

1 **Detrital Thermochronometry - Recorder of Earth's Dynamic Past**

2 **Daniel F. Stockli¹ and Yani M. R. Najman^{2,3}**

3 ¹Dept. of Geological Sciences, University of Texas at Austin, USA

4 ²Lancaster University, UK, and ³University of Colorado, Boulder, USA

5 **Abstract**

6 Advances in detrital noble gas thermochronometry by $^{40}\text{Ar}/^{39}\text{Ar}$ and (U-Th)/He dating are
7 improving the resolution of sedimentary provenance reconstructions and providing new insights
8 into the evolution of the Earth's surface through time. While detrital petrography and
9 geochronology can illuminate sedimentary provenance and reconstructions of ancient drainage
10 networks and sedimentary basins, they can be hampered by sediment recycling or monotonous
11 source signatures and often do not provide quantitative insights into the tectonic or erosional
12 evolution of drainage basins. Detrital thermochronometry has the added ability to quantify
13 tectonic unroofing or erosion, temporal and dynamic connections between source and sink,
14 sediment lag-times and transfer rates, depositional timing, and post-depositional burial heating.
15 Hence, detrital thermochronometry provides the unique ability to use the detrital record in basins
16 to reconstruct the Earth's dynamic long-term landscape evolution and coupling of basins and
17 their hinterland.

18 **Keywords:** thermochronology, detrital, provenance, source-to-sink, lag times, basin evolution

19

20 **Introduction**

21 The Earth's surface has been modified through the interplay between lithospheric,
22 hydrologic, and atmospheric processes throughout geological times. These dynamic interactions
23 have been archived in sedimentary records preserved in the sedimentary basins of the world.

24 These archives of material eroded from continents and mountain belts provide an invaluable,
25 long-term record of tectonism and erosion in the hinterland – a record that is often no longer
26 accessible in the bedrock of the source region due to progressive tectonic or metamorphic
27 overprinting or subsequent erosion.

28 While petrographic studies have long been used to unravel sedimentary provenance and
29 thus hinterland tectonics and erosion, single-grain detrital isotopic techniques are increasingly
30 used as a potent tool for reconstructing hinterland tectonics, past drainage systems, landscapes,
31 linkages between sediment sources and depositional sinks, or hydrocarbon reservoir
32 characterizations. This revolution is largely attributable to technological and analytical advances,
33 in particular zircon U-Pb laser-ablation inductively-coupled-plasma mass-spectrometry (LA-
34 ICP-MS), making data more affordable and more readily available in large numbers (e.g.,
35 Gehrels, 2014). However, interpretations based on zircon U-Pb crystallization ages can be
36 hampered by non-diagnostic source signatures and sediment recycling, preventing discrimination
37 between different source terranes . Thus, although these traditional provenance tools can shed
38 light on the evolution of drainage basins, low-temperature thermochronometers can deliver
39 quantitative insights into the dynamic and thermal evolution related to tectonic unroofing or
40 erosional denudation of different portions of the source area (e.g., Ehlers and Farley, 2003;
41 Reiners and Ehlers, 2005). Hence, detrital thermochronometry affords the unique ability to
42 constrain the history of source regions in terms of both the timing and rates of exhumation or
43 erosion.

44 Detrital thermochronometry has been applied on very different spatial scales, ranging
45 from individual drainages to elucidate geomorphic or erosional processes, to regional drainage
46 systems to reconstruct source-to-sink linkages and tectonic processes in the hinterland, to

47 continental scale reconstruction of drainage system evolution or mantle-driven dynamic
48 topography. Detrital thermochronometry, based on either noble gas or fission track
49 measurements, produced by radiogenic alpha particles, has mainly been used to decode cooling
50 histories recorded by K-bearing rock-forming and U-Th bearing accessory minerals that are
51 common detrital phases in modern and ancient siliciclastic sedimentary deposits (e.g., Bernet and
52 Spiegel, 2004; Hodges et al., 2005). Although fission track techniques have low precision single
53 grain ages compared to $^{40}\text{Ar}/^{39}\text{Ar}$ and (U-Th)/He dating, interpretations are never based on an
54 age of a single grain, but always populations, age peaks, or weighted averages of a number of
55 grains to shed light on the exhumational history of different portions of the source area or reveal
56 abrupt changes in the source area. These are characterized by nominal closure temperatures for
57 these minerals of $\sim 350^\circ\text{C}$ for muscovite and $\sim 350\text{-}180^\circ\text{C}$ for K-feldspar $^{40}\text{Ar}/^{39}\text{Ar}$, $\sim 250^\circ\text{C}$ for
58 zircon and $\sim 100^\circ\text{C}$ for apatite fission track, and $\sim 180^\circ\text{C}$ for zircon and $\sim 60^\circ\text{C}$ for apatite (U-
59 Th)/He (e.g., Reiners, 2009). See Gautheron and Zeitler (this issue) for further discussion of this
60 concept. These different moderate- to low-temperature sensitivity windows can make them either
61 excellent detrital recorders of tectonic and erosional signals in the source area or sensitive to
62 post-depositional burial heating and resetting. Overall, improved and novel noble gas
63 thermochronometric methods and their applications, integrated with innovative new conceptual
64 and numerical models, provide an improved quantitative understanding of timing, rates, and
65 spatial patterns of long-term landscape evolution, erosion budgets, and tectonic and geodynamic
66 processes.

