

1 Substorm onset latitude and the steadiness of 2 magnetospheric convection

S. E. Milan^{1,2}, M.-T. Walach³, J. A. Carter¹, H. Sangha¹, and B. J.

Anderson⁴

Corresponding author: S. E. Milan, Department of Physics and Astronomy, University of Leicester, Leicester LE1 7RH, UK. (steve.milan@le.ac.uk)

¹Department of Physics and Astronomy,
University of Leicester, Leicester, UK.

²Birkeland Centre for Space Science,
University of Bergen, Bergen, Norway.

³Physics Department, Lancaster
University, Lancaster, UK.

⁴Johns Hopkins University Applied
Physics Laboratory, Laurel, Maryland,
USA.

Abstract.

We study the role of substorms and steady magnetospheric convection (SMC) in magnetic flux transport in the magnetosphere, using observations of field-aligned currents (FACs) by the Active Magnetosphere and Planetary Electrodynamics Response Experiment (AMPERE). We identify two classes of substorm, with onsets above and below 65° magnetic latitude, which display different nightside FAC morphologies. We show that the low-latitude onsets develop a poleward-expanding auroral bulge, and identify these as substorms that manifest ionospheric convection-braking in the auroral bulge region [Grocott et al., 2009]. We show that the high-latitude substorms, which do not experience braking, can evolve into SMC events if the interplanetary magnetic field (IMF) remains southwards for a prolonged period following onset. We conclude that during periods of ongoing driving, the magnetosphere displays repeated substorm activity or SMC depending on the rate of driving and the open magnetic flux content of the magnetosphere prior to onset. We speculate that sawtooth events are an extreme case of repeated onsets, and that substorms triggered by northward-turnings of the IMF mark the cessation of periods of SMC. Our results provide a new explanation for the differing modes of response of the terrestrial system to solar wind-magnetosphere-ionosphere coupling by invoking friction between the ionosphere and atmosphere.

Key points

- AMPERE observations reveal two classes of substorm: high- and low-latitude onsets which are weak and intense, respectively
- Intense substorms experience convection-braking in the auroral bulge; weak onsets can develop into SMC
- These results suggest a framework within which different magnetospheric modes, including sawtooth events, can be understood

1. Introduction

The dynamics of the magnetosphere are driven primarily by the interaction of the solar wind and embedded interplanetary magnetic field (IMF) with the terrestrial field through the process of magnetic reconnection. During periods of southward-directed IMF this excites the Dungey cycle of circulation – or convection – of the field and plasma within the magnetosphere, in which reconnection at the subsolar magnetopause creates open magnetic flux and reconnection in the magnetotail closes flux again, with a general antisunwards transport of open flux and sunwards return flow of closed flux [Dungey, 1961]. This transport is communicated to the polar ionosphere by an electrical current system linking the magnetopause, ionosphere, and ring current [e.g., Iijima and Potemra, 1976; Cowley, 2000], resulting in an ionospheric twin-cell convection pattern [e.g., Heppner and Maynard, 1987, and references therein], which comprises antisunwards plasma drift in the footprint of open field lines (known as the polar cap) and sunwards plasma drift at lower latitudes.

44 The rate of magnetopause (or dayside) reconnection is controlled by conditions in the
45 solar wind [e.g., *Milan et al.*, 2012, and references therein], most importantly the ori-
46 entation of the IMF. The factors controlling the occurrence and rate of magnetotail (or
47 nightside) reconnection are less well understood, but are thought to be determined by the
48 conditions in the plasmashet and pressure exerted on the magnetotail by the solar wind
49 [e.g., *Boudouridis et al.*, 2003; *Milan et al.*, 2004, 2007; *Hubert et al.*, 2006b]. Dayside
50 and nightside reconnection can occur independently of one another, leading to changes
51 in the open magnetic flux content of the magnetosphere, with attendant changes in the
52 size of the ionospheric polar caps; the flux transport and convection associated with these
53 changes is described by the expanding/contracting polar cap (ECPC) model [e.g., *Siscoe*
54 *and Huang*, 1985; *Cowley and Lockwood*, 1992; *Hubert et al.*, 2006a; *Milan*, 2015].

55 Often, changes in open flux content can be linked with the substorm cycle [e.g., *Lockwood*
56 *and Cowley*, 1992; *Milan et al.*, 2003, 2007; *Lockwood et al.*, 2009]. Substorm growth phase
57 is the accumulation of open flux in the magnetotail lobes by dayside reconnection. The
58 near-Earth neutral line (NENL) model of substorm onset [e.g., *Hones*, 1984; *Baker et*
59 *al.*, 1996] asserts that substorm expansion phase (often referred to as “substorm onset”)
60 corresponds to the formation of a reconnection X-line within the closed flux of the plasma
61 sheet, and that this closed flux must first be pinched off (forming a plasmoid) before
62 reconnection proceeds to close the open flux of the tail lobes. Subsequently, the recovery
63 phase marks the antisunwards motion of the NENL to form a distant neutral line (DNL).
64 The tailward motion of the NENL is thought to be associated with the pile-up of newly
65 closed flux in the near-Earth tail, but what provokes this is unclear. The present study
66 addresses this question.

67 *Milan et al.* [2009a] and *Grocott et al.* [2009] studied the auroral intensity and the
68 convection response of substorms with different onset latitudes, that is substorms that
69 accumulated different amounts of open magnetic flux prior to onset. They found that
70 high-latitude substorms (onset above 65° magnetic latitude) tend to have a weak auroral
71 response but lead to enhanced convection in the nightside auroral zone. On the other
72 hand, low-latitude substorms (onset below 65° magnetic latitude) have a more intense
73 auroral response, but counterintuitively lead to a braking of the convection flow. It was
74 suggested by *Grocott et al.* [2009] that this convection-braking was produced by enhanced
75 conductance in the more intense auroral bulge, a mechanism earlier discussed by *Morelli*
76 *et al.* [1995].

77 At other times the magnetosphere can achieve similar dayside and nightside recon-
78 nection rates, leading to steady magnetospheric convection (SMC) or balanced reconnection
79 intervals (BRI) in which the open magnetic flux content remains uniform for an extended
80 period [e.g., *Sergeev et al.*, 1996; *DeJong et al.*, 2008, 2018; *McWilliams et al.*, 2008;
81 *Kissinger et al.*, 2012, and references therein]. *Sergeev et al.* [1996], *DeJong et al.* [2008],
82 and *Kissinger et al.* [2012] noted that periods of SMC are often bracketed by substorm
83 activity, so *Walach and Milan* [2015] and *Walach et al.* [2017] examined the relationship
84 between substorms and SMC events (SMCs) in more detail. They concluded that some
85 SMCs are substorms that have their expansion phase prolonged by continued southwards
86 IMF, so-called “driven”-substorms, whereas “classic” or “isolated” substorms are those
87 during which the IMF turns northwards shortly after onset. There is also debate as to
88 whether northward-turnings of the IMF can trigger substorm onset (see discussion in
89 *Wild et al.* [2009]). It is the purpose of the current study to reexamine the link between

90 changes in the IMF, substorms and SMCs, and the onset latitude dependence of substorm
91 intensity.

92 To monitor solar wind-magnetosphere coupling, convection, and substorms, we em-
93 ploy measurements of the magnetosphere-ionosphere coupling currents, also known as
94 field-aligned currents (FACs) or Birkeland currents, made by the Active Magnetosphere
95 and Planetary Electrodynamics Response Experiment (AMPERE) [*Anderson et al.*,
96 2000, 2002; *Waters et al.*, 2001; *Coxon et al.*, 2018]. The magnitude of the FACs measured
97 by AMPERE, of which the region 1 and region 2 (R1/R2) currents identified by *Iijima*
98 *and Potemra* [1976] are the main component, are a measure of convection strength and
99 ionospheric conductance [*Coxon et al.*, 2016; *Milan et al.*, 2017], whereas the location of
100 the R1/R2 currents is related to the open magnetic flux content of the magnetosphere
101 [*Iijima and Potemra*, 1978; *Clausen et al.*, 2012]. AMPERE measurements have been
102 used to study the large-scale magnetospheric response to solar wind driving [e.g., *Coxon*
103 *et al.*, 2014a; *Anderson et al.*, 2014, 2018; *Milan et al.*, 2017] and substorms [e.g., *Clausen*
104 *et al.*, 2013a, b; *Coxon et al.*, 2014b; *Murphy et al.*, 2013; *Forsyth et al.*, 2018].

105 *Milan et al.* [2015] applied principal component analysis (PCA) to AMPERE current
106 maps to determine the dominant modes of response of FACs to solar wind driving. Subse-
107 quently, *Milan et al.* [2018] applied PCA separately to the dayside and nightside portions
108 of the polar FAC pattern, allowing the temporal response of currents to magnetopause
109 and magnetotail drivers to be examined. The same technique is employed in the current
110 study.

2. Methodology

111 The Active Magnetosphere and Planetary Electrodynamics Response Experiment (AM-
 112 PERE) [Anderson *et al.*, 2000, 2002; Waters *et al.*, 2001; Coxon *et al.*, 2018] measures
 113 the FAC density in both northern and southern hemispheres, at geomagnetic latitudes
 114 above 40° with a latitudinal resolution of 1° , in 24 magnetic local time (MLT) sectors, at
 115 a cadence of 2 min. The data used in this study cover the period 2010 to 2016.

116 The application of principal component analysis to AMPERE observations has been
 117 described by Milan *et al.* [2015, 2017, 2018]. An automated algorithm fits a circle to
 118 the boundary between the region 1 and 2 current ovals and the current density maps
 119 are transformed to be the same size and centred on the geomagnetic pole. The radius
 120 of the fitted circle, Λ° , measured in degrees of colatitude, is later used as a proxy for
 121 the size of the polar cap. Each current map is then described by two vectors \mathbf{J}^D and
 122 \mathbf{J}^N , being the dayside and nightside portions respectively, each of $M = 440$ elements
 123 (40 colatitudinal bins and 11 MLT sectors centred on noon and midnight). PCA is then
 124 performed separately on the dayside and nightside currents, producing two sets of basis
 125 vectors \mathbf{D}_i and \mathbf{N}_i , $i = 1, 2, 3, \dots, M$, which are the eigenvectors of the covariance matrices
 126 of \mathbf{J}^D and \mathbf{J}^N , respectively. These basis vectors, which we term eigenFACs, are those that
 127 most efficiently describe variations in the observations. There are as many dayside and
 128 nightside eigenFACs as there are elements in the original vectors, but only the first few
 129 are significant. The first dayside and the seven most important nightside eigenFACs are
 130 presented in Figure 1. For the time being, we note that \mathbf{D}_1 and \mathbf{N}_1 resemble the dayside
 131 and nightside portions of the R1 and R2 current systems.

