = 1, 2, \dots, m+1,$$
(13)

where \bar{X}_i and \bar{t}_i are as before. These are again very close to the maximum likelihood estimators (Lee and Lund 2012). The BIC and MDL penalties are listed in Table 1.

A GA was used to estimate this configuration, which is plotted against the data in Figure 3. For the full CET series, both BIC and MDL flag a single mean shift in 1988, while the single detected shift moves to 1990 in the truncated series (post 1772). Fewer changepoints are detected in this model than with the trend shift models of the previous section, but the time of the single change detected here is consistent with the last changepoint found in the trend shifts models. Since the BIC and MDL penalized likelihoods in Tables 2 and 3 are larger for the constant slope model than for the regime-varying trend slope model, the inference is that regime-varying slopes are preferable.

314 c. Joinpin models

There is debate over whether trend models should impose continuity in $E[X_t]$ at the changepoint times in temperature series (Rahmstorf et al. 2017). These so-called joinpin models require $E[X_t] = f(t)$ to be continuous in time *t*. Here, we compare a joinpin model to the trend shifts and fixed slope mean shift models fitted in the previous sections. Unfortunately, it is not clear what an appropriate MDL penalty is for this case, nor does this seem to be an easy matter to rectify; hence, we proceed with BIC penalties only.

To fit a joinpin model, the package in Maidstone et al. (2017a) was used. We fit the same model as (4), but with additional constraints to force continuity at the changepoint time(s). A simple way to enforce this continuity is to view the slopes as determined from $E[X_t]$ at the start and end of each regime. This enforces continuity within a simple form foregoing additional constraints. This formulation fits the model

$$X_t = \gamma_{\tau_i} + \frac{\gamma_{\tau_{i+1}} - \gamma_{\tau_i}}{\tau_{i+1} - \tau_i} (t - \tau_i) + \epsilon_t, \tag{14}$$

where γ_i is the value of the mean at time *i*. This formulation is equivalent to (4) with an additional continuity constraint at the changepoint locations. Based on Maidstone et al. (2017a), the BIC for the joinpin model is

BIC =
$$N \log(\hat{\sigma}^2) + N + N \log(2\pi) + (2m+1) \log(N),$$
 (15)

329 where

$$\hat{\sigma}^2 = \frac{1}{N} \sum_{i=1}^{m+1} \sum_{t=\tau_i}^{\tau_{i+1}} \left[X_t - \frac{\gamma_{\tau_{i+1}} - \gamma_{\tau_i}}{\tau_{i+1} - \tau_i} (t - \tau_i) \right]^2.$$

In the formulation of Maidstone et al. (2017a), the white noise variance is fixed and needs to 330 be estimated. While median absolute deviations could be used for this purpose, we instead use 331 the estimated error variance of 0.29 (Table 2), taken from the discontinuous model fits and BIC 332 penalties of the last section, This fit assumes IID errors, which seems plausible given the results of 333 the previous sections. The fitted model flags a single changepoint in 1973 in the full CET series 334 and none in the truncated series; see Tables 2-3 and Figure 4. These fits are stable against changes 335 from 0.29 in the white noise variance. Compared to our previous model fits, the joinpin model has 336 a much higher BIC than the trend shift and fixed slope mean shifts models (Tables 2-3). As such, 337 joinpin models do not appear to be competitive. 338

³³⁹While a changepoint seems plausible towards the end of the record due to an increased warming ³⁴⁰rate, the joinpin fit to the earliest data is poor, similar to the fixed slope mean shifts model. This ³⁴¹is graphically evident in the Figure 4 fits, but is also reflected by the higher BIC scores in Tables ³⁴²2-3. A joinpin model should be used when a discontinuous mean function is unlikely or physically ³⁴³implausible. With the CET series, it is not evident whether the estimated mean function should ³⁴⁴be continuous or discontinuous. Elaborating, for series containing "only a single station", mean ³⁴⁵discontinuities are physically expected. However, when more and more station records are averaged ³⁴⁶into a composite record, mean function discontinuities are reduced, becoming less pronounced with ³⁴⁷ an increasing number of stations. Should a discontinuous mean function be deemed possible, a ³⁴⁸ trend shift model provides greater flexibility since it can simultaneously approximate a joinpin ³⁴⁹ continuous structure as well as discontinuous shifts (Beaulieu and Killick 2018).

350 d. Long-memory models

A body of climate literature argues that climate time series exhibit long-memory, where the series' autocorrelation decays slowly in lag, often via a power law (Yuan et al. 2015; Blender and Fraedrich 2003; Franzke 2012). Long-memory correlation and changepoint features can inject similar run properties into a climate series, which is appreciated in the statistical and econometric literatures (Diebold and Inoue 2001; Granger and Hyung 2004; Mills 2007; Yau and Davis 2012). The daily CET series may exhibit long-memory (Syroka and Toumi 2001; Franzke 2012).

To compare our changepoint models to a long-memory model, we fit an autoregressive fractionally integrated moving-average (ARFIMA) model to the CET series. In particular, ARFIMA models with no moving-average component, an integration parameter d with 0 < d < 0.5, and an autoregressive component of orders zero and one, are considered. The AR(1) long-memory model is characterized as

$$X_t = (1 - B)^d (1 - \phi B)^{-1} \epsilon_t, \tag{16}$$

where *B* is the backshift operator applied to X_t .