67 **Detrital Thermochronometry**

68 The sedimentary records of basins have long been linked to the erosion of mountain belts
69 to document hinterland erosion and tectonic activity and quiescence. At the same time, bedrock

70 thermochronometry has become a routine approach to recover thermal histories of mountain
71 belts and thus to quantify their short- and long-term tectonic, exhumational and erosional
72 unroofing histories (Fig. 1). Detrital thermochronometric methods combine these two approaches
73 and have been applied to sedimentary archives in order to reconstruct erosional and tectonic
74 processes in large-scale orogenic systems (e.g., Bernet and Spiegel, 2004; Malusà and Fitzgerald,
75 2019). Early studies focused mainly on giant orogenic systems, such as the Himalayas and
76 associated foreland basin and large-scale depocenters of the Indus and Bengal Rivers, in light of
77 high to ultra-high exhumation rates (>5 km/Myrs) in the orogenic hinterland and the substantial
78 sedimentary archives (Najman, 2006). These early studies relied mainly on detrital $^{40}\text{Ar}/^{39}\text{Ar}$
79 thermochronometry, as this could be accomplished by high-precision, single-grain $^{40}\text{Ar}/^{39}\text{Ar}$
80 dating of detrital muscovite and K-feldspar, and the fact that exhumation rates and magnitudes
81 were sufficient to reset $^{40}\text{Ar}/^{39}\text{Ar}$ ages in the source terranes. Practically, this methodology also
82 benefited from the possibility of $^{40}\text{Ar}/^{39}\text{Ar}$ dating large numbers of individual detrital grains by
83 laser, yielding statistically robust datasets. However, the cost and speed of analysis (notably the
84 need for neutron irradiations) represents an impediment for detrital $^{40}\text{Ar}/^{39}\text{Ar}$ dating.
85 Furthermore, the relatively high closure temperatures of $^{40}\text{Ar}/^{39}\text{Ar}$ minerals ($>300^\circ\text{C}$) limits their
86 thermal sensitivity to upper-crustal processes in the source area.

87 Detrital thermochronometry of refractory accessory phases, such as zircon, titanite, or
88 apatite, by (U-Th)/He dating has experienced dramatic growth over the last decade. Many of
89 these detrital low-temperature thermochronometers leverage their $<200^\circ\text{C}$ closure temperatures
90 to reveal upper-crustal (<6 km) tectonic activity and erosion in sediment source areas. Besides
91 refining provenance, detrital thermochronometry is able to recover source terrane exhumation
92 rates by using the concept of lag-time – a measure of the temporal difference between cooling

93 age of the detrital mineral grain in the source region, and the depositional age of the sedimentary
94 host rock (Figs. 1 and 2). Whereas application of apatite fission-track dating in detrital studies
95 has been hampered by both post-depositional burial resetting and low single-grain precision, thus
96 relying on multi-grain pooled ages, detrital fission track dating of zircon has been more widely
97 applied in mega-orogenic source to sink studies (e.g. Carter et al., 2019). Most detrital zircon
98 fission-track provenance studies exploit its low closure temperature ($\sim 250^{\circ}\text{C}$) to provide insights
99 into source area denudation and temporal variations in exhumation rates in the Himalayas and
100 North American Cordillera (Cerveny et al., 1988, Garver and Brandon, 1994). Increasingly, these
101 studies couple zircon fission track dating with U-Pb dating of the same crystal to refine
102 provenance identification, and identification of volcanic grains (U-Pb age = low-temperature
103 thermochronometric age) (Fig 2). This double-dating approach can provide more detailed
104 insight into the cooling history of specific hinterland terranes (e.g., Carter and Moss, 1999).
105 However, the applicability of zircon fission track dating has been limited by analytical
106 complexities (e.g. etching), single-grain precision, and uncertainties in thermal sensitivity due to
107 lack of a robust track annealing model (Malusà and Fitzgerald, 2019).