Each of the original vectors \mathbf{J}^D or \mathbf{J}^N can be reconstructed as a linear combination of the eigenFACs:

$$\mathbf{J}^D = \sum_{i=1}^m \alpha_i \mathbf{D}_i, \quad \mathbf{J}^N = \sum_{i=1}^m \beta_i \mathbf{N}_i, \quad (1)$$

where α_i and β_i are coefficients which can be determined by finding the projection of \mathbf{D}_i and \mathbf{N}_i on \mathbf{J}^D and \mathbf{J}^N :

$$\alpha_i = \mathbf{J}^D \cdot \mathbf{D}_i, \quad \beta_i = \mathbf{J}^N \cdot \mathbf{N}_i. \quad (2)$$

132 For the reconstruction to be exact, all eigenFACs must be included in the summations,
 133 that is $m = M$, though in practice reasonable fidelity can be achieved with $m \ll M$.
 134 The coefficients α_i and β_i then provide a means of quantifying a complex dataset using
 135 a handful of numbers, a technique known as dimensionality reduction. In this study we
 136 use $m = 1$ on the dayside and $m = 7$ on the nightside. The significance of each eigenFAC
 137 (the amount of variance in the original data that it represents) is given by the ratio of
 138 its corresponding eigenvalue to the sum of the eigenvalues of all the eigenFACs, indicated
 139 in Figure 1 of *Milan et al.* [2018]. As shown in that figure, eigenFACs \mathbf{N}_1 to \mathbf{N}_7 capture
 140 $\sim 80\%$ of the variance in the nightside FAC patterns, with the contribution of the first
 141 10 eigenFACs to the variance being 45.8, 9.2, 6.3, 5.2, 4.1, 3.3, 3.0, 2.7, 2.5, and 2.2%.
 142 There is no clear change in significance between \mathbf{N}_7 and \mathbf{N}_8 , and the choice of this cut-
 143 off in our analysis is somewhat arbitrary; however, we find that no new information for
 144 the present study is contributed by including \mathbf{N}_8 and beyond. We expect α_1 and β_1 to
 145 quantify the strengths of the dayside and nightside portions of the R1/R2 current system
 146 in each AMPERE FAC map. As we will demonstrate, β_3 to β_7 are related to substorm
 147 phenomena.

148 Supporting data are provided by the 1-min OMNI dataset [*King and Papitashvili, 2005*],
 149 the SuperMAG geomagnetic index dataset, including SML, SMU, and SMR, equivalent to
 150 AL, AU, and SYM-H [*Newell and Gjerloev, 2011, 2012; Gjerloev, 2012*], and the substorm
 151 onset list derived from SuperMAG observations [*Newell and Gjerloev, 2011*].

3. Observations and Discussion

152 This paper focusses on three aspects of solar wind-magnetosphere coupling, substorms,
 153 and substorm-related FACs: how does the FAC response vary with substorm onset lati-
 154 tude?; what is the relationship between substorms and steady magnetospheric convection?;
 155 and what do the FAC systems tell us about magnetosphere-ionosphere coupling during
 156 substorms? We investigate these three themes in turn.

157 Figure 2 presents a superposed epoch analysis of substorms, from 2 hours before to 6
 158 hours after substorm onset. Panels (a) to (c) show IMF B_Z , the electrojet indices SML
 159 and SMU, and the radius of the boundary between the R1/R2 current ovals, Λ° , a proxy
 160 for polar cap size. Panels (d) to (j) show the coefficients α_1 and β_1 to β_7 , neglecting β_2 .
 161 As described by *Milan et al. [2018]*, β_2 quantifies IMF B_Y asymmetries in the nightside
 162 FACs and is not of interest to the present study. The substorms are categorised by the
 163 value of Λ at the time of onset, $\Lambda(t = 0)$ or Λ_0 , and the corresponding traces colour-coded
 164 from $\Lambda_0 = 18^\circ$ (black) to $\Lambda_0 = 26^\circ$ (red) in steps of 1° . In total, 8896 substorms are
 165 included in the analysis. For clarity, the traces do not show the standard error on the
 166 mean, though these are in general low due to the relatively large number of substorms in
 167 each category.

168 On average, IMF B_Z is close to zero or negative throughout the period of analysis.
 169 This is because substorms tend to occur during periods of southwards IMF which lead

170 to magnetopause reconnection, and substorm growth phase. More negative values of
171 B_Z are associated with lower-latitude (higher- Λ_0) onsets. That is, substorms tend to
172 occur on an expanded auroral oval, corresponding to high polar cap flux, when B_Z is
173 strongly southwards. B_Z becomes more negative as onset is approached, associated with
174 substorm growth phase, and less negative afterwards, as there is no longer a requirement
175 for continued creation of open flux after the substorm has commenced. The SMU and
176 SML indices (the SuperMAG equivalents of the AU and AL electrojet indices) show
177 substorm growth, expansion, and recovery phase signatures, as expected, though the
178 magnitude of the variations are larger for high- Λ_0 onsets. In all categories except the
179 lowest- Λ_0 substorms, Λ increases prior to onset, a signature of growth phase, and decreases
180 thereafter. The beginning of the contraction of the polar cap can be delayed by almost
181 an hour after onset for the low- Λ_0 substorms. Panels (d) and (e) show the strength of
182 the dayside and nightside R1/R2 currents, as quantified by α_1 and β_1 , which quantify the
183 rate of convection on the dayside and nightside [e.g., *Milan, 2013; Clausen et al., 2013a*].
184 The magnitude of the R1/R2 FACs is well-ordered by Λ_0 , indicating that magnetospheric
185 convection is enhanced for low-latitude onsets. The dayside R1/R2 tends to grow during
186 the growth phase, and then steps up following onset, before decaying after a few 10s of
187 minutes. The nightside R1/R2 FACs remain roughly constant during the growth phase
188 but again increase around the time of onset. These results are broadly consistent with
189 previous studies of the latitude dependence of substorms [e.g., *Milan et al., 2009a; Clausen*
190 *et al., 2013b; Coxon et al., 2014b*].

191 Panels (f) to (j) show the nightside response of the FACs as quantified by β_3 to β_7 .
192 We defer a discussion of what the eigenFACs \mathbf{N}_3 to \mathbf{N}_7 actually signify until later in

193 the paper. For the time-being we note that all five parameters show substorm-related
194 variations, i.e. their behaviour shows marked changes before, during, and after onset.
195 Their variations are also ordered by Λ_0 : specifically, there appear to be two classes of
196 behaviour displayed by low- and high- Λ_0 substorms. For instance, for $\Lambda_0 < 21^\circ$ (black to
197 cyan traces) β_4 decreases from 0 in the 30 mins before onset and increases back to 0 in
198 the 30 mins after onset; for $\Lambda_0 > 21^\circ$ (green to red traces) β_4 shows little variation prior
199 to onset, but increases for an hour or so after onset. Similar, clear differences between
200 these two latitudinal classes can be seen in the variations of β_6 and β_7 .

201 *Grocott et al.* [2009] also identified two classes of substorm: those that experience
202 convection-braking (onsets below 65° magnetic latitude) and those that don't (onsets
203 above 65°). Our value of Λ_0 of 21° is consistent with 65° at midnight, as the auroral
204 oval is on average displaced antisunwards from the geomagnetic pole by 4° . Their inter-
205 pretation was that enhanced conductance, associated with enhanced auroral luminosity
206 for low-latitude onsets [*Milan et al.*, 2009a], leads to frictional coupling between the iono-
207 sphere and atmosphere such that the convection is arrested. Following on from the results
208 of *Grocott et al.* [2009], in the remainder of this study we assume that our low-latitude
209 category of onsets experience convection-braking, whereas our high-latitude onsets do
210 not. We will go on to demonstrate that high-latitude onsets can evolve into periods of
211 SMC, but that low-latitude onsets cannot. (We note that *DeJong et al.* [2018] presented
212 a counter-example to this hypothesis, in which a case study of the conductance during an
213 isolated substorm and an SMC event indicated higher conductance during the latter.)

214 We next investigate the nature of convection associated with substorms during which
215 IMF B_Z remains southwards for different lengths of time following onset. In general

216 $B_Z < 0$ nT during the growth phase – it is generally accepted that ongoing loading of
217 open flux into the magnetosphere is a prerequisite for substorm onset, unless an external
218 perturbation such as a solar wind pressure pulse triggers onset [e.g., *Boudouridis et al.*,
219 2003; *Milan et al.*, 2004; *Hubert et al.*, 2006b] – but once onset has commenced, the IMF
220 orientation can change. Figure 3 presents a superposed epoch analysis of substorms in
221 the same format as Fig. 2 (though note that the vertical scales differ between the two
222 figures). In this analysis, substorms are categorised by the length of time that IMF B_Z
223 remains negative after onset. In each category, we require that $B_Z < -2$ nT for 90%
224 of data points from 20 mins before onset to 30, 60, 90, ..., 360 mins after onset, with
225 traces colour-coded from black to red. Substorms which do not match these criteria are
226 indicated by a dashed line.

227 The resulting B_Z traces show the expected variation, becoming increasingly negative
228 prior to onset, and then remaining negative for a different length of time post-onset before
229 turning positive. The corresponding SMU and SML traces show the expected substorm
230 growth, expansion, and recovery phases, except that the duration of the substorm bay
231 in SML is prolonged by the length of time that B_Z remains southwards. The radius of
232 the current ovals, Λ , increases during the growth phase prior to onset and begins to fall
233 about 20 mins after onset, but remains elevated for the duration of the B_Z -southwards
234 phase. Similar behaviours are seen for the dayside and nightside R1/R2 FAC magnitudes
235 as quantified by α_1 and β_1 : increasing before onset and remaining elevated during the
236 period of southwards IMF before falling to pre-growth phase levels, that is, convection
237 strength increases during the growth phase and is maintained while the magnetosphere

continues to be driven. At this point, with regards to the variations of β_3 to β_7 , we note that these are similar to each other for all categories of B_Z -southwards duration.

On the face of it, these results would seem to support the conclusions of *Walach and Milan* [2015], that continuing southwards IMF after substorm onset can lead to a period of steady magnetospheric convection, which only subsides once the IMF turns northwards. However, we have not considered the possibility that with continued southwards IMF a series of substorms could be triggered, and that averaging over many such substorms could lead to the results presented in Fig. 3. To investigate further, we repeat the superposed epoch analysis, but now limit the events to those substorms for which there is no subsequent substorm in the following 6 hours. This significantly reduces the number of events in the analysis, so we relax our B_Z criterion to be that $B_Z < -1$ nT (rather than $B_Z < -2$ nT) for 90% of data points from 20 mins before onset to 30, 60, 90, ..., 360 mins after onset. The results are presented in Figure 4.

The B_Z traces are similar to Fig. 3, though B_Z is not as negative as before. The SML traces for each category are similar to each other, indicating a substorm bay that lasts 90 mins in each case – a single expansion phase lasting approximately 1 hour irrespective of the duration of the B_Z -southwards phase. However, after the expansion and contraction of the polar cap associated with the onset, Λ remains elevated for the duration of the B_Z -southward phase. Similarly, the dayside and nightside R1/R2 current magnitudes are also elevated for the duration of the B_Z -southwards phase. These results do appear to confirm the conclusions of *Walach and Milan* [2015], that nightside reconnection can be maintained at the end of the expansion phase of a substorm, and steady convection can ensue, if the IMF remains southwards.