To fit ARFIMA models, the R package fracdiff (Maechler 2020) was used. A BIC penalty was calculated and is listed in Table 1. An MDL penalty is not informative since this model does not have any changepoints. Long-memory model fits to the full and truncated CET series are described in Tables 2-3). The long-memory models have the largest BIC score among all models compared on the full CET time series. On the truncated series, they are also amongst the least plausible,

although joinpin models have higher BIC scores. These results suggest that changepoints, rather 368 than long-memory, are more plausible in the CET series. For additional evidence that changepoints 369 are preferred over long-memory features, we applied the time varying wavelet spectrum methods in 370 Norwood and Killick (2018) to the CET series. These methods were used on surface temperatures 371 in Beaulieu et al. (2020) and shown to discriminate changepoint and long-memory models well in 372 long series. The results confirm that a changepoint model is more appropriate than a long-memory 373 model. The fitted model of autoregressive order zero was also preferred to the fitted model of order 374 one, reinforcing that correlation aspects in the CET series are minimal. 375

³⁷⁶ e. Model selection uncertainty

Among the six models compared, the trend shift model with white noise is judged the most 377 plausible, as suggested by both BIC and MDL scores. The BIC posterior probabilities for all 378 models fitted above are presented in Table 4. For the full series, the model probability for the 379 trend shift model with white noise is 0.64, followed by the joinpin model with probability 0.12 and 380 the trend shift model with AR(1) errors with probability 0.11. The three other models all have a 381 posterior probability of 0.05 or less. This highlights the uncertainty in the model selected, although 382 the trend shifts models with AR(1) and white noise errors are very similar (the autocorrelation 383 estimated in the AR(1) model is small and both configurations identify the same shifts). As for the 384 joinpin model, the fit at the start of the record seems poor. 385

Moving to the truncated series, the trend shift model with white noise has a posterior probability of 0.68. The next most plausible models are the fixed slope mean shift model with AR(1) errors and the trend shift model with AR(1) errors, having posterior probabilities of 0.1 and 0.09, respectively (Table 4). These models are similar in that estimated changepoint times are very close, giving further evidence for a shift in the late 1980s. However, this suggests that a fixed slope model should ³⁹¹ not be entirely discarded. Unlike results for the full CET series, the joinpin model ranks very low ³⁹² (0.02) on the truncated CET series. This is not surprising given that no changepoint is detected ³⁹³ under the joinpin model in the truncated series (Figure 4).

5. Trends vs Mean Shifts

The simplest changepoint analysis is arguably that of mean shifts. This is the most common model in the changepoint literature and has been widely used to analyze climate series. While this structure is inappropriate for series having trends (such as the CET analyzed here), we include this model here for comparative purposes. The mean shifts model posits $f(\cdot)$ to have form

$$f(t) = \begin{cases} \mu_1, & 1 \le t \le \tau_1, \\ \mu_2, & \tau_1 + 1 \le t \le \tau_2, \\ \vdots \\ \mu_{m+1}, & \tau_m + 1 \le t \le N. \end{cases}$$
(17)

The model's mean structure is compactly written as $f(t) = E[X_t] = \mu_{r(t)}$, where $r(t) \in \{1, 2, ..., m + 1\}$ denotes the regime being used at time *t*; for example, r(t) = 1 for $1 \le t \le \tau_1$. Given *m* and the changepoint times $\tau_1, ..., \tau_m$, mean parameters are first estimated via segment

Given *m* and the changepoint times τ_1, \ldots, τ_m , mean parameters are first estimated via segment averages:

$$\hat{\mu}_i = \frac{1}{\tau_i - \tau_{i-1}} \sum_{t=\tau_{i-1}+1}^{\tau_i} X_t, \qquad i = 1, 2, \dots, m+1.$$
(18)

While sample means are not the exact maximum likelihood estimators of the mean parameters for correlated series, they are typically very close and are easy to compute (unlike maximum likelihood estimators). Next, the regime-wise mean estimated in (18) is subtracted from the series by computing $D_t = X_t - \hat{f}(t) = X_t - \hat{\mu}_{r(t)}$. The variance $\hat{\sigma}^2$ is then estimated as in (9). We do not fit this model with AR(1) errors based on the results from the previous sections. The BIC and MDL ⁴⁰⁸ penalized likelihoods for this model are

409

BIC =
$$N \log(\hat{\sigma}^2) + N + N \log(2\pi) + (3m+3) \log(N);$$
 (19)

$$MDL = N\log(\hat{\sigma}^2) + N + N\log(2\pi) + \log(N) + 2\log(m) + 2\sum_{i=1}^{m+1}\log(\tau_i - \tau_{i-1}) + 2\sum_{i=2}^{m+1}\log(\tau_i).$$
(20)

We discuss only results on the full series here, but conclusions are consistent (i.e., the same changepoints are detected post 1772) if we repeat the analysis on the truncated series only. Fitting this model, seven changepoints are flagged with both MDL and BIC (Figure 5).

Both penalties pinpoint 1989 as a changepoint time, which is consistent with results of the previous section. Here, MDL and BIC both deem the "cold year" in 1740 an outlier, bracketing this time by two changepoints. Because MDL methods are based on information theory (Rissanen 1978) and not large sample statistical asymptotics, they often flag outliers. Shifts are more frequent at the beginning of the record, perhaps suggesting that the data during these times is less reliable. Evident in the fits is that the last three regimes act to move the series higher in a "staircase", which is expected for a series experiencing a long-term warming trend (Figure 5).