108 The higher precision of single-grain (U-Th)/He dating of accessory phases overcomes
109 some of the limitations of detrital fission track dating. In particular, zircon (U-Th)/He dating,
110 with a single-grain age precision ($\sim 8\%$) and a nominal closure temperature of $\sim 180^{\circ}\text{C}$, has
111 become widely utilized for revealing upper-crustal exhumation histories and tectonic activity in
112 source terranes (e.g., Rahl et al., 2003; Reiners, 2005). Importantly, zircon is also very conducive
113 to U-Pb-He double dating of the same detrital grain, enabling more differentiated and
114 sophisticated provenance interpretations by linking source area crystallization ages with source
115 area exhumation histories (e.g., Reiners et al., 2005; Thomson et al., 2017; Carter, 2019). Detrital

116 apatite (U-Th)/He dating can also be a powerful tool in modern geomorphic and active tectonic
117 studies, given its low to very low thermal sensitivity window (e.g., Stock et al., 2006), although
118 limitations arise due to the need for intact, inclusion-free apatite, and its propensity for
119 undergoing mechanical abrasion during transport and post-depositional burial (partial) resetting.

120

121 **Earth's Dynamic Past in a Grain of Sand**

122 Detrital thermochronometry has become more versatile and sophisticated through an expanded
123 analytical repertoire, innovative conceptual approaches, and numerical modeling, making it more
124 applicable to a wider spectrum of tectonic and geomorphic environments. These data can
125 constrain the timing and rates of hinterland erosion and tectonic unroofing or maximum
126 depositional ages and presence of first-cycle volcanic zircons, as illustrated in the following
127 sections

128 **Improved Sedimentary Provenance**

129 Zircon U-Pb geochronology has exploded as a sedimentary provenance tool over the past
130 decade with the advent of LA-ICP-MS analysis, making data available in staggering numbers.
131 Detrital zircon U-Pb provenance studies, based on source crystallization ages, have worked
132 exceedingly well in geological settings with highly-variable crustal formation ages, such as
133 North and South America, and the Himalayas. However, as shown by modern river studies, this
134 is often not the case, due to either monotonous source signatures or significant recycling of older
135 sedimentary strata.

136 U-Pb-He double dating can differentiate between different sources with the same or
137 similar zircon U-Pb age signatures (Fig. 2). For example, in the Paleogene foreland basin of the
138 southern Pyrenees, Thomson and others (2019) showed that identical Variscan U-Pb ages from

139 the foreland and the orogenic wedge can readily be differentiated by (U-Th)/He double dating.
140 While Variscan zircons from the foreland are characterized by Permian zircon He ages, Variscan
141 zircons from the Pyrenees yielded early Paleogene He cooling ages reflecting rapid tectonic
142 exhumation of the fold and thrust belt (Fig. 3). Similarly, Xu and others (2018) showed that U-
143 Pb-He double dating of Miocene strata in the northern Gulf of Mexico is able to differentiate
144 detrital zircons with non-diagnostic, invariant late Mesoproterozoic (Grenvillian) ages by
145 leveraging their different Appalachian, Laramide, and Cenozoic zircon He cooling ages and thus
146 improve sediment provenance reconstructions on the basis of otherwise non-diagnostic zircons.

147 **Hinterland's Dynamic Past**

148 In addition to refining sedimentary provenance, detrital thermochronometry potentially
149 makes lag time calculations possible, with the aim of quantifying hinterland bedrock exhumation
150 rates from the detrital signature. Lag time, as defined above, is a measure of the time elapsed
151 between a mineral grain's passage through its closure isotherm to its surface exposure, erosion
152 and deposition (e.g., Ruiz et al., 2004; Saylor et al., 2012). Hence, assuming negligible transport
153 and recycling duration, minimum lag times reflect the rapidity of source terrane exhumation rates
154 and give a measure of the dynamic evolution of the hinterland (Fig. 1). While short lag times
155 reflect rapid unroofing, long lag-times are indicative of slow exhumation or long-term
156 intermediate storage. The temporal evolution and variation in hinterland exhumation rates can be
157 assessed by comparing lag time estimates from different stratigraphic horizons. A temporal
158 decrease in lag time can be interpreted as an increase in source terrane exhumation rate, while an
159 increase in lag time might signal either a decrease in exhumation rate or a change in sediment
160 source (e.g. Saylor et la., 2012).