261 We now compare other differences between the substorms of Figs. 3 and 4. Firstly,
 262 Fig. 3 has an average value of $\Lambda_0 > 21^\circ$ while for Fig. 4, on average $\Lambda_0 < 21^\circ$. The
 263 difference in the two average values is marginal, but does place the two sets of substorms
 264 in the high- and low- Λ_0 categories discussed in relation to Fig. 2. Moreover, in Fig. 4
 265 both β_4 and β_6 become negative at the time of onset and shortly afterwards, respectively,
 266 whereas this negative excursion is not so significant in Fig. 3. This reinforces the link
 267 between the high- and low- Λ_0 categories and Figs. 3 and 4, respectively, that is the dip
 268 in β_4 and β_6 distinguishes those substorms that do not experience convection-braking at
 269 onset from those that do.

270 Figure 5 presents a schematic of the two scenarios we envisage for substorms occurring
 271 during prolonged periods of solar wind-magnetosphere coupling, with high-latitude sub-
 272 storm onsets on the left, panels (a)-(c), and low-latitude onsets on the right. The figure
 273 has a format similar to Fig. 3 of *Cowley and Lockwood* [1992]. Panels (a) and (d) show
 274 substorm growth phase for the two cases, followed by the expansion phase in panels (b)
 275 and (e). We suggest that substorms that can transition to periods of SMC (panel (c))
 276 are those that do not experience convection-braking, whereas substorms that do experi-
 277 ence braking cannot lead to a laminar convection state, but must result in a sequence of
 278 onsets if the IMF remains southwards (panel (f)). An expected consequence of convection-
 279 braking is the formation of a pronounced substorm auroral bulge following onset, with a
 280 significant poleward motion of the nightside open/closed field line boundary (OCB), as
 281 magnetotail reconnection erodes the open flux of the polar cap (panel (e)). As the bulge
 282 begins to dim and the brake is released, the substorm enters a recovery phase in which
 283 the polar cap returns to a circular shape through the redistribution of open and closed

284 flux in the ionosphere (panel (f)), before the cycle begins again (panel (d)). Conversely,
285 we expect that substorms with no convection-braking can maintain a roughly circular
286 polar cap through continuous redistribution of flux, such that the substorm appears as
287 a brightening of the nightside auroral oval rather than a poleward-growing bulge (panel
288 (b)). There is evidence in the wideband imaging camera (WIC) observations of Fig. 5 of
289 *Milan et al.* [2009a] to support this suggestion that low-latitude onsets have a much more
290 significant poleward progression of the substorm auroras than high-latitude onsets, and
291 this is also consistent with the nightside auroral observations of an isolated substorm and
292 an SMC event presented by *DeJong et al.* [2018].

293 Convection-braking is also expected to have ramifications for the dynamics in the mag-
294 netotail. The substorm onset marks the formation of a near-Earth neutral line (NENL)
295 in the closed plasma sheet [e.g., *Hones*, 1984; *Baker et al.*, 1996]. Once the NENL has
296 pinched off the closed flux, open flux is closed and the polar cap contracts. If the redis-
297 tribution of magnetic flux in the ionosphere required by the ECPC is unimpeded, then
298 the NENL reconnection rate can adjust to match the dayside reconnection rate and a
299 period of SMC, that is a balanced reconnection interval (BRI), can ensue. On the other
300 hand, if convection-braking occurs, ongoing tail reconnection will lead to the formation
301 of a poleward-progressing auroral bulge. This will be associated with flux pileup in the
302 near-tail, as newly-closed flux cannot convect sunwards, and this pileup will push the
303 NENL down-tail until reconnection ceases. Reconnection can only recommence by the
304 formation of a new NENL within the region of newly-closed field lines, signalled by a new
305 substorm AL bay. In this manner, a sequence of onsets is required if the IMF remains
306 directed southwards. In both substorm and SMC cases, once the IMF turns northwards

307 dayside reconnection ceases but nightside reconnection continues until the tail reaches a
308 relaxed configuration.

309 So far we have discussed statistical results. We now present some case examples. To
310 aid with event selection, we developed an algorithm to identify potential periods of SMC
311 from the SMU and SML indices, using criteria similar to *McWilliams et al.* [2008] and
312 *Kissinger et al.* [2012]. We then discarded events during which the IMF was variable or
313 the FAC ovals showed large changes in radius (see also *Walach and Milan* [2015]). Many
314 events were found, of which some typical examples are shown in Figures 6 to 8. We first
315 discuss the 18-hour period beginning 04 UT, 6 October 2010, presented in Fig. 6.

316 Panels (a) and (b) show AMPERE FAC densities along the dawn-dusk meridians of
317 the northern and southern hemispheres, in which the upward/downward FAC pairs (the
318 R1/R2 FACs) can be seen at dawn and dusk. The radii of the FAC ovals, Λ , are shown
319 in panel (c), followed by IMF B_Y and B_Z and solar wind speed and density in panels
320 (d) and (e). Panel (f) shows the dayside reconnection rate, Φ_D , estimated from the solar
321 wind observations using eq. (15) of *Milan et al.* [2012], and the time integral of Φ_D . This
322 integral shows the amount of open flux that would accumulate in the polar caps if no
323 nightside reconnection took place. Typically the polar caps contain 0.5 GWb of open
324 flux, rising to ~ 1 GWb during extreme conditions [*Milan et al.*, 2007]. Below this are:
325 panel (g) the SML and SMU indices, panel (h) the SMR ring current index, and panel
326 (i) the PC index which measures convection strength in the polar regions. Vertical green
327 lines indicate substorm onset identified by SuperMAG. Vertical red lines, labelled i , ii ,
328 etc., indicate times which will be discussed below; if a red line corresponds to a substorm
329 onset, it has been displaced slightly for clarity.

330 Two events occurred during this time period. Between i and $viii$, the IMF was predom-
 331 inantly directed southwards (panel (d)), the R1/R2 FACs were enhanced (panels (a) and
 332 (b)), and the PC index was elevated (panel (i)). Following the southwards turning at i ,
 333 dayside reconnection was elevated leading to expansions of the polar caps (panel (c)). At
 334 ii , SuperMAG identified a substorm onset. Thereafter the IMF remained southwards until
 335 iv , and between iii and iv SMC ensued, indicated by the horizontal orange bar, during
 336 which SMU and SML were approximately constant (panel (g)), the FAC radius remained
 337 uniform, and PC indicated steady convection. By v the IMF had turned southwards
 338 again, growth phase signatures were seen in SMU/SML and Λ , followed by a substorm
 339 onset at vi . Associated with the substorm bay, the polar caps initially contracted, but
 340 by vii they stabilised and varied only gradually during a second period of SMC, again
 341 accompanied by steady PC. At $viii$ the IMF rotated so that it was no longer southward,
 342 and the SMC petered out. We consider both these periods of SMC to be examples of the
 343 driven-substorm SMC described by *Walach and Milan* [2015]. We note that Λ remained
 344 below 21° throughout almost the entire period.

345 To summarize these two events, we identify the following intervals as: ($i - ii$) growth
 346 phase, ($ii - iii$) expansion phase, ($iii - iv$) SMC, ($iv - v$) recovery phase, ($v - vi$) growth
 347 phase, ($vi - vii$) expansion phase, ($vii - viii$) SMC, ($viii -$) recovery phase. In both of
 348 these cases, $\int \Phi_D dt \sim 0.1$ GWb of open flux accumulated during the growth phase of the
 349 substorm and ~ 0.1 GWb during the expansion phase. During the two periods of SMC,
 350 ~ 0.15 and ~ 0.5 GWb of open flux were open and closed, that is, transported through
 351 the system, in the latter case equivalent to refreshing a typical polar cap.

352 Fig. 7 presents the 28-hour period beginning 19 UT, 28 May 2010. At i the IMF
 353 turned southwards and remained so for almost 23 hours. Following i the the polar caps
 354 expanded (Λ), before a substorm onset was detected at ii , which by iii developed into a
 355 period of SMC. IMF B_Z became increasingly negative such that the dayside reconnection
 356 rate increased and exceeded the nightside rate resulting in gradually expanding polar caps
 357 (Λ) and stronger convection (SML/SMU and PC). Around iv , Λ grew beyond 21° and
 358 thereafter multiple substorm onset or substorm intensification signatures were identified
 359 by SuperMAG. By v , the IMF was no longer so strongly southwards, the polar caps
 360 had contracted, and SMC resumed, until the IMF turned northwards at vi . A substorm
 361 occurred at this time and the polar caps rapidly contracted. This event shows that SMC
 362 can occur when the polar caps are contracted, but if the radii grow too large, repeated
 363 substorm activity results. We also note that this period of elevated polar cap size is
 364 associated with an enhanced ring current (SMR, similar to SYM-H), as suggested by
 365 *Milan et al.* [2009b]. $\int \Phi_D dt \sim 0.2$ GWb during both the growth and expansion phases of
 366 the initial substorm, and ~ 0.8 and ~ 0.4 GWb during the two phases of SMC. Between
 367 iv and v , $\int \Phi_D dt \sim 4.5$ GWb, or approximately 0.2 GWb for each intensification. It is
 368 debatable if each substorm onset identified by SuperMAG in the interval iv to v is a true
 369 substorm or rather a substorm intensification; however, this clearly is not a period of
 370 SMC, and SML indicates intense fluctuations in nightside precipitation which would be
 371 expected to give rise to convection-braking.

372 Finally, Fig. 8 presents two similar events: the 9-hour period after 10 UT, 4 September
 373 2011, and the 13-hour period after 00:30 UT, 21 January 2012. In both examples, during
 374 ongoing southward IMF growth phase signatures were observed between i and ii , followed

375 by substorm onset at *ii*, transitioning into SMC at *iii*, and then ending as the IMF
 376 turned northwards at *iv*. In both cases a substorm onset was observed at *iv*, which lead
 377 to rapid contractions of the polar caps. Several studies have indicated that periods of
 378 SMC often end with a substorm [e.g., *Sergeev et al.*, 1996], and these examples (and
 379 arguably that in Fig. 7) conform to this. In none of these cases is there a clear solar wind
 380 cause for the triggering of a substorm, except a reduction in the dayside reconnection
 381 rate. We suggest that if the tail is in a stressed state at the end of a period of SMC,
 382 a substorm can be triggered to return it to a relaxed state. In all three cases the solar
 383 wind density is relatively high, exceeding 10 cm^{-3} , which may exacerbate this stressed
 384 state. Interestingly, such cases may inform the debate regarding the existence of substorms
 385 triggered by northward turnings of the IMF [e.g., *Wild et al.*, 2009, and references therein].
 386 In both these cases, $\int \Phi_D dt \sim 0.1 \text{ GWb}$ during the growth and expansion phases of the
 387 initial substorms (that is, $\sim 0.2 \text{ GWb}$ associated with each substorm), and ~ 0.3 and
 388 $\sim 0.7 \text{ GWb}$ during the two phases of SMC.