The BIC and MDL scores obtained on the full CET series are 648.17 and 656.09, respectively. 420 Should this model be included in our main comparison, one would still prefer the trend shift model 421 should the MDL penalty be used to make conclusions. However, the BIC mean shift score is 422 smaller than the BIC trend shift score in the previous section, indicating preference for the mean 423 shift model. A model containing only mean shifts will flag a sequence of shifts in an attempt to 424 follow a long-term trend should the data have a trend and it not be included in the model. If the 425 trend is not steep, as is the case here, it is especially challenging to distinguish between trends 426 and mean shifts. To illustrate this, we conducted a simulation study where 500 synthetic series 427 with the same trend magnitude and variability (as estimated in the truncated CET time series over 428 1772-2020) were generated. The mean shifts plus white noise and trend shifts plus white noise 429

⁴³⁰ models were fitted to each series. In only 18% of the synthetic series, the correct model with ⁴³¹ a long-term trend was selected by BIC. Figure 6 presents a histogram of the difference between ⁴³² the two fitted models' BIC scores, further demonstrating the bias BIC has for the erroneous mean ⁴³³ shifts model. Should there be any suspicion about a trend or "staircase feature" in the record, we ⁴³⁴ recommend using techniques that incorporate trends, as done here.

6. Comments, Conclusions, and Discussion

This study compared and contrasted several common changepoint model fits for data containing trends, as well as a long-memory autocovariance model, to the CET time series. To our knowledge, this is the first time a detailed changepoint analysis has been conducted on this long record. Starting with a trend shift model, several different changepoint structures were fitted, illustrating the techniques and salient points of changepoint analyses.

Tables 2-3 present the log-likelihood, BIC, and MDL scores of all model fits. Depending on the model configuration, we detect either three changepoints (trend shifts models) or one changepoint (fixed slope mean shifts and joinpin models) in the full series. This changepoint count discrepancy traces to the large variations in the series during roughly the first century of the record.

Most models agree on a change to a rapidly warming regime circa 1988, except for the joinpin 445 model (this is also true for the truncated series). Among all fitted models, the optimal one has trend 446 shifts in 1700, 1739, and 1988 (full series), and one in 1988 (truncated series). Table 5 provides 447 estimates of the best fitting model's intercept and slope parameters by regime. While the best fitting 448 model is the trend shifts model, other models are also plausible (Table 4). Models with higher 449 posterior probabilities tend to be consistent in their flagged changepoint times, but highlight that a 450 fixed slope model (as opposed to the varying slopes in the trend shifts models) may be plausible. 451 Long-memory models yield the highest BIC scores, and are less plausible than all other models 452

compared. The results of the full and truncated CET series are consistent, showing that our post
1772 changepoint inferences are not overly sensitive to inclusion of the first century of the series.
Having both BIC and MDL penalties agree on the model type and changepoint configuration
adds robustness to our conclusions, suggesting that the fitted segmentations are stable. According
to Lavielle (2005), changepoint segmentations that are stable over a range of penalty values should
be preferred. Overall, models with shifts were deemed preferable to models having autocorrelated
errors.

While our aim is not necessarily directed to the causes of the detected shifts, we provide some 460 interpretations here. Shifts flagged during the first century of the record are likely due to inferior 461 data quality over this early period (Hillebrand and Projetti 2017). Due to lack of overlapping 462 instrumentation coverage before 1722, non-instrumental weather diaries were used to adjust the 463 series (Parker et al. 1992). Observations were generally collected in unheated rooms until 1760, 464 and adjusted by calibrating indoor and outdoor observations later (Parker et al. 1992). Even with 465 the most careful adjustments, one cannot guarantee that all biases were removed from the data. 466 Some authors omit the first century of data altogether due to this issue (Hillebrand and Proietti 467 2017). 468

The trend shifts model on the earlier part of the data detects two changepoints in 1700 and 1739, 469 characterizing a steep cooling trend followed by a warming trend. The mean shifts model fitted 470 on the earlier part of the data flags multiple changes (1691, 1699, 1727, 1740, 1741), calling for a 471 closer examination of the earlier part of the record. In data with inhomogeneities, BIC penalties 472 favor mean shift models over trend shift models, even if the trend shifts model is truth. A mean 473 shift model characterizes a warming trend as a staircase of increasing steps. This issue can be 474 troublesome if the trend in the data is weak, as demonstrated in our simulation study (see Figure 475 6). 476

23

The changepoint flagged in 1988 (from multiple models and in both the full and truncated CET series) is not surprising given the warming seen on the global level in the 1960/70s in a range of surface temperature records, as discussed in studies using both trend shift and joinpin models (Cahill et al. 2015; Beaulieu and Killick 2018; Rahmstorf et al. 2017; Ruggieri 2013). While the more recent part of the CET series is considered more reliable and has been adjusted for inhomogeneities, we cannot entirely discard issues in this era either. Overall, it is possible that a combination of natural and artificial causes contribute to shifts in the CET series.

To further rule out artificial changes, one could subtract all $\binom{15}{2} = 105$ pairs of series from one another and examine these differences for changepoints. Then, one can distinguish artificially caused changepoints from those due to natural climate change and variability. See Menne and Williams Jr. (2009) for more details on this procedure. Artificial changes can then be corrected before long-term trends are analyzed. Changes that are not considered artificial can be further investigated through an attribution study (Hartmann et al. 2013).