161 Detailed zircon (U-Th)/He and apatite fission track analyses from the northern Alpine
162 Molasse foreland basin in central Switzerland illustrate the insights that can be gleaned from
163 high-resolution lag-time reconstructions (Fig. 4). The Swiss Molasse basin is characterized by
164 two major shallowing-upward Oligo-Miocene sedimentary sequences, divided by a marine
165 transgression in the early Miocene (Schlunegger et al., 1997). While it has been suggested that
166 thrusting and erosion might have waned during the transgression, thermochronometric lag times
167 tell a different story. Lag times are uniformly ~5-10 Myrs during deposition of the Lower
168 Freshwater Molasse, but decrease abruptly to ~0 Myrs at the transgressive base of the Upper
169 Freshwater Molasse at ~19 Ma, signaling a major phase of accelerated exhumation of the fold-
170 and-thrust belt, possibly due to either out-of-sequence thrusting or syn-orogenic hinterland
171 extension. Lag times again increase to >10 Myrs at ~12 Ma, interpreted as a decrease in
172 shortening in the Helvetic thrust belt, the onset of shortening in the Jura Mountains, and the
173 Molasse basin becoming a piggy-back basin. Hence, high-resolution lag time studies have great
174 potential to elucidate tectonic and erosional evolution of an orogenic hinterland (Fig. 4).

175 Recently Malusà and Fitzgerald (2020) explored some of the assumptions and limitations
176 of quantitative lag-time interpretations. They pointed out that for reliable lag-time quantification,
177 ages need to reflect mineral closure ages resulting from a simple cooling history (no partial
178 resetting), source terrane isotherms should be in steady-state through time, and sediment
179 transport time from source to sink should be negligible. Importantly, for reliable lag-time
180 estimates, it is critical to exclude first-cycle volcanic grains as they do not track the exhumation
181 path from closure to erosion. This is best accomplished by U-Pb-He double-dating and the
182 exclusion of volcanic grains with identical crystallization and cooling ages (Saylor et al., 2012).

183 **Short- and Long-Term Landscape Evolution**

184 Detrital geo- and thermochronometry are potent tools for reconstructing past drainages
185 and landscapes or continental paleogeographies, and elucidating their long-term erosional and
186 landscape evolution. Detrital zircon U-Pb dating has been the principle workhorse for these
187 reconstructions, but detrital thermochronometry has both refined these reconstructions and, more
188 importantly, provided unique quantitative insights into the dynamic evolution of continents.

189 Detrital thermochronometry is not limited in its application to large-scale, deep-time
190 reconstructions; studies have also used detrital thermochronometric data to track geomorphic
191 processes, such as the locus of erosion within a catchment or the nature and efficiency of erosive
192 agents (e.g., glaciers). This approach is most effective if the distribution of bedrock
193 thermochronometric ages within a drainage are spatially variable and well-defined, for
194 pinpointing where sediment came from, and attributing detrital ages to a specific portion or
195 elevation within a drainage.

196 Stock and others (2006) utilized apatite (U-Th)/He to investigate glacial, fluvial, and
197 hillslope erosion processes in the Sierra Nevada. By using statistical comparison of observed and
198 predicted age distributions, based on catchment hypsometry, they were able to point to the
199 locations of sediment generation and storage in mountain-scale catchments. Using a similar
200 comparison of catchment erosion in the Nepal Himalayas using detrital muscovite $^{40}\text{Ar}/^{39}\text{Ar}$
201 ages, Ruhl and Hodges (2005) reconstructed catchment-averaged erosion rates to evaluate spatial
202 uniformity or heterogeneity of erosion, and to explore possible transience in erosional processes
203 and their departure from steady-state over the million-year timescale. Finally, detrital
204 thermochronometry from modern catchments has proven useful in mapping late Cenozoic
205 exhumation patterns in large, inaccessible regions such as the ice-covered mountains of Alaska
206 (Lease et al., 2016).

207

208 **Stratigraphic Age Constraints**

209 U-Pb dating of first-cycle detrital volcanic zircons has been extensively used to determine
210 depositional ages in strata that lack biostratigraphic age constraints (Dickinson and Gehrels,
211 2009). Although this methodology has been shown to work well in tectonic basins adjacent to
212 magmatic arcs, such as forearc and back- or retro-arc basins, it is commonly not applicable in
213 basins associated with continent-continent collisions, rifts, or continental margins due to the lack
214 of new magmatic zircon generation. In these cases, the youngest mode of detrital
215 thermochronological cooling ages, assuming no post-depositional resetting, can be used to
216 estimate maximum depositional ages. This approach relies on the fact that, in rapidly exhuming
217 regions, the youngest mode of cooling ages has very short or negligible lag times and therefore
218 approximates the stratigraphic age.