389 We now turn to our final question, regarding the nature of the FAC patterns associated
 390 with substorm onset, and especially the difference between the high- and low- Λ_0 onsets.
 391 Remembering that a general FAC pattern can be reproduced as a linear combination of
 392 the eigenFACs, we briefly describe the FAC morphology associated with \mathbf{N}_i , $i = 1, \dots, 7$,
 393 and their contribution to the summation in eq. (1).

394 $\beta_1 \mathbf{N}_1$: The nightside portion of the large-scale R1/R2 current system. β_1 is always
 395 found to be positive, as this corresponds to the expected polarity of the R1/R2 FACs.
 396 This eigenFAC is roughly symmetric about the midnight meridian, but we note that

397 the upward FACs link up across midnight, the average configuration associated with the
 398 Harang discontinuity [e.g., *Iijima and Potemra*, 1978].

399 $\beta_2\mathbf{N}_2$: As discussed by *Milan et al.* [2018], this eigenFAC controls the local time at which
 400 the polarities of the R1/R2 currents reverse, which we can identify with the location of
 401 the nightside convection throat. For $\beta_2 > 0$ and $\beta_2 < 0$ the convection throat moves pre-
 402 and post-midnight, respectively. *Milan et al.* [2018] showed that the value of β_2 is related
 403 to the B_Y component of the IMF.

404 $\beta_3\mathbf{N}_3$: This eigenFAC controls how the pre- and post-midnight portions of the R1/R2
 405 FACs link up across midnight. If $\beta_3 < 0$ then the strength of the upwards current in the
 406 Harang discontinuity region is enhanced. If $\beta_3 > 0$ then upwards current is diminished or
 407 downwards current intrudes into this region. We observe that β_3 tends to be positive for
 408 high- Λ_0 substorms, which is consistent with Fig. 15 of *Iijima and Potemra* [1978].

409 $\beta_4\mathbf{N}_4$: When $\beta_4 > 0$ this eigenFAC leads to a strengthening and poleward motion of the
 410 R1/R2 FACs, especially at midnight and in the post-midnight sector; $\beta_4 < 0$ leads to a
 411 thinning of these currents. That $\beta_4 > 0$ and $\beta_4 < 0$ for high- and low- Λ_0 substorms (e.g.,
 412 Fig. 2) is consistent with our assertion that the auroral bulge is enhanced and protrudes
 413 polewards during high- Λ_0 substorms.

414 $\beta_5\mathbf{N}_5$: In the midnight sector this eigenFAC is morphologically similar to \mathbf{N}_3 (though of
 415 opposite polarity), so we expect it to play a role in modulating the Harang discontinuity
 416 currents.

417 $\beta_6\mathbf{N}_6$: When $\beta_6 > 0$ this eigenFAC contributes upwards FAC at high latitudes, espe-
 418 cially in the midnight and post-midnight regions, and in this respect is similar to \mathbf{N}_4 .

419 $\beta_7 \mathbf{N}_7$: When $\beta_7 > 0$ this eigenFAC contributes upwards FAC at high latitudes, across
420 both pre- and post-midnight regions.

421 In summary, those eigenFACs which tend to be enhanced during high- Λ_0 substorms
422 contribute FAC at high latitudes in the pre-, post-, and midnight regions, consistent
423 with our expectations that these substorms have an enhanced auroral bulge that will
424 lead to convection-braking; indeed, the polewards-growth of the bulge is a consequence of
425 this convection-braking, requiring a poleward motion of the ionospheric projection of the
426 nightside reconnection X-line as open magnetic flux is eroded.

4. Conclusions

427 We have examined the field-aligned current strength and morphology during substorms,
428 using observations from AMPERE, focussing on two main questions: what influence does
429 onset latitude have on the FAC response?; and what is the relationship between substorm
430 onset, prolonged IMF B_Z -southward conditions, and steady magnetospheric convection?

431 *Milan et al.* [2009a] demonstrated that substorms occurring at low latitudes (high- Λ_0
432 substorms in the terminology of this paper) are more intense than high-latitude sub-
433 storms, and *Grocott et al.* [2009] demonstrated that these experience convection-braking,
434 possibly associated with the high conductance of the bright auroral bulge [e.g., *Morelli*
435 *et al.*, 1995]. *Walach and Milan* [2015] showed that a significant number of steady mag-
436 netospheric convection events are initially substorms, but substorms for which the IMF
437 remains southwards for a prolonged period after onset. Our results support both of these
438 conclusions, but we go further to suggest that those substorms which can evolve into
439 SMC are those that occur at high latitudes and do not experience convection-braking, as
440 illustrated in Fig. 5(a)-(c). In this case, once a substorm commences, associated with the

441 onset of magnetic reconnection in the tail at a near-Earth neutral line (NENL), that re-
442 connection can persist if new open flux continues to be supplied by dayside reconnection.
443 Typical substorm signatures, such as the SML bay and substorm-associated FAC mor-
444 phologies, last 60 to 90 mins after onset, but these die away even if NENL reconnection
445 continues thereafter. We suggest that substorms which experience braking and associated
446 flux pile-up in the near-tail pushing the NENL down-tail, cannot segue into a laminar
447 convection state, and instead a staccato sequence of substorm onsets results, each with
448 a recovery phase that represents the release of the brake, as illustrated in Fig. 5(d)-(f).
449 Substorms that experience braking will be those that develop poleward-growing auroral
450 bulges, whereas high-latitude, non-braking substorms will have less-pronounced bulges,
451 maintaining a circular polar cap through frictionless redistribution of magnetic flux.

452 In the examples presented, between 0.2 and 0.4 GWb of open flux transport were
453 associated with the growth and expansion phases of each precursor substorm, with between
454 0.15 and 0.8 GWb during the following period of SMC. This latter value depends on how
455 long the IMF remains southwards following the initial onset, that is, the duration of
456 the SMC. When the polar caps grew sufficiently large that a sequence of substorms or
457 intensifications was triggered, each effected 0.2 GWb of flux transport. We note that
458 periods of SMC can lead to a complete refreshment of the open flux of the polar caps,
459 that is straight through-put of open flux from the dayside to the nightside X-lines and
460 convection from the dayside OCB to the nightside OCB.

461 In the example presented in Fig. 7, repeating substorms occurred with a repetition
462 rate of ~ 30 min. Indeed, these may not be true substorms, but substorm intensifications
463 caused by convection-braking. We speculate that sawtooth events, ~ 3 hr quasi-periodic

464 intense substorms [*Belian et al.*, 1995], may be an extreme case of braking substorms
465 occurring during strongly-driven intervals associated with geomagnetic storms [*Walach*
466 *and Milan*, 2015; *Walach et al.*, 2017]. In this case, we would place SMCs, repeating
467 substorms, and sawtooth events as a spectrum of responses to periods of prolonged low to
468 high solar wind-magnetosphere coupling; we note that this spectrum of behaviour agrees
469 with the ordering of *Hubert et al.* [2017]. On the other hand, isolated substorms are
470 associated with periods when the IMF is sporadically turning northwards and southwards.

471 We have not addressed the question of why some substorms commence at high latitudes
472 and others at low latitudes, though *Milan et al.* [2009b] suggested that this was associated
473 with the magnetic perturbation produced by the ring current dipolarizing the near-Earth
474 tail. This will be investigated further in a subsequent study.

475 **Acknowledgments.** SEM and JAC were supported by the Science and Technology Fa-
476 cilities Council (STFC), UK, grant no. ST/N000749/1; HS was supported by an STFC stu-
477 dentship. MTW was supported by the Natural Environment Research Council (NERC),
478 UK, grant no. NE/P001556/1. The work at the Birkeland Centre for Space Centre, Uni-
479 versity of Bergen, Norway, was supported by the Research Council of Norway/CoE un-
480 der contract 223252/F50. We thank the AMPERE team and the AMPERE Science
481 Center for providing the Iridium-derived data products; AMPERE products are avail-
482 able at <http://ampere.jhuapl.edu>. The OMNI data, including solar wind parameters
483 and geomagnetic indices, were obtained from the GSFC/SPDF OMNIWeb interface at
484 <http://omniweb.gsfc.nasa.gov>. The SuperMAG indices and substorm list was downloaded
485 from <http://supermag.jhuapl.edu>. For the ground magnetometer data from which these
486 were derived, we gratefully acknowledge: Intermagnet; USGS, Jeffrey J. Love; CARISMA,

487 PI Ian Mann; CANMOS; The S-RAMP Database, PI K. Yumoto and Dr. K. Shiokawa;
488 The SPIDR database; AARI, PI Oleg Troshichev; The MACCS program, PI M. Engebret-
489 son, Geomagnetism Unit of the Geological Survey of Canada; GIMA; MEASURE, UCLA
490 IGPP and Florida Institute of Technology; SAMBA, PI Eftyhia Zesta; 210 Chain, PI K.
491 Yumoto; SAMNET, PI Farideh Honary; The institutes who maintain the IMAGE mag-
492 netometer array, PI Eija Tanskanen; PENGUIN; AUTUMN, PI Martin Connors; DTU
493 Space, PI Dr. Juergen Matzka; South Pole and McMurdo Magnetometer, PI's Louis J.
494 Lanzarotti and Alan T. Weatherwax; ICESTAR; RAPIDMAG; PENGUIn; British Artarc-
495 tic Survey; MacMac, PI Dr. Peter Chi; BGS, PI Dr. Susan Macmillan; Pushkov Institute
496 of Terrestrial Magnetism, Ionosphere and Radio Wave Propagation (IZMIRAN); GFZ, PI
497 Dr. Juergen Matzka; MFGI, PI B. Heilig; IGFPAS, PI J. Reda; University of L'Aquila,
498 PI M. Vellante; SuperMAG, PI Jesper W. Gjerloev.