Residual analyses were conducted to ensure that the underlying assumptions of the model were 490 met. With the CET series, residuals of the trend shift model fit were judged to be uncorrelated 491 (white noise). However, climate time series often exhibit autocorrelation that should be taken into 492 account. We stress the importance of verifying the underlying assumptions in any changepoint 493 model. Indeed, neglecting positive autocorrelation raises the risk of detecting spurious shifts. 494 Also, the series' autocorrelation may be more complex than an AR(1) process and may itself 495 contain shifts (Beaulieu et al. 2012; Beaulieu and Killick 2018). Some climate series may also 496 contain long-memory autocorrelations (Vyushin et al. 2012). An additional challenge lies with 497 the ambiguity between long-memory and changepoint models: both features can produce series 498 with similar run structures. Because of this, a long-memory model was included as part of our 499 comparison. We found that the CET time series is best represented by a multiple trend shift 500

changepoint structure and not a long-memory model. Such a comparison is not possible for all climate series since lengthy records are required to analyze long-memory series (Beaulieu et al. 2020). The CET time series, which is the longest publicly available surface temperature series, enables this comparison. Other assumptions that were made include constant variance temperatures and normally distributed observations. Both assumptions cannot be rejected in any models fitted (Tables 2-3).

Model selection based on a criteria does not guarantee that the selected model is "truth". All 507 models are an approximation of reality and multiple models can plausibly represent the data. To 508 quantify this, one can calculate posterior model probabilities with BIC that each fitted model is 509 the "quasi-truth". This assumes that all models included in the comparison have the same prior 510 weight, which may not be reasonable. One must also note that this measure is relative to the 511 models included in the comparison, and does not reflect the uncertainty that the "true" model may 512 not be part of the model set. Similarly, uncertainty in the total number of changepoints and their 513 individual occurrence times is a difficult statistics problem. Bayesian methods, which were not 514 considered here, can in principle place uncertainty margins on the number of changepoints and 515 their locations. When several distinct models have similar penalized likelihood scores, inferences 516 about the number of changepoints are likely to be less reliable. Recent statistics work is now 517 studying this issue (Li et al. 2019; Cappello et al. 2021). 518

⁵¹⁹ Ultimately the choice of "best model" should be arrived at from a judgment made by the ⁵²⁰ researcher(s) based on objective statistical metrics, such as presented in this work, combined with ⁵²¹ understanding of the data recording practices and physics of the natural system.

25

Acknowledgments. Rebecca Killick gratefully acknowledges funding from EP/R01860X/1 and NE/T006102/1. Robert Lund and Xueheng Shi thank funding from NSF DMS-2113592. The comments of three referees and the editor substantially improved this manuscript.

⁵²⁵ Data availability statement. The Central England data used in this study is available at https: ⁵²⁶ //www.metoffice.gov.uk/hadobs/hadcet/. We used the annual means from 1659-2020.

527 **References**

- Bai, J., and P. Perron, 1998: Estimating and testing linear models with multiple structural changes.
 Econometrica, 66, 47–78.
- Bai, J., and P. Perron, 2003: Computation and analysis of multiple structural change models.
 Journal of Applied Econometrics, 18 (1), 1–22.
- ⁵³² Barry, D., and J. A. Hartigan, 1993: A Bayesian analysis for change point problems. *Journal of* ⁵³³ *the American Statistical Association*, **88 (421)**, 309–319.
- Beaulieu, C., J. Chen, and J. L. Sarmiento, 2012: Change-point analysis as a tool to detect abrupt
 climate variations. *Philosophical Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences*, 370 (1962), 1228–1249, doi:10.1098/rsta.2011.0383.
- Beaulieu, C., and R. Killick, 2018: Distinguishing trends and shifts from memory in climate data.
 Journal of Climate, **31 (23)**, 9519 9543.
- Beaulieu, C., R. Killick, D. Ireland, and B. Norwood, 2020: Considering long-memory when
 testing for changepoints in surface temperature: A classification approach based on the time varying spectrum. *Environmetrics*, **31** (1), e2568.

- ⁵⁴² Blender, R., and K. Fraedrich, 2003: Long time memory in global warming simulations. *Geophys*-⁵⁴³ *ical Research Letters*, **30** (14).
- Brockwell, P. J., and R. A. Davis, 1991: *Time Series: Theory and Methods*. 2nd ed., SpringerVerlag.
- ⁵⁴⁶ Burnham, K. P., and D. R. Anderson, 2004: Multimodel inference: Understanding AIC and ⁵⁴⁷ BIC in model selection. *Sociological Methods & Research*, **33** (2), 261–304, doi:10.1177/ ⁵⁴⁸ 0049124104268644.
- ⁵⁴⁹ Cahill, N., S. Rahmstorf, and A. C. Parnell, 2015: Change points of global temperature. *Environ- mental Research Letters*, **10 (8)**, 084 002.
- ⁵⁵¹ Cappello, L., O. H. M. Padilla, and J. A. Palacios, 2021: Scalable Bayesian change point detection
 ⁵⁵² with spike and slab priors. *arXiv preprint arXiv:2106.10383*.
- ⁵⁵³ Caussinus, H., and O. Mestre, 2004: Detection and correction of artificial shifts in climate series.
- Journal of the Royal Statistical Society: Series C (Applied Statistics), **53** (**3**), 405–425.
- ⁵⁵⁵ Chernoff, H., and S. Zacks, 1964: Estimating the current mean of a normal distribution which is ⁵⁵⁶ subjected to changes in time. *The Annals of Mathematical Statistics*, **35** (**3**), 999–1018.
- ⁵⁵⁷ Chib, S., 1998: Estimation and comparison of multiple change-point models. *Journal of Econo-*⁵⁵⁸ *metrics*, **86 (2)**, 221–241.
- ⁵⁵⁹ Chow, G. C., 1960: Tests of equality between sets of coefficients in two linear regressions. *Econometrica: Journal of the Econometric Society*, **28**, 591–605.
- ⁵⁶¹ Davis, R. A., T. C. M. Lee, and G. A. Rodriguez-Yam, 2006: Structural break estimation for ⁵⁶² nonstationary time series models. *Journal of the American Statistical Association*, **101** (**473**), ⁵⁶³ 223–239.