219 In fold and thrust belts this thermochronometric approach to determining depositional
220 ages has been applied in a number of studies to date syn-tectonic continental siliciclastic
221 deposits. Najman and other (1997) employed laser muscovite $^{40}\text{Ar}/^{39}\text{Ar}$ dating in NW India to
222 corroborate the syn-tectonic nature of early Himalayan foreland basin sediments, and to
223 constrain their maximum depositional ages. Short-lag time detrital ages reflected the very rapid,
224 Oligo-Miocene exhumation of mid-crustal, muscovite-bearing rocks from the crystalline core of
225 the Himalayas. While maximum depositional age constraints, based on muscovite $^{40}\text{Ar}/^{39}\text{Ar}$
226 dating, require ultra-rapid exhumation in the tectonically-active hinterland, (U-Th)/He data can
227 yield reliable depositional ages, based on short lag-time ages, even with limited magnitudes of
228 upper-crustal exhumation. Thomson and others (2017) demonstrated that, despite the absence of
229 volcanic input, the youngest zircon (U-Th)/He age modes from foreland basin strata in the

230 southern Pyrenees yielded maximum depositional age constraints for most of the Paleocene to
231 Eocene section, similar to biostratigraphic and magnetostratigraphic age estimates. In fact, even
232 samples with 5 Myrs lag time yield maximum depositional ages that would only be an
233 overestimation of 10% for sediments that otherwise lack any depositional age information.

234 This approach has proven particularly powerful to resolve the stratigraphy of the Nubian
235 sandstone of NE Africa and Arabia – a massive, undifferentiated mega-sequence of Paleozoic
236 and Mesozoic quartz-arenite sandstones with a monotonous and non-diagnostic detrital zircon
237 signatures, dominated by Pan-African and older zircons. In contrast, the youngest mode of
238 detrital zircon (U-Th)/He ages allows for an amazingly clear identification of different
239 Carboniferous, Triassic, and Cretaceous stratigraphic packages (Fig. 5). Age-diagnostic detrital
240 zircon (U-Th)/He signatures record the tectonic evolution of source terranes in response to
241 different tectonic events, such as Hercynian block faulting, Triassic and Jurassic Tethyan
242 extension and rifting, and Late Cretaceous inversion. These age constraints allow for both
243 chronostratigraphic unraveling of a km-scale monotonous sedimentary package as well as
244 genetic attribution of different units to tectonic events affecting the region (Fig. 5). Szymanski
245 and others (2016) showed how this approach can also be applied to relatively small continental
246 extensional basins. Their work on a Saudi Red Sea rift basin differentiated between pre- and syn-
247 rift strata and provided direct age constraints on undated continental syn-rift strata in a small (<3
248 km) half-graben by documenting the presence of reset early Miocene apatite (U-Th)/He ages
249 derived from the exhumed extensional footwall.

250 **Conclusions and Prospectives**

251 Detrital thermochronometry has evolved into a powerful and versatile approach that is no
252 longer only limited to large-scale mega-tectonics. In particular, detrital noble gas

253 thermochronometry has reinvented itself through new analytical tools, more readily available
254 high-resolution and large datasets of precise single-grain ages, and analytical, conceptual, and
255 numerical integration, allowing for a higher thermal, temporal, and spatial resolution of tracking
256 and quantification of geological processes.

257 Detrital thermochronometry is a potent means to provide lag time estimates that
258 illuminate hinterland erosion and tectonic exhumation magnitudes and rates, and thus quantify
259 the dynamic evolution of sedimentary source terranes with considerable temporal and spatial
260 resolution. Furthermore, geo- and thermochronometric double dating (e.g., U-Pb-He double
261 dating) can provide more accurate sedimentary provenance information. This approach allows
262 for a better characterization of sediment contributions from volcanic, first-cycle basement, or
263 recycled sedimentary input into source-to-sink systems on the basis of both lag time and age
264 difference between crystallization and cooling ages. Lastly, a rapid growth area of detrital noble
265 gas thermochronometry lies in its application of delineating maximum depositional ages in areas
266 that lack input of volcanic zircons. The use of ultra-short lag time cooling ages in foreland
267 basins, continental rifts, or continental margins is proving to be a reliable way of providing
268 stratigraphic age constraints, a means to differentiate monotonous continental sedimentary
269 packages, and to illuminate their associated tectonic drivers.

270 Detrital thermochronometry has been applied in a range of tectonic environments, from
271 orogenic systems to continental rifts. High-octane orogenic systems and fold-and-thrust belts,
272 such as the Himalayas or Andes, have long been the playground for detrital provenance and
273 thermochronometry studies, as the thermochronometric signals, like the mountains themselves,
274 are big. Analytical advances have dramatically improved our ability to resolve the timing and
275 rates of hinterland exhumation, making it possible to tie detrital fluxes and depositional records

276 to specific thrust sheets, and to provide a detailed record of the dynamic past of an orogen. In
277 extensional rifts and continental margins, bedrock thermochronometry has been a compelling
278 technique to elucidate the temporal and exhumation histories, while isotopic provenance studies
279 have focused on large-scale source provenance reconstructions, basin fill history, and reservoir
280 characterizations. Detrital thermochronometry is starting to play a more important role in
281 reconstructing rifts and continental margins, and elucidating the dynamic interactions of crustal
282 and lithospheric tectonics. Application of lag-times, maximum depositional age estimates, and
283 better quantification of post-depositional thermal effects, have allowed detrital
284 thermochronometry to become a more sophisticated tool in reconstructing the dynamic evolution
285 of continental rifting and break-up.