References

- 499 Anderson, B. J., Takahashi, K., & Toth, B. A. (2000). Sensing global Birkeland currents
500 with Iridium engineering magnetometer data. *Geophysical Research Letters*, *27*, 4045–
501 4048. doi: 10.1029/2000GL000094
- 502 Anderson, B. J., Takahashi, K., Kamei, T., Waters, C. L., & Toth, B. A. (2002).
503 Birkeland current system key parameters derived from Iridium observations: Method
504 and initial validation results. *Journal of Geophysical Research*, *107*, 1079. doi:
505 10.1029/2001JA000080
- 506 Anderson, B. J., Korth, H., Waters, C. L., Green, D. L., Merkin, V. G., Barnes, R. J.,
507 & Dyrud, L. P. (2014). Development of large-scale Birkeland currents determined from

- 508 the Active Magnetosphere and Planetary Electrodynamics Response Experiment. *Geo-*
509 *physical Research Letters*, *41*. doi: 10.1002/2014GL059941
- 510 Anderson, B. J., Olson, C. N., Korth, H., Barnes, R. J., Waters, C. L., & Vines, S. K.
511 (2018). Temporal and spatial development of global Birkeland currents. *Journal of Geo-*
512 *physical Research: Space Physics*, *123*. doi: 10.1029/2018JA025254
- 513 Baker, D. N., Pulkkinen, T. I., Angelopoulos, V., Baumjohann, W., & McPherron, R. L.
514 (1996). Neutral line model of substorms: Past results and present view. *Journal of*
515 *Geophysical Research*, *101*, 12975–13010.
- 516 Belian, R. D., Cayton, T. E., & Reeves, G. D. (1995). Quasi-periodic global substorm
517 generated flux variations observed at geosynchronous orbit. In *Space Plasma: Coupling*
518 *Between Small and Medium Scale Processes, Geophys. Monogr. Ser.*, vol. 86, eds. M.
519 Ashour-Abdalla, T. Chang, and P. Dusenbery, pp. 143–148, AGU, Washington, D. C.
520 doi:10.1029/GM086p0143
- 521 Boudouridis, A., Zesta, E., Lyons, R., Anderson, P. C., & Lummerzheim, D. (2003). Effect
522 of solar wind pressure pulses on the size and strength of the auroral oval. *Journal of*
523 *Geophysical Research*, *108*, 8012. doi: 10.1029/2002JA009373
- 524 Clausen, L. B. N., Baker, J. B. H., Ruohoniemi, J. M., Milan, S. E., & Anderson, B. J.
525 (2012). Dynamics of the region 1 Birkeland current oval derived from the Active Mag-
526 netosphere and Planetary Electrodynamics Response Experiment (AMPERE). *Journal*
527 *of Geophysical Research*, *117*, A06233. doi:10.1029/2012JA017666
- 528 Clausen, L. B. N., Baker, J. B. H., Ruohoniemi, J. M., Milan, S. E., Coxon, J. C., Wing,
529 S., Ohtani, S., & Anderson, B. J. (2013a). Temporal and spatial dynamics of the region
530 1 and 2 Birkeland currents during substorms. *Journal of Geophysical Research: Space*

- 531 *Physics*, 118. doi: 10.1002/jgra.50288
- 532 Clausen, L. B. N., Milan, S. E., Baker, J. B. H., Ruohoniemi, J. M., Glassmeier, K.-
533 H., Coxon, J. C., & Anderson, B. J. (2013b). On the influence of open magnetic flux
534 on substorm intensity: ground- and space-based observations. *Journal of Geophysical*
535 *Research: Space Physics*, 118. doi: 10.1002/jgra.50308
- 536 Cowley, S. W. H. (2000). Magnetosphere-ionosphere interactions: A tutorial review. In
537 *Magnetospheric Current Systems, Geophys. Monogr. Ser.*, vol. 118, eds. S. Ohtani et
538 al., pp. 91–106, AGU, Washington, D.C. doi: 10.1029/GM118p0091
- 539 Cowley, S. W. H., & Lockwood, M. (1992). Excitation and decay of solar wind-driven
540 flows in the magnetosphere-ionosphere system. *Annales Geophysicae*, 10, 103–115.
- 541 Coxon, J. C., Milan, S. E., Clausen, L. B. N., Anderson, B. J., & Korth, H. (2014a).
542 The magnitudes of the Birkeland currents observed by AMPERE and their role in
543 solar wind-magnetosphere-ionosphere coupling. *Journal of Geophysical Research: Space*
544 *Physics*, 119. doi: 10.1002/2014JA020138
- 545 Coxon, J. C., Milan, S. E., Clausen, L. B. N., Anderson, B. J., & Korth, H. (2014b). A
546 superposed epoch analysis of the Region 1 and Region 2 Birkeland currents observed
547 by AMPERE during substorms. *Journal of Geophysical Research: Space Physics*, 119.
548 doi: 10.1002/2014JA020500
- 549 Coxon, J. C., Milan, S. E., Carter, J. A., Clausen, L. B. N., Anderson, B. J., & Korth,
550 H. (2016). Seasonal and diurnal variations in AMPERE observations of the Birkeland
551 currents compared to modelled results. *Journal of Geophysical Research: Space Physics*,
552 121. doi: 10.1002/2015JA022050

- 553 Coxon, J. C., Milan, S. E., Korth, H., & Anderson, B. J. (2018). Ampère's Law: a review
554 of Birkeland current research using the Iridium constellation. In *Electric Currents in*
555 *GeoSpace and Beyond, AGU Geophysical Monograph 235*, eds. A. Keiling et al., Wiley
556 and Sons.
- 557 DeJong, A. D., Ridley, A. J., & Clauer, C. R. (2008). Balanced reconnection intervals:
558 Four case studies. *Annales Geophysicae*, *26*, 3897–3912.
- 559 DeJong, A. D., Bell, J. M., & Ridley, A. J. (2018). Comparison of the ionosphere during
560 an SMC initiating substorm and an isolated substorm. *Journal of Geophysical Research:*
561 *Space Physics*, *123*, 4939–4951. <https://doi.org/10.1029/2017JA025055>
- 562 Dungey, J. W. (1961). Interplanetary magnetic fields and the auroral zones. *Physical*
563 *Review Letters*, *6*, 47–48.
- 564 Forsyth, C., Shortt, M., Coxon, J. C., Rae, I. J., Freeman, M. P., Kalmoni, N.
565 M. E., Jackson, C. M., Anderson, B. J., Milan, S. E., & Burrell, A. G. (2018).
566 Seasonal and temporal variations of field-aligned currents and ground magnetic de-
567 flections during substorms. *Journal of Geophysical Research: Space Physics*, *123*.
568 <https://doi.org/10.1002/2017JA025136>
- 569 Gjerloev, J. W. (2012). The SuperMAG data processing technique. *Journal of Geophysical*
570 *Research*, *117*, A09213. doi: 10.1029/2012JA017683.
- 571 Grocott, A., Wild, J. A., Milan, S. E., & Yeoman, T. K. (2009). Superposed epoch analysis
572 of the ionospheric convection evolution during substorms: onset latitude dependence.
573 *Annales Geophysicae*, *27*, 591–600.
- 574 Heppner, J. P., & Maynard, N. (1987). Empirical high-latitude electric field models. *Jour-*
575 *nal of Geophysical Research*, *92*, 4467–4489. doi: 10.1029/JA092iA05p04467

- 576 Hones, E. W., Jr. (1984). Plasma sheet behavior during substorms, in *Magnetic Recon-*
577 *nection in Space and Laboratory Plasmas*, Geophys. Monogr. Ser., vol. 30, edited by E.
578 W. Hones Jr., pp. 178-184, AGU, Washington, D.C.
- 579 Hubert, B., Milan, S. E., Grocott, A., Blockx, C., Cowley, S. W. H., & Gérard, J.-C.
580 (2006a). Dayside and nightside reconnection rates inferred from IMAGE FUV and Super
581 Dual Auroral Radar Network data. *Journal of Geophysical Research*, *111*, A03217. doi:
582 10.1029/2005JA011140
- 583 Hubert, B., Palmroth, M., Laitinen, T. V., Janhunen, P., Milan, S. E., Grocott, A.,
584 Cowley, S. W. H., Pulkkinen, T., & Gérard, J.-C. (2006b). Compression of the Earth's
585 magnetotail by interplanetary shocks drives magnetic flux closure. *Geophysical Research*
586 *Letters*, *33*, L10105. doi: 10.1029/2006GL026008
- 587 Hubert, B., Gérard, J.-C., Milan, S. E., & Cowley, S. W. H. (2017). Magnetic recon-
588 nection during steady magnetospheric convection and magnetospheric modes. *Annales*
589 *Geophysica*, *35*, 505–524. doi: 10.5194/angeo-35-505-2017
- 590 Iijima, T., & Potemra, T. A. (1976). The amplitude distribution of field-aligned currents
591 at northern high latitudes observed by Triad. *Journal of Geophysical Research*, *81*,
592 2165–2174. doi: 10.1029/JA081i013p02165
- 593 Iijima, T., & Potemra, T. A. (1978). Large-scale characteristics of field-aligned cur-
594 rents associated with substorms. *Journal of Geophysical Research*, *83*, 599–615. doi:
595 10.1029/JA083iA02p00599
- 596 King, J. H., & Papitashvili, N. E. (2005). Solar wind spatial scales in and comparisons of
597 hourly Wind and ACE plasma and magnetic field data. *Journal of Geophysical Research*,
598 *110*, A02209. doi: 10.1029/2004JA010804

- 599 Kissinger, J. E., McPherron, R. L., Hsu, T.-S., Angelopoulos, V., & Chu, X. (2012).
600 Necessity of substorm expansions in the initiation of steady magnetospheric convection,
601 *Geophysical Research Letters*, *39*, L15105. doi: 10.1029/2012GL052599
- 602 Lockwood, M., & Cowley, S. W. H. (1992). Ionospheric convection and the substorm cycle.
603 In *Proceedings of the International Conference on Substorms (ICS-1)*, 99–109.
- 604 Lockwood, M., Hairston, M., Finch, I., & Rouillard, A. (2009). Transpolar voltage and
605 polar cap flux during the substorm cycle and steady convection events. *Journal of*
606 *Geophysical Research*, *114*, A01210. doi: 10.1029/2008JA013697
- 607 McWilliams, K. A., Pfeifer, J. B., & McPherron, R. L. (2008). Steady magnetospheric con-
608 vection selection criteria: Implications of global SuperDARN convection measurements,
609 *Geophysical Research Letters*, *35*, L09102 doi: 10.1029/2008GL033671
- 610 Milan, S. E., Lester, M., Cowley, S. W. H., Oksavik, K., Brittnacher, M., Greenwald, R. A.,
611 Sofko, G., & Villain, J.-P. (2003). Variations in polar cap area during two substorm
612 cycles. *Annales Geophysicae*, *21*, 1121–1140.
- 613 Milan, S. E., Cowley, S. W. H., Lester, M., Wright, D. M., Slavin, J. A., Fillingim, M.,
614 Carlson, C. W., and Singer, H. J. (2004). Response of the magnetotail to changes in
615 open flux content of the magnetosphere. *Journal of Geophysical Research*, *109*, A04220.
616 doi: 10.1029/2003JA010350
- 617 Milan, S. E., Provan, G., & Hubert, B. (2007). Magnetic flux transport in the Dungey
618 cycle: A survey of dayside and nightside reconnection rates. *Journal of Geophysical*
619 *Research*, *112*, A01209. doi: 10.1029/2006JA011642
- 620 Milan, S. E., Grocott, A., Forsyth, C., Imber, S. M., Boakes, P. D., & Hubert, B. (2009a).
621 A superposed epoch analysis of auroral evolution during substorm growth, onset and