- ⁵⁶⁴ Diebold, F. X., and A. Inoue, 2001: Long memory and regime switching. *Journal of Econometrics*, ⁵⁶⁵ **105** (1), 131–159.
- Eichinger, B., and C. Kirch, 2018: A MOSUM procedure for the estimation of multiple random change points. *Bernoulli*, **24** (1), 526–564.
- Fearnhead, P., 2006: Exact and efficient Bayesian inference for multiple changepoint problems.
 Statistics and Computing, 16 (2), 203–213.
- Franzke, C., 2012: Nonlinear trends, long-range dependence, and climate noise properties of
 surface temperature. *Journal of Climate*, 25 (12), 4172–4183, URL http://www.jstor.org/stable/
 26191996.
- Fryzlewicz, P., 2014: Wild binary segmentation for multiple change-point detection. *Annals of Statistics*, 42 (6), 2243–2281.
- ⁵⁷⁵ Gallagher, C. M., R. Killick, R. Lund, and X. Shi, 2021: Autocovariance estimation in the presence
 ⁵⁷⁶ of changepoints. *arXiv preprint arXiv:2102.10669*.
- Granger, C. W. J., and N. Hyung, 2004: Occasional structural breaks and long memory with an application to the S&P 500 absolute stock returns. *Journal of Empirical Finance*, **11 (3)**, 399–421.
- Hartmann, D., and Coauthors, 2013: *Observations: Atmosphere and Surface*, book section 2,
 159–254. Cambridge University Press, Cambridge, United Kingdom.
- Harvey, D. I., and T. C. Mills, 2003: Modelling trends in Central England temperatures. *Journal* of Forecasting, 22 (1), 35–47.
- Hasselmann, K., 1976: Stochastic climate models part I. Theory. *Tellus*, **28** (6), 473–485.

- Hewaarachchi, A. P., Y. Li, R. Lund, and J. Rennie, 2017: Homogenization of daily temperature
 data. *Journal of Climate*, **30** (**3**), 985–999.
- ⁵⁸⁷ Hillebrand, E., and T. Proietti, 2017: Phase changes and seasonal warming in early instrumental
 ⁵⁸⁸ temperature records. *Journal of Climate*, **30** (17), 6795–6821.
- Hsu, D.-A., 1977: Tests for variance shift at an unknown time point. *Journal of the Royal Statistical Society: Series C*, 26 (3), 279–284.
- Jandhyala, V., S. Fotopoulos, I. MacNeill, and P. Liu, 2013: Inference for single and multiple change-points in time series. *Journal of Time Series Analysis*, **34** (**4**), 423–446.
- Karoly, D. J., and P. A. Stott, 2006: Anthropogenic warming of Central England temperature.
 Atmospheric Science Letters, 7 (4), 81–85.
- Kendon, M., M. McCarthy, S. Jevrejeva, A. Matthews, T. Sparks, and J. Garforth, 2021: State of
- the UK climate 2020. International Journal of Climatology, **41** (S2), 1–76, doi:https://doi.org/
- ⁵⁹⁷ 10.1002/joc.7285, URL https://rmets.onlinelibrary.wiley.com/doi/abs/10.1002/joc.7285, https:

//rmets.onlinelibrary.wiley.com/doi/pdf/10.1002/joc.7285.

- Killick, R., P. Fearnhead, and I. A. Eckley, 2012: Optimal detection of changepoints with a linear
 computational cost. *Journal of the American Statistical Association*, **107** (**500**), 1590–1598.
- Lavielle, M., 2005: Using penalized contrasts for the change-point problem. *Signal Processing*, **85 (8)**, 1501–1510.
- Lee, J., and R. Lund, 2012: A refined efficiency rate for ordinary least squares and generalized least
 squares estimators for a linear trend with autoregressive errors. *Journal of Time Series Analysis*,
 33 (2), 312–324.

- ⁶⁰⁶ Li, S., and R. Lund, 2012: Multiple changepoint detection via genetic algorithms. *Journal of* ⁶⁰⁷ *Climate*, **25** (2), 674–686.
- ⁶⁰⁸ Li, Y., R. Lund, and A. Hewaarachchi, 2019: Multiple changepoint detection with partial informa-⁶⁰⁹ tion on changepoint times. *Electronic Journal of Statistics*, **13** (**2**), 2462–2520.
- ⁶¹⁰ Lu, Q., and R. Lund, 2007: Simple linear regression with multiple level shifts. *Canadian Journal* ⁶¹¹ *of Statistics*, **35 (3)**, 447–458.
- ⁶¹² Lu, Q., R. Lund, and T. C. M. Lee, 2010: An MDL approach to the climate segmentation problem. ⁶¹³ *Annals of Applied Statistics*, **4**, 299–319.
- ⁶¹⁴ Lund, R., and J. Reeves, 2002: Detection of undocumented changepoints: A revision of the ⁶¹⁵ two-phase regression model. *Journal of Climate*, **15** (**17**), 2547–2554.
- Lund, R., and X. Shi, 2020: Comments on "Detecting possibly frequent change-points: wild binary
 segmentation 2 and steepest-drop model selection". *Journal of the Korean Statistical Society*, **49** (4), 1090–1095.
- ⁶¹⁹ Lund, R., X. L. Wang, Q. Q. Lu, J. Reeves, C. M. Gallagher, and Y. Feng, 2007: Changepoint ⁶²⁰ detection in periodic and autocorrelated time series. *Journal of Climate*, **20** (**20**), 5178–5190.