286 Looking to the future, as detrital noble gas thermochronometry and its combination with
287 U-Pb double dating is experiencing rapid growth, more time- and cost-effective data acquisition
288 is needed - the development of detrital laser (U-Th)/He dating should provide this boost (e.g.,
289 Horne et al., 2016). In addition, improved recovery of thermal information from single detrital
290 grains would also be desirable. While this has been potentially possible, though rarely utilized, in
291 the case of single-grain detrital $^{40}\text{Ar}/^{39}\text{Ar}$ step-heating, it should also be feasible in the case of
292 $^4\text{He}/^3\text{He}$ thermochronometry (e.g., Shuster and Farley, 2005). To date, no detrital $^4\text{He}/^3\text{He}$
293 thermochronometric studies have been attempted. However, statistically more robust large-
294 sample size studies and single-grain thermal history recovery in the future will also require better
295 integration of numerical modeling, such as coupled thermo-mechanical, landscape, and
296 geodynamic modeling, to warrant increased analytical effort.

297 **Acknowledgements**

298 We would like to thank Devon Orme, Richard Lease, and Jon Blundy for helpful and
299 constructive reviews and Marissa Tremblay, Emily Cooperdock, and Peter Zeitler for editorial
300 handing. Stockli would like to acknowledge support by the Chevron (Gulf) Centennial
301 Professorship. This paper was written while Najman was a visiting scholar at University of
302 Colorado, Boulder, supported by the CIRES Visiting Fellow Program funded by NOAA
303 agreement no. NA17OAR4320101.

304 **Reference list (30):**

305 Bernet, M., and Spiegel, C., eds., 2004, Detrital thermochronology - Provenance analysis,
306 exhumation, and landscape evolution of mountain belts: Boulder, Colorado, Geological Society
307 of America Special Paper 378, p. 37–50.

308
309 Carter, A., 2019, Thermochronology on sand and sandstones for stratigraphic and provenance
310 studies. In Fission-Track Thermochronology and its Application to Geology (pp. 259-268).
311 Springer, Cham.

312
313 Carter, A. and Moss, S.J., 1999, Combined detrital-zircon fission-track and U-Pb dating: A new
314 approach to understanding hinterland evolution. *Geology*, 27(3), pp.235-238.

315
316 Cervený, P.F., Naeser, N.D., Zeitler, P.K., Naeser, C.W. and Johnson, N.M., 1988. History of
317 uplift and relief of the Himalaya during the past 18 million years: Evidence from fission-track
318 ages of detrital zircons from sandstones of the Siwalik Group. In *New perspectives in basin*
319 *analysis* (pp. 43-61). Springer, New York, NY.

320
321 Ehlers, T.A. and Farley, K.A., 2003, Apatite (U–Th)/He thermochronometry: methods and
322 applications to problems in tectonic and surface processes. *Earth and Planetary Science*
323 *Letters*, 206(1-2), pp.1-14.

324
325 Garver, J.I. and Brandon, M.T., 1994, Erosional denudation of the British Columbia Coast
326 Ranges as determined from fission-track ages of detrital zircon from the Tofino basin, Olympic
327 Peninsula, Washington. *Geological Society of America Bulletin*, 106(11), pp.1398-1412.

328
329 Gautheron and Zeitler, this issue

330
331 Gehrels, G., 2014, Detrital Zircon U-Pb Geochronology Applied to Tectonics. *Annual Review of*
332 *Earth and Planetary Sciences* 2014 42:1, 127-149

333
334 Hodges, K.V., Ruhl, K.W., Wobus, C.W. and Pringle, M.S., 2005, $^{40}\text{Ar}/^{39}\text{Ar}$ thermochronology
335 of detrital minerals. *Reviews in Mineralogy and Geochemistry*, 58(1), pp.239-257.