- 622 recovery: open magnetic flux control of substorm intensity. *Annales Geophysicae*, *27*,
623 659–668.
- 624 Milan, S. E., Hutchinson, J., Boakes, P. D., & Hubert, B. (2009b). Influences on the radius
625 of the auroral oval. *Annales Geophysicae*, *27*, 2913–2924.
- 626 Milan, S. E., Gosling, J. S., & Hubert, B. (2012). Relationship between interplan-
627 etary parameters and the magnetopause reconnection rate quantified from observa-
628 tions of the expanding polar cap. *Journal of Geophysical Research*, *117*, A03226. doi:
629 10.1029/2011JA017082
- 630 Milan, S. E. (2013), Modelling Birkeland currents in the expanding/contracting polar cap
631 paradigm. *Journal of Geophysical Research*, *118*. doi:10.1002/jgra.50393
- 632 Milan, S. E. (2015). Sun et Lumière: solar wind-magnetosphere coupling as deduced
633 from ionospheric flows and polar auroras. In *Magnetospheric Plasma Physics: The*
634 *Impact of Jim Dungey's Research, Astrophysics and Space Science Proceedings 41*, eds.
635 D. Southwood et al., Springer. doi: 10.1007/978-3-319-18359-6_2
- 636 Milan, S. E., Carter, J. A., Korth, H., & Anderson, B. J. (2015). Principal Component
637 Analysis of Birkeland currents determined by the Active Magnetosphere and Plane-
638 tary Electrodynamics Response Experiment. *Journal of Geophysical Research: Space*
639 *Physics*, *120*. doi: 10.1002/2015JA021680
- 640 Milan, S. E., Clausen, L. B. N., Coxon, J. C., Carter, J. A., Walach, M.-T., Laun-
641 dal, K., Østgaard, N., Tenfjord, P., Reistad, J., Snekvik, K., Korth, H., & Ander-
642 son, B. J. (2017). Overview of solar wind-magnetosphere-ionosphere-atmosphere cou-
643 pling and the generation of magnetospheric currents. *Space Science Reviews*, *206*. doi:
644 10.1007/s11214-017-0333-0

- 645 Milan, S. E., Carter, J. A., Sangha, H., Laundal, K., Østgaard, N., Tenfjord, P., Reistad,
646 J., Snekvik, K., Coxon, J. C., Korth, H., & Anderson, B. J. (2018). Timescales of day-
647 side and nightside field-aligned current response to changes in solar wind-magnetosphere
648 coupling. *Journal of Geophysical Research: Space Physics, Journal of Geophysical Re-*
649 *search: Space Physics, 123*. <https://doi.org/10.1029/2018JA025645>.
- 650 Morelli, J. P., Bunting, R. J., Cowley, S. W. H., Farrugia, C. J., Freeman, M. P., Friis-
651 Christensen, E., Jones, G. O. L., Lester, M., Lewis, R. V., Lühr, H., Orr, D., Pinnock,
652 M., Reeves, G. D., Williams, P. J. S., & Yeoman, T. K. (1995). Radar observations of
653 auroral zone flows during a multiple-onset substorm. *Annales Geophysicae, 13*, 1144.
654 <https://doi.org/10.1007/s00585-995-1144-2>
- 655 Murphy, K. R., Mann, I. R., Rae, I. J., Waters, C. L., Frey, H. U., Kale, A., Singer, H. J.,
656 Anderson, B. J., & Korth, H. (2013). The detailed spatial structure of field-aligned cur-
657 rents comprising the substorm current wedge. *Journal of Geophysical Research: Space*
658 *Physics, 118*, 7714–7727. doi: 10.1002/2013JA018979
- 659 Newell, P. T., & Gjerloev, J. W. (2011). Evaluation of SuperMAG auroral electrojet
660 indices as indicators of substorms and auroral power. *Journal of Geophysical Research,*
661 *116*, A12211. doi: 10.1029/2011JA016779
- 662 Newell, P. T. & Gjerloev, J. W. (2012). SuperMAG-Based Partial Ring Current Indices
663 *Journal of Geophysical Research, 117*, A12211. doi:10.1029/2012JA017586
- 664 Sergeev, V. A., Pellinen, R. J., & Pulkkinen, T. I. (1996). Steady magnetospheric convec-
665 tion: A review of recent results. *Space Science Reviews, 75*, 551–604.
- 666 Siscoe, G. L., & Huang, T. S. (1985). Polar cap inflation and deflation. *Journal of Geo-*
667 *physical Research, 90*, 543–547.

- 668 Walach, M.-T., & Milan, S. E. (2015). Are steady magnetospheric convection events
669 prolonged substorms? *Journal of Geophysical Research: Space Physics*, *120*. doi:
670 10.1002/2014JA020631
- 671 Walach, M.-T., Milan, S. E., Murphy, K. R., Carter, J. A., Hubert, B. A., & Grocott, A.
672 (2017). Comparative study of large-scale auroral signatures of substorms, steady mag-
673 netospheric convection events, and sawtooth events. *Journal of Geophysical Research:*
674 *Space Physics*, *122*. doi: 10.1002/2017JA023991.
- 675 Waters, C. L., Anderson, B. J., & Liou, K. (2001). Estimation of global field aligned
676 currents using the Iridium® system magnetometer data. *Geophysical Research Letters*,
677 *28*, 2165–2168.
- 678 Wild, J. A., Woodfield, E. E., & Morley, S. K. (2009). On the triggering of auroral sub-
679 storms by northward turnings of the interplanetary magnetic field. *Annales Geophysicae*,
680 *27*, 3559–3570. www.ann-geophys.net/27/3559/2009/

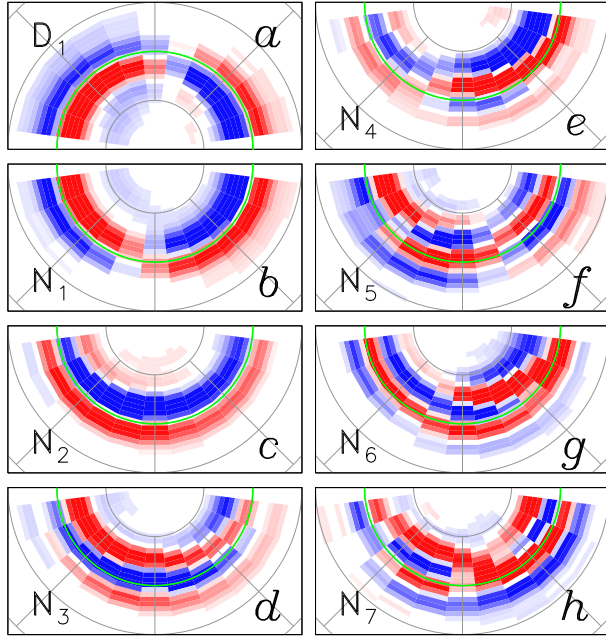


Figure 1. The first dayside (\mathbf{D}_1) and first seven nightside (\mathbf{N}_i , $i = 1, \dots, 7$) eigenFACs derived from AMPERE field-aligned current distributions. Each panel is presented in a magnetic local time and magnetic latitude frame, where magnetic latitudes are scaled such that the boundary between R1 and R2 FACs occurs at 70° (green semicircle). Blue and red colours correspond to positive and negative values.

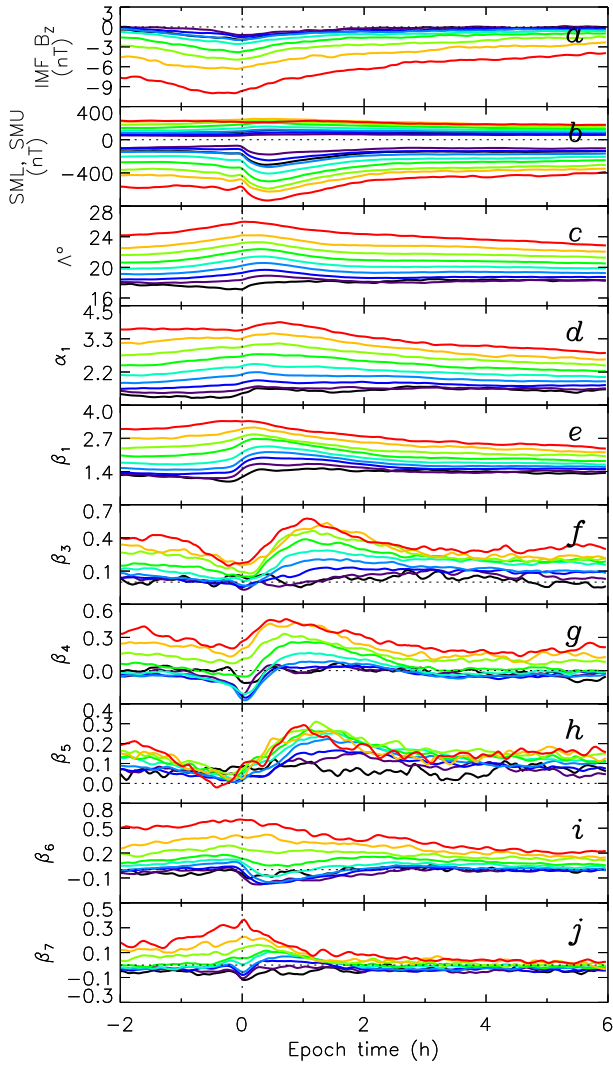


Figure 2. Superposed epoch analysis of substorms from 2 hours prior to 6 hours after onset. (a) IMF B_z , (b) SMU and SML electrojet indices, (c) radius of the AMPERE current pattern, Λ , and (d) to (j) coefficients associated with the eigenFACs presented in Fig. 1 (except N_2). The substorms are categorised by the value of Λ at onset ($t = 0$) or Λ_0 , in 1° steps from 18° (black trace) to 26° (red trace); the number of substorms in each category is 976, 510, 1282, 1681, 1527, 1108, 789, 501, and 522, respectively.

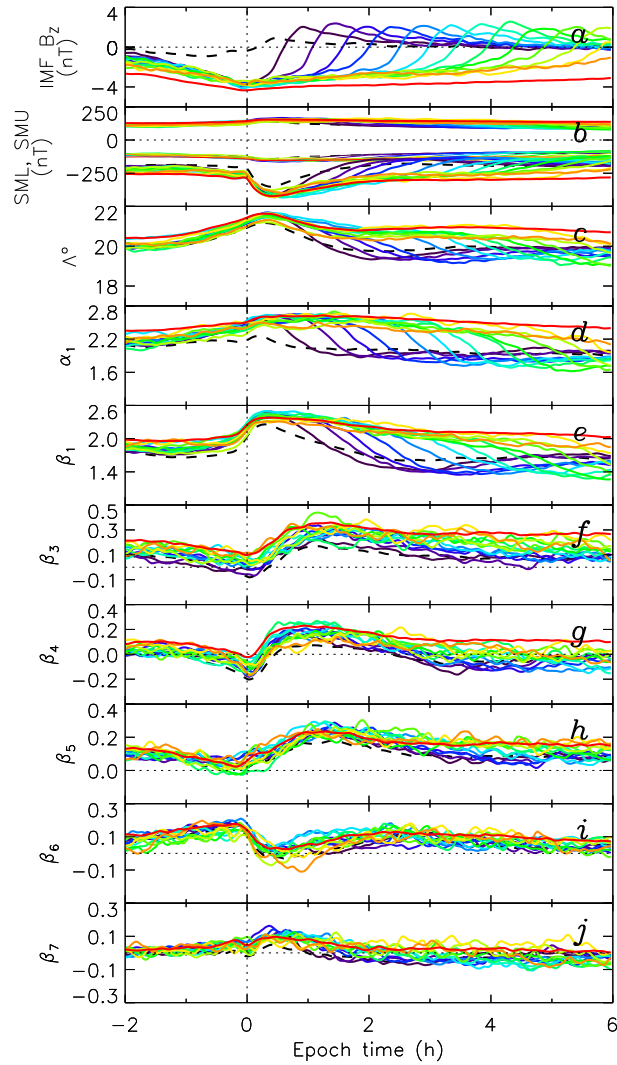


Figure 3. Superposed epoch analysis of substorms in the same format as Fig. 2. Substorms are categorised by the length of time that the IMF remains southwards after onset: over 30, 60, 90, ..., 360 mins (blue to red traces). The number of substorms in each category is 699, 495, 472, 344, 337, 290, 253, 229, 234, 184, 222, 151, 189, and 2786, respectively. Substorms which do not fit the selection criteria (2409) are shown as a dashed line.