⁶²¹ Maechler, M., 2020: fracdiff: Fractionally Differenced ARIMA aka ARFIMA(p,d,q) Models. URL

https://CRAN.R-project.org/package=fracdiff, R package version 1.5-1.

- Maidstone, R., P. Fearnhead, and A. Letchford, 2017a: Detecting changes in slope with an L_0 penalty. *Journal of Computational and Graphical Statistics*, **28**, 265–275.
- ⁶²⁵ Maidstone, R., T. Hocking, G. Rigaill, and P. Fearnhead, 2017b: On optimal multiple changepoint ⁶²⁶ algorithms for large data. *Statistics and Computing*, **27** (**2**), 519–533.

- Manley, G., 1953: The mean temperature of Central England, 1698–1952. *Quarterly Journal of the Royal Meteorological Society*, **79 (340)**, 242–261.
- Manley, G., 1974: Central England temperatures: monthly means 1659 to 1973. *Quarterly Journal* of the Royal Meteorological Society, **100** (**425**), 389–405.
- ⁶³¹ Menne, M. J., and C. N. Williams Jr., 2009: Homogenization of temperature series via pairwise ⁶³² comparisons. *Journal of Climate*, **22** (**7**), 1700–1717.
- Mills, T. C., 2007: Time series modelling of two millennia of Northern Hemisphere temperatures:
- long memory or shifting trends? Journal of the Royal Statistical Society: Series A (Statistics in
- 635 *Society*), **170** (**1**), 83–94.
- ⁶³⁶ Mitchell Jr., J. M., 1953: On the causes of instrumentally observed secular temperature trends. ⁶³⁷ *Journal of Atmospheric Sciences*, **10** (**4**), 244–261.
- Norwood, B., and R. Killick, 2018: Long memory and changepoint models: a spectral classification
 procedure. *Statistics and Computing*, 28 (2), 291–302.
- Page, E. S., 1954: Continuous inspection schemes. *Biometrika*, **41** (1), 100–115.
- Parker, D., and B. Horton, 2005: Uncertainties in central england temperature 1878–2003 and
 some improvements to the maximum and minimum series. *International Journal of Climatology*,
 25 (9), 1173–1188.
- Parker, D. E., T. P. Legg, and C. K. Folland, 1992: A new daily central england temperature series,
 1772–1991. *International Journal of Climatology*, **12** (**4**), 317–342.
- Plaut, G., M. Ghil, and R. Vautard, 1995: Interannual and interdecadal variability in 335 years of
- ⁶⁴⁷ Central England temperatures. *Science*, **268** (**5211**), 710–713.

- ⁶⁴⁸ Quandt, R. E., 1958: The estimation of the parameters of a linear regression system obeying two ⁶⁴⁹ separate regimes. *Journal of the American Statistical Association*, **53** (**284**), 873–880.
- Rahmstorf, S., G. Foster, and N. Cahill, 2017: Global temperature evolution: recent trends and
- some pitfalls. *Environmental Research Letters*, **12** (**5**), 054 001.
- Rissanen, J., 1978: Modeling by shortest data description. Automatica, 14 (5), 465–471.
- Robbins, M., C. Gallagher, R. Lund, and A. Aue, 2011: Mean shift testing in correlated data.
 Journal of Time Series Analysis, 32 (5), 498–511.
- Robbins, M. W., C. M. Gallagher, and R. Lund, 2016: A general regression changepoint test for
 time series data. *Journal of the American Statistical Association*, **111 (514)**, 670–683.
- ⁶⁵⁷ Rodionov, S. N., 2004: A sequential algorithm for testing climate regime shifts. *Geophysical* ⁶⁵⁸ *Research Letters*, **31** (9), doi:10.1029/2004GL019448.
- ⁶⁵⁹ Ruggieri, E., 2013: A Bayesian approach to detecting change points in climatic records. *Interna-* ⁶⁶⁰ *tional Journal of Climatology*, **33 (2)**, 520–528.
- Scott, A. J., and M. Knott, 1974: A cluster analysis method for grouping means in the analysis of
 variance. *Biometrics*, **30**, 507–512.
- Scrucca, L., 2013: GA: a package for genetic algorithms in R. *Journal of Statistical Software*,
 53 (4), 1–37.
- Serinaldi, F., and C. G. Kilsby, 2016: The importance of prewhitening in change point analysis
- ⁶⁶⁶ under persistence. *Stochastic Environmental Research and Risk Assessment*, **30** (2), 763–777.