336

337 Horne, A.M., van Soest, M.C., Hodges, K.V., Tripathy-Lang, A. and Hourigan, J.K., 2016,
338 Integrated single crystal laser ablation U/Pb and (U–Th)/He dating of detrital accessory
339 minerals–Proof-of-concept studies of titanites and zircons from the Fish Canyon
340 tuff. *Geochimica et Cosmochimica Acta*, 178, pp.106-123.
341
342 Lease, R.O., Haeussler, P.J. and O'Sullivan, P., 2016, Changing exhumation patterns during
343 Cenozoic growth and glaciation of the Alaska Range: Insights from detrital thermochronology
344 and geochronology. *Tectonics*, 35(4), pp.934-955.
345
346 Malusà, M.G. and Fitzgerald, P.G., 2019, Application of thermochronology to geologic
347 problems: bedrock and detrital approaches. In *Fission-Track Thermochronology and its*
348 *Application to Geology* (pp. 191-209). Springer, Cham.
349
350 Miller, J.C., 2012, Detrital thermochronology of the Alpine foreland basin in Central
351 Switzerland: Insights into tectonic and erosion history of the North Central Alps. University of
352 Kansas, Unpublished M.S. thesis p. 169.
353
354 Najman, Y.M.R., Pringle, M.S., Johnson, M.R.W., Robertson, A.H.F. and Wijbrans, J.R., 1997,
355 Laser $^{40}\text{Ar}/^{39}\text{Ar}$ dating of single detrital muscovite grains from early foreland-basin sedimentary
356 deposits in India: Implications for early Himalayan evolution. *Geology*, 25(6), pp.535-538.
357
358 Najman, Y., 2006, The detrital record of orogenesis: A review of approaches and techniques used
359 in the Himalayan sedimentary basins. *Earth-Science Reviews*, 74, 1-72.
360
361 Pujols, E.J., 2011, Temporal and thermal evolution of extensional faulting in the central Gulf of
362 Suez and detrital zircon (U-Th)/He constraints on the thermo-tectonic Paleozoic and Mesozoic
363 history of the Sinai, Egypt. University of Kansas, Unpublished M.S. thesis p. 247.
364
365 Rahl, J.M., Reiners, P.W., Campbell, I.H., Nicolescu, S. and Allen, C.M., 2003, Combined
366 single-grain (U-Th)/He and U/Pb dating of detrital zircons from the Navajo Sandstone,
367 Utah. *Geology*, 31(9), pp.761-764.
368
369 Reiners, P.W., 2005. Zircon (U-Th)/He thermochronometry. *Reviews in Mineralogy and*
370 *Geochemistry*, 58(1), pp.151-179.
371
372 Reiners, P.W., Campbell, I.H., Nicolescu, S., Allen, C.M., Hourigan, J.K., Garver, J.I.,
373 Mattinson, J.M. and Cowan, D.S., 2005, (U-Th)/(He-Pb) double dating of detrital
374 zircons. *American Journal of Science*, 305(4), pp.259-311.
375
376 Reiners, P.W. and Ehlers, T.A. eds., 2018, *Low-Temperature Thermochronology:: Techniques,*
377 *Interpretations, and Applications* (Vol. 58). Walter de Gruyter GmbH & Co KG.
378
379 Reiners, P.W., 2009, Nonmonotonic thermal histories and contrasting kinetics of multiple
380 thermochronometers. *Geochimica et Cosmochimica Acta* 73(12), 3612–3629.
381

- 382 Ruhl, K.W. and Hodges, K.V., 2005, The use of detrital mineral cooling ages to evaluate steady
383 state assumptions in active orogens: an example from the central Nepalese Himalaya. *Tectonics*,
384 24, TC4015. dx. doi. org/10.1029.
- 385
- 386 Ruiz, G.M.H., Seward, D. and Winkler, W., 2004, Detrital thermochronology—a new perspective
387 on hinterland tectonics, an example from the Andean Amazon Basin, Ecuador. *Basin*
388 *Research*, 16(3), pp.413-430.
- 389
- 390 Saylor, J.E., Stockli, D.F., Horton, B.K., Nie, J. and Mora, A., 2012, Discriminating rapid
391 exhumation from syndepositional volcanism using detrital zircon double dating: Implications for
392 the tectonic history of the Eastern Cordillera, Colombia. *Bulletin*, 124(5-6), pp.762-779.
- 393
- 394 Shuster, D.L. and Farley, K.A., 2005, $4\text{He}/3\text{He}$ thermochronometry: theory, practice, and
395 potential complications. *Reviews in Mineralogy and Geochemistry*, 58(1), pp.181-203.
- 396
- 397 Schlunegger, F., Jordan, T.E. and Klaper, E.M., 1997, Controls of erosional denudation in the
398 orogen on foreland basin evolution: the Oligocene central Swiss Molasse Basin as an
399 example. *Tectonics*, 16(5), pp.823-840.
- 400
- 401 Stock, G.M., Ehlers, T.A. and Farley, K.A., 2006, Where does sediment come from? Quantifying
402 catchment erosion with detrital apatite (U-Th)/He thermochronometry. *Geology*, 34(9), pp.725-
403 728.
- 404
- 405 Szymanski, E., Stockli, D.F., Johnson, P.R. and Hager, C., 2016, Thermochronometric evidence
406 for diffuse extension and two-phase rifting within the Central Arabian Margin of the Red Sea
407 Rift. *Tectonics*, 35(12), pp.2863-2895.
- 408
- 409 Thomson, K.D., Stockli, D.F., Clark, J.D., Puigdefàbregas, C. and Fildani, A., 2017, Detrital
410 zircon (U-Th)/(He-Pb) double-dating constraints on provenance and foreland basin evolution of
411 the Ainsa Basin, south-central Pyrenees, Spain. *Tectonics*, 36(7), pp.1352-1375
- 412
- 413 Thomson, K.D., Stockli, D.F., Odlum, M.L., Tolentino, P., Puigdefàbregas, C., Clark, J. and
414 Fildani, A., 2019, Sediment provenance and routing evolution in the Late Cretaceous–Eocene
415 Ager Basin, south-central Pyrenees, Spain. *Basin Research*.
- 416
- 417 Xu, J., Stockli, D.F. and Snedden, J.W., 2017, Enhanced provenance interpretation using
418 combined U–Pb and (U–Th)/He double dating of detrital zircon grains from lower Miocene
419 strata, proximal Gulf of Mexico Basin, North America. *Earth and Planetary Science Letters*, 475,
420 pp.44-57.