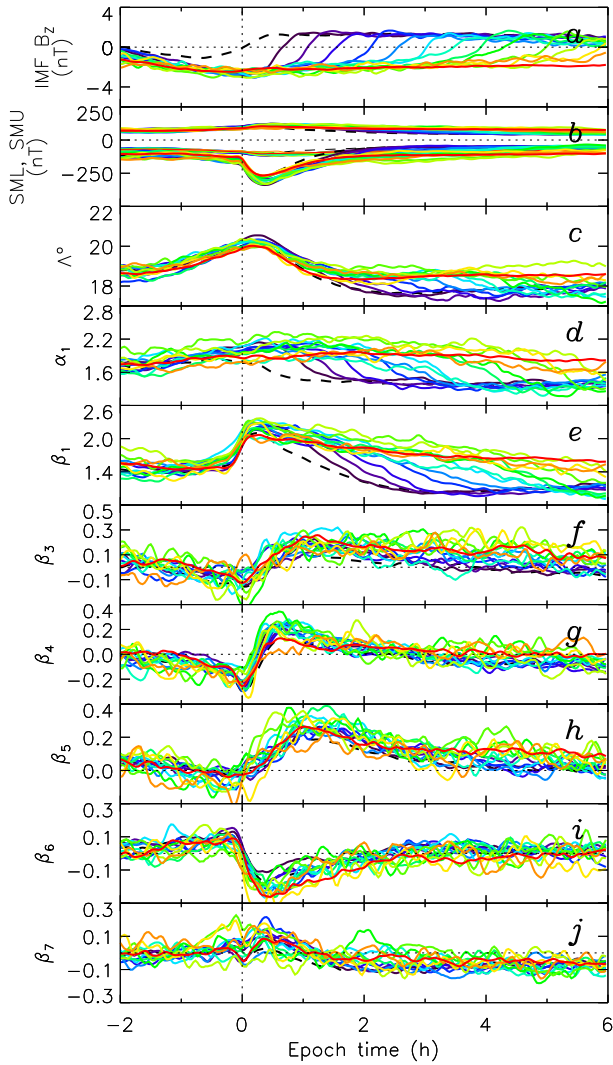


Figure 4. Superposed epoch analysis of substorms in the same format as Fig. 2, and selected in the same way, though restricted to those substorms which are not followed by a subsequent substorm for at least 6 hours. The number of substorms in each category is 59, 180, 128, 85, 97, 58, 48, 65, 37, 35, 27, 26, 31, 306, respectively. Substorms which do not fit the selection criteria (847) are shown as a dashed line.

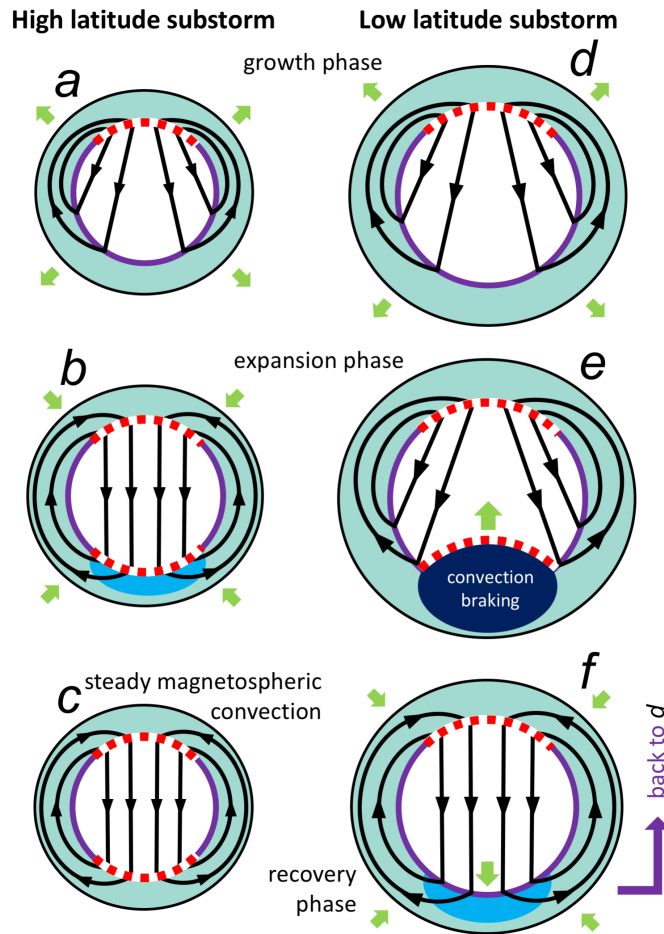


Figure 5. A schematic of the development of high- and low-latitude substorms, panels (a) to (c) and (d) to (f), respectively, in response to prolonged solar wind-magnetosphere coupling. Each panel has noon at the top. Black arrowed lines show convection streamlines, the purple circle is the open/closed field line boundary enclosing the polar cap. Red dashed lines show portions of the OCB mapping to active reconnection X-lines at the magnetopause or in the magnetotail. Blue regions indicate the location of the auroral bulge and whether it is of high (dark blue) or low (light blue) ionospheric conductance; along these portions of the OCB the ionospheric flow crosses the boundary, along other portions the flow is adiaroic. Green arrows indicate expansion or contraction of the auroral zone and polar cap, or motions of the OCB.

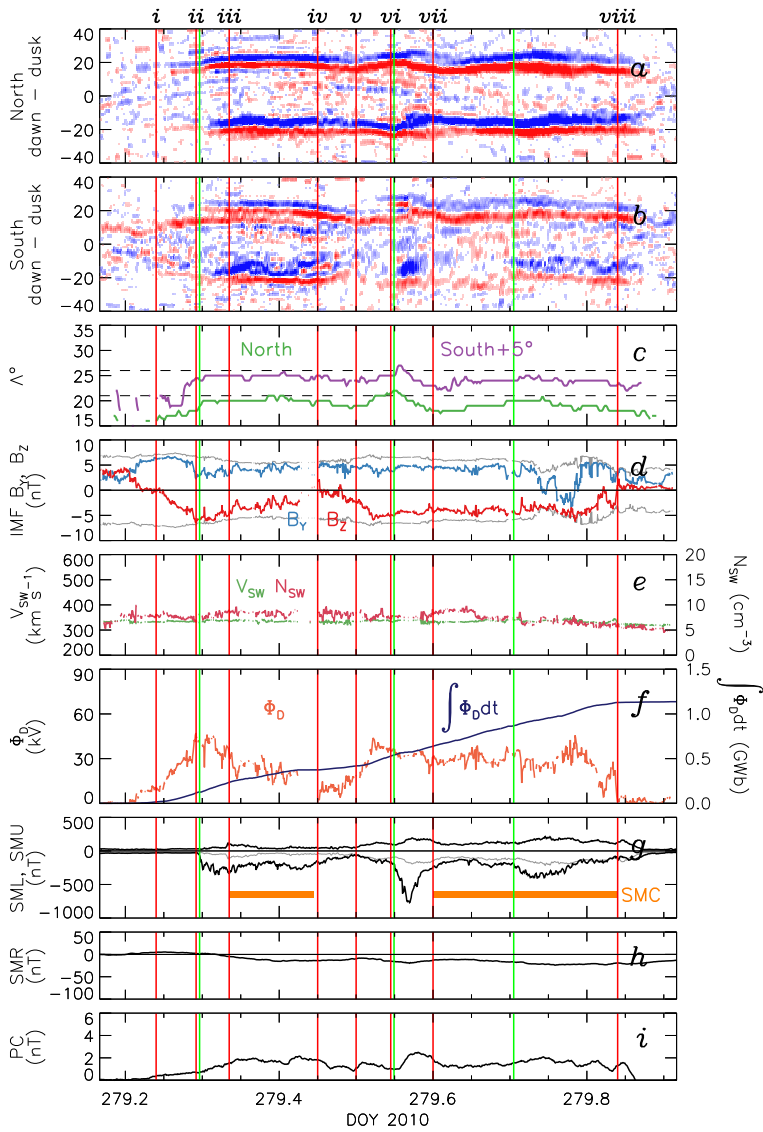


Figure 6. Solar wind and magnetospheric dynamics for the period 04 to 22 UT, 6 October 2010. (a) and (b) Field-aligned current density along the dawn-dusk meridian in the northern and southern hemispheres, with red and blue indicating upwards and downwards FAC, with the colour scale saturating at $0.5 \mu\text{A m}^{-2}$. (c) Radius, Λ , of the northern and southern FAC ovals, with the southern hemisphere being displaced by 5° for clarity; the horizontal dashed lines indicate $\Lambda = 21^\circ$. (d) IMF B_Y and B_Z . (e) Solar wind speed and density. (f) Dayside reconnection rate, Φ_D , and the time integral of Φ_D . (g) The electrojet indices SML and SMU, with periods of steady magnetospheric convection indicated by orange bars. (h) The ring current index SMR. (i) The PC index. Vertical green lines show SuperMAG substorm onsets, and vertical red lines

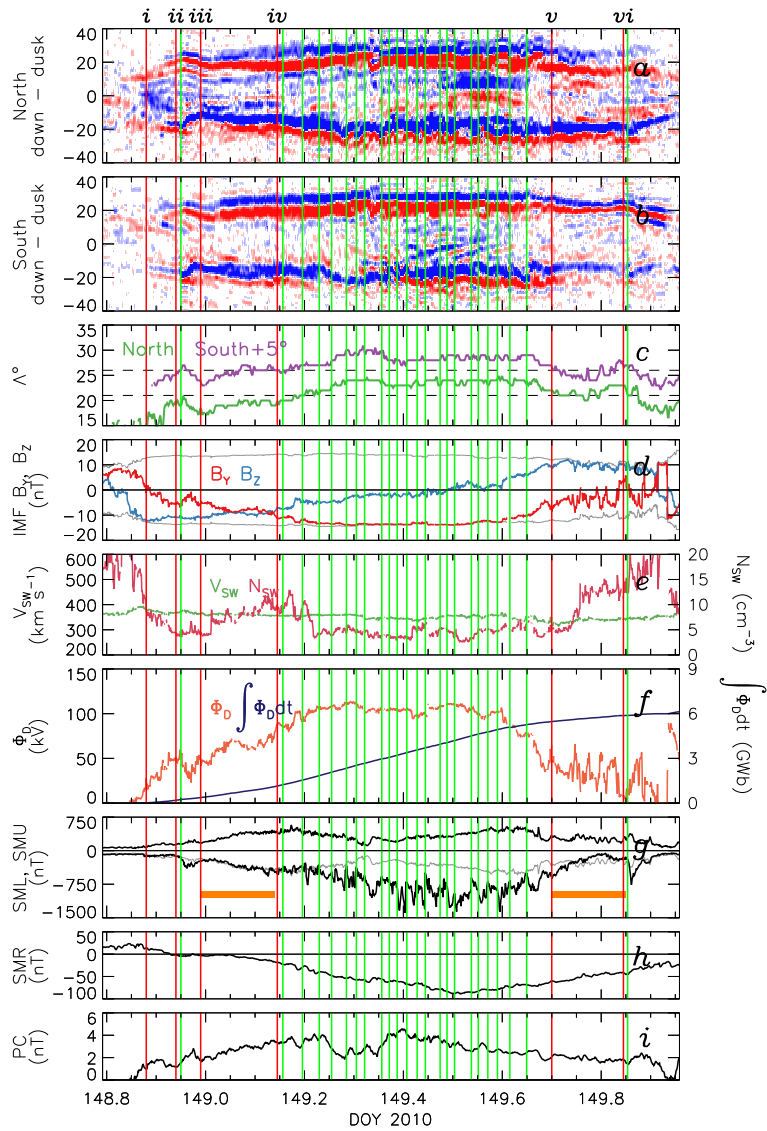


Figure 7. Solar wind and magnetospheric dynamics for the period 19 UT, 28 May 2010, to 23 UT on the following day, in the same format as Fig. 6.