- ⁶⁶⁷ Shi, X., C. Gallagher, R. Lund, and R. Killick, 2022: A comparison of single and multiple ⁶⁶⁸ changepoint techniques for time series data. *Computational Statistics & Data Analysis*, **170**, ⁶⁶⁹ 107 433.
- Syroka, J., and R. Toumi, 2001: Scaling of Central England temperature fluctuations? *Atmospheric Science Letters*, 2 (1), 143–154, doi:https://doi.org/10.1006/asle.2002.0047.
- ⁶⁷² Trewin, B., and Coauthors, 2020: An updated long-term homogenized daily temperature data set ⁶⁷³ for Australia. *Geoscience Data Journal*, **7** (**2**), 149–169.
- ⁶⁷⁴ Vincent, L. A., M. M. Hartwell, and X. L. Wang, 2020: A third generation of homogenized
- temperature for trend analysis and monitoring changes in Canada's climate. *Atmosphere-Ocean*,
 58 (3), 173–191.
- ⁶⁷⁷ Vyushin, D. I., P. J. Kushner, and F. Zwiers, 2012: Modeling and understanding persistence of ⁶⁷⁸ climate variability. *Journal of Geophysical Research: Atmospheres*, **117** (**D21**).
- ⁶⁷⁹ Wang, X. L., 2003: Comments on "Detection of undocumented changepoints: A revision of the ⁶⁸⁰ two-phase regression model". *Journal of Climate*, **16** (**20**), 3383–3385.
- Yau, C. Y., and R. A. Davis, 2012: Likelihood inference for discriminating between long-memory
 and change-point models. *Journal of Time Series Analysis*, 33 (4), 649–664.
- Yuan, N., M. Ding, Y. Huang, Z. Fu, E. Xoplaki, and J. Luterbacher, 2015: On the long-term
 climate memory in the surface air temperature records over Antarctica: A nonnegligible factor
 for trend evaluation. *Journal of Climate*, 28 (15), 5922–5934.
- Zeileis, A., F. Leisch, K. Hornik, C. Kleiber, B. Hansen, E. C. Merkle, and M. A. Zeileis, 2015:
- Package 'strucchange'. *R package version*, 1–5.

688 LIST OF TABLES

689 690 691 692	Table 1.	Penalized likelihoods. The boxed terms are the penalties, with the unboxed terms constituting $-2\log(L^*)$. Here, N denotes the length of series, m the number of changepoints, τ_i is the time of the <i>i</i> th changepoint, and $\hat{\sigma}^2$ is the estimated white noise variance.	35
693 694 695 696	Table 2.	Model fitting results. Here, $\hat{\sigma}^2$ denotes the estimated variance of the white noise (* is assumed rather than estimated). Bolded values are the smallest penalized score. All model residuals have been checked for normality (Shapiro-Wilk's & Kolmogorov-Smirnov test) and constant variance (Levene's test).	36
697 698 699 700 701	Table 3.	Model fitting results based on truncated CET series. Here, $\hat{\sigma}^2$ denotes the estimated variance of the white noise (* is assumed rather than estimated). Bolded values are the smallest penalized score. All model residuals have been checked for normality (Shapiro-Wilk's & Kolmogorov-Smirnov test) and constant variance (Levene's test).	37
702	Table 4.	BIC posterior probabilities for models fitted to the full and truncated CET series	38
703	Table 5.	Parameter estimates of the best fitting model: trend shifts with white noise errors .	39

TABLE 1: Penalized likelihoods. The boxed terms are the penalties, with the unboxed terms constituting $-2\log(L^*)$. Here, N denotes the length of series, m the number of changepoints, τ_i is the time of the *i*th changepoint, and $\hat{\sigma}^2$ is the estimated white noise variance.

Criteria Objective Function

BIC
$$N \log(\hat{\sigma}^2) + N + N \log(2\pi) + (3m+4)\log(N)$$

MDL $N \log(\hat{\sigma}^2) + N + N \log(2\pi) + 2\log(N) + 2\log(m) + 2\sum_{i=1}^{m+1} \log(\tau_i - \tau_{i-1}) + 2\sum_{i=2}^{m+1} \log(\tau_i)$

(a) Penalized likelihoods for the trend shift model with $\mbox{AR}(1)$ errors

Criteria Objective Function

BIC	$N\log(\hat{\sigma}^2) + N + N\log(2\pi) +$	$(3m+3)\log(N)$
MDL	$N\log(\hat{\sigma}^2) + N + N\log(2\pi) +$	$\log(N) + 2\log(m) + 2\sum_{i=1}^{m+1}\log(\tau_i - \tau_{i-1}) + 2\sum_{i=2}^{m+1}\log(\tau_i)$

(b) Penalized likelihoods for the trend shift model with white noise errors

Criteria Objective function

BIC
$$N \log(\hat{\sigma}^2) + N + N \log(2\pi) + (2m+4)\log(N)$$

MDL $N \log(\hat{\sigma}^2) + N + N \log(2\pi) + 3\log(N) + 2\log(m) + \sum_{i=1}^{m+1} \log(\tau_i - \tau_{i-1}) + 2\sum_{i=2}^{m+1} \log(\tau_i)$

(c) Penalized likelihoods for the fixed slope mean shift with AR(1) errors

Criteria Objective function

BIC
$$N \log(\hat{\sigma}^2) + N + N \log(2\pi) + (2m+1) \log(N)$$

(d) Penalized likelihoods for the Joinpin model with white noise errors

Criteria Objective function

BIC
$$N\log(\hat{\sigma}^2) + N + N\log(2\pi) + 4\log(N)$$

(e) Penalized likelihoods for the long memory model with AR(1) errors. Minus log(N) for white noise errors.