421 425 **Figure Captions**

426
427 **Figure 1.** Diagram illustrating the partial path from cooling, exhumation, erosion, transport,
428 deposition, and burial in an extensional source-to-sink system. Detrital thermochronometric ages
429 and their lag time (tL) can be used as a measure of the rate of exhumation in a source area by
430 estimating the time difference between cooling through a closure isotherm (T_c) at time t_c and

431 erosion at the surface (T_s) at time t_e and depositional age (t_d) at temperature (T_d), assuming a
432 negligible transport time (e.g., Ruiz et al., 2004). For obtaining accurate lag time estimates it is
433 essential to exclude volcanic zircons (t_v), identified as having identical crystallization and
434 cooling ages (t_v , U-Pb=He) (e.g., Saylor et al., 2012).

435
436 **Figure 2.** 3D tectonic cartoon landscape illustrating the power of U-Pb-He double-dating in
437 detrital provenance studies by leveraging spatially variable cooling histories of different tectonic
438 domains underlain by monotonous, non-diagnostic crystalline basement. While U-Pb ages (red
439 line) fail to differentiate different sources, zircon He ages (shaded) allow for distinguishing
440 between different basement terranes and proper attribution of detrital provenance. U-Pb-He
441 double dating diagram graphically deconvolves sink detrital signature from basal sink (E) and
442 helps identify different provenance components.

443
444 **Figure 3.** Case study from the southern Pyrenean foreland basin demonstrating power of U-Pb-
445 He double-dating in differentiating sediment input from foreland (red arrow) and orogenic
446 hinterland (red arrow) in time and space despite non-diagnostic zircon U-Pb signature (modified
447 after Thomson et al., 2019). While Cretaceous strata are dominated by Permo-Triassic cooling
448 ages, the onset of Pyrenean provenance is marked by the abrupt shift to Paleocene-Eocene
449 cooling ages derived from the exhuming Pyrenean fold-and-thrust belt.

450
451 **Figure 4.** Application of temporal variations in thermochronometric lag times in the Cenozoic
452 Northern Alpine Foreland basin in central Switzerland. These data show how the detrital record
453 can record the dynamic tectonic evolution of the orogenic hinterland – a rapid decrease in lag
454 time (t_L), temporally coincident with the Burdigalian transgression in the foredeep, signals the
455 onset of rapid Early Miocene thrust belt exhumation, while a Middle Miocene increase signals
456 the onset of shortening in the Jura Mountains and the Molasse basin becoming a piggy-back
457 basin (data from Miller, 2012). The diagonal lines are contours of constant lag times ($t_L=0, 10,$
458 $20,$ and 30 Myrs).

459
460 **Figure 5.** Application of Maximum Depositional age (MDA) estimates derived from rapidly-
461 exhumed zircon He cooling ages with ultra-short lag times. The Nubian Sandstone covering
462 much large parts of NE Africa and Arabia and spanning most of the Phanerozoic has largely
463 lacked reliable bio- and chronostratigraphic constraints. Short-lag time detrital zircon He cooling
464 ages (ZHe) provide crucial new chronostratigraphic constraints, allowing for stratigraphic
465 differentiation and correlation, and link different stratigraphic packages genetically to tectonic
466 events affecting the region (data from Pujols, 2011).