TABLE 2: Model fitting results. Here, $\hat{\sigma}^2$ denotes the estimated variance of the white noise (* is assumed rather than estimated). Bolded values are the smallest penalized score. All model residuals have been checked for normality (Shapiro-Wilk's & Kolmogorov-Smirnov test) and constant variance (Levene's test).

Model	Penalty	Flagged Changepoints	$\hat{\sigma}^2$	Log-likelihood	Penalized Score
Trend shifts $+ AP(1)$	BIC	1700,1739,1988	0.290	-288.80	654.19
	MDL	1700,1739,1988	0.290	-288.80	656.52
Trend shifts+WN	BIC	1700,1739, 1988	0.291	-290.02	650.74
ficht shifts+ with	MDL	1700,1739, 1988	0.291	-290.02	653.07
Fixed slope mean shift $\Delta \mathbf{P}(1)$	BIC	1988	0.325	-310.11	655.79
	MDL	1988	0.325	-310.11	658.93
Joinpin	BIC	1973	0.291*	-321.19	654.17
Long-memory+AR(1)	BIC	-	0.579	-316.59	656.75
Long-memory	BIC	-	0.584	-319.31	655.93

TABLE 3: Model fitting results based on truncated CET series. Here, $\hat{\sigma}^2$ denotes the estimated variance of the white noise (* is assumed rather than estimated). Bolded values are the smallest penalized score. All model residuals have been checked for normality (Shapiro-Wilk's & Kolmogorov-Smirnov test) and constant variance (Levene's test).

Model	Penalty	Flagged Changepoints	$\hat{\sigma}^2$	Log-likelihood	Penalized Score
Trend shifts $\perp \Delta \mathbf{R}(1)$	BIC	1987	0.305	-205.44	449.51
	MDL	1987	0.305	-205.44	450.70
Trend shifts+WN	BIC	1987	0.308	-206.13	445.36
field shifts+ wiv	MDL	1987	0.308	-206.13	446.55
Fixed slope mean shift $\Delta \mathbf{R}(1)$	BIC	1990	0.306	-208.06	449.23
	MDL	1990	0.306	-208.06	452.51
Joinpin	BIC	-	0.308*	-220.72	452.47
Long-memory+AR(1)	BIC	-	0.333	-217.01	450.57
Long-memory	BIC	_	0.340	-219.41	449.85

Model	Full	Truncated
Trend shifts + AR(1)	0.11	0.08
Trend shifts + WN	0.64	0.68
Fixed slope +mean shifts+AR(1)	0.05	0.10
Joinpin	0.12	0.02
Long-memory+AR(1)	0.03	0.05
Long-memory	0.05	0.07

 TABLE 4: BIC posterior probabilities for models fitted to the full and truncated CET series

Segment	Slope (°C/yr)					
1659-1699	-0.027					
1700-1738	0.026					
1739-1987	0.002					
1988-2020	0.011					
(a) Full CET						
Segment	Slope (°C/yr)					
1772-1986	0.002					
1987-2020	0.016					

TABLE 5: Parameter estimates of the best fitting model: trend shifts with white noise errors

(b) Truncated CET

704 LIST OF FIGURES

705	Fig. 1.	Station locations and annual average temperatures of Central England.	41
706 707 708	Fig. 2.	Estimated CET trend shift structure. BIC and MDL flag the same changepoints in both the CET series (1700, 1739, 1988, red solid line) and truncated CET (1987, blue dashed line) series when assuming either AR(1) or white noise errors.	42
709 710 711	Fig. 3.	The estimated CET trend shift structure for the full (red solid line) and truncated CET (blue dashed line) series when a constant regime trend slope is imposed. Both BIC and MDL flag a single changepoint in 1988 for the full series and 1990 for the truncated series.	43
712 713	Fig. 4.	Estimated CET joinpin shift structure for full (red solid line) and truncated (blue dashed line) series. BIC flags one shift in 1973 in the full series and and none for the truncated series.	44
714 715 716	Fig. 5.	The estimated CET mean shift structure for full (red solid line) and truncated (blue dashed line) series. BIC and MDL detect the same changepoints for both the CET and truncated CET series assuming white noise errors.	45
717 718 719	Fig. 6.	Histogram of differences in BIC scores between the trend and mean-shift models. The correct model is the trend-shift model; however, BIC selects the mean-shift model the majority of the time.	46



FIG. 1: Station locations and annual average temperatures of Central England.

(a) Locations of weather stations.



(b) Annual average temperature of Central England.

FIG. 2: Estimated CET trend shift structure. BIC and MDL flag the same changepoints in both the CET series (1700, 1739, 1988, red solid line) and truncated CET (1987, blue dashed line) series when assuming either AR(1) or white noise errors.



FIG. 3: The estimated CET trend shift structure for the full (red solid line) and truncated CET (blue dashed line) series when a constant regime trend slope is imposed. Both BIC and MDL flag a single changepoint in 1988 for the full series and 1990 for the truncated series.





FIG. 4: Estimated CET joinpin shift structure for full (red solid line) and truncated (blue dashed line) series. BIC flags one shift in 1973 in the full series and and none for the truncated series.

FIG. 5: The estimated CET mean shift structure for full (red solid line) and truncated (blue dashed line) series. BIC and MDL detect the same changepoints for both the CET and truncated CET series assuming white noise errors.



FIG. 6: Histogram of differences in BIC scores between the trend and mean-shift models. The correct model is the trend-shift model; however, BIC selects the mean-shift model the majority of the time.



Difference in BIC Scores (Trend model – Mean model)