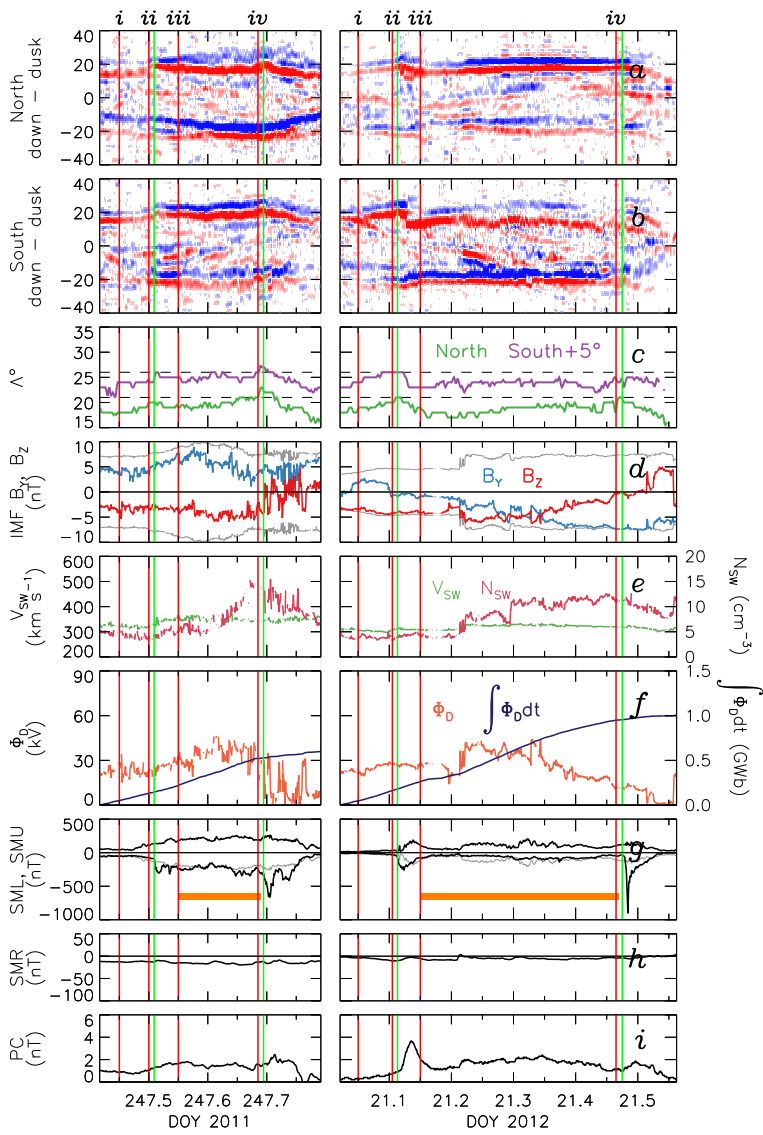


Figure 8. Solar wind and magnetospheric dynamics for the period 10 to 19 UT, 4 September 2011, and 00:30 to 13:30 UT, 21 January 2012, in the same format as Fig. 6.

Figure 1.

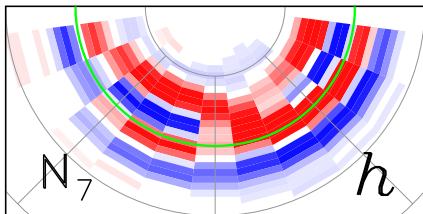
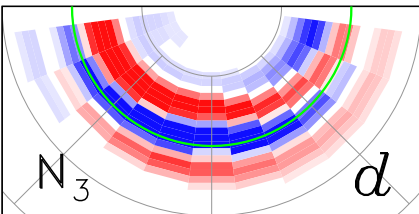
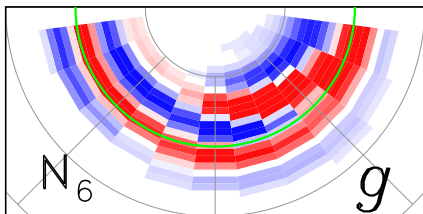
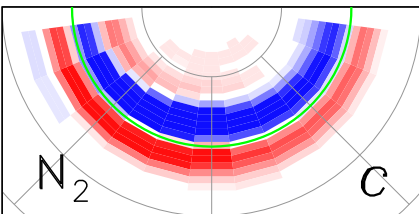
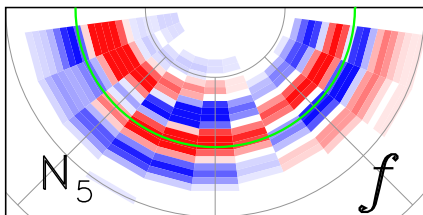
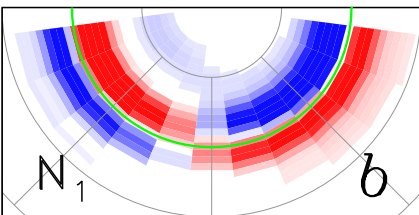
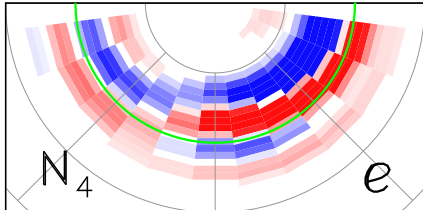
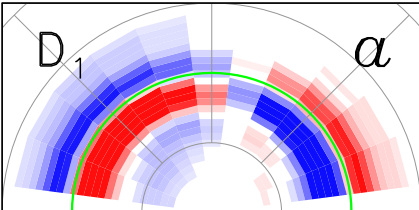


Figure 2.

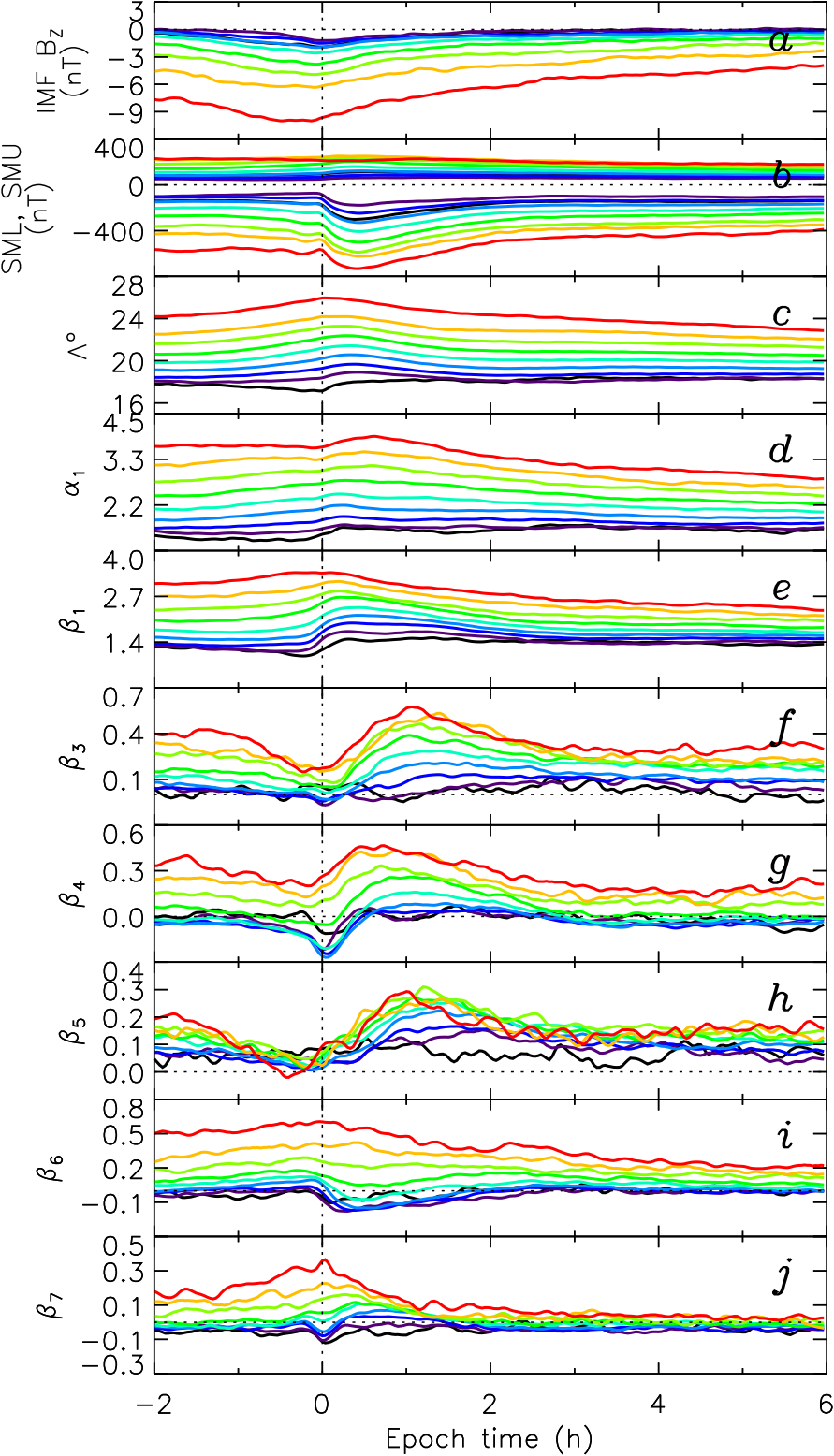


Figure 3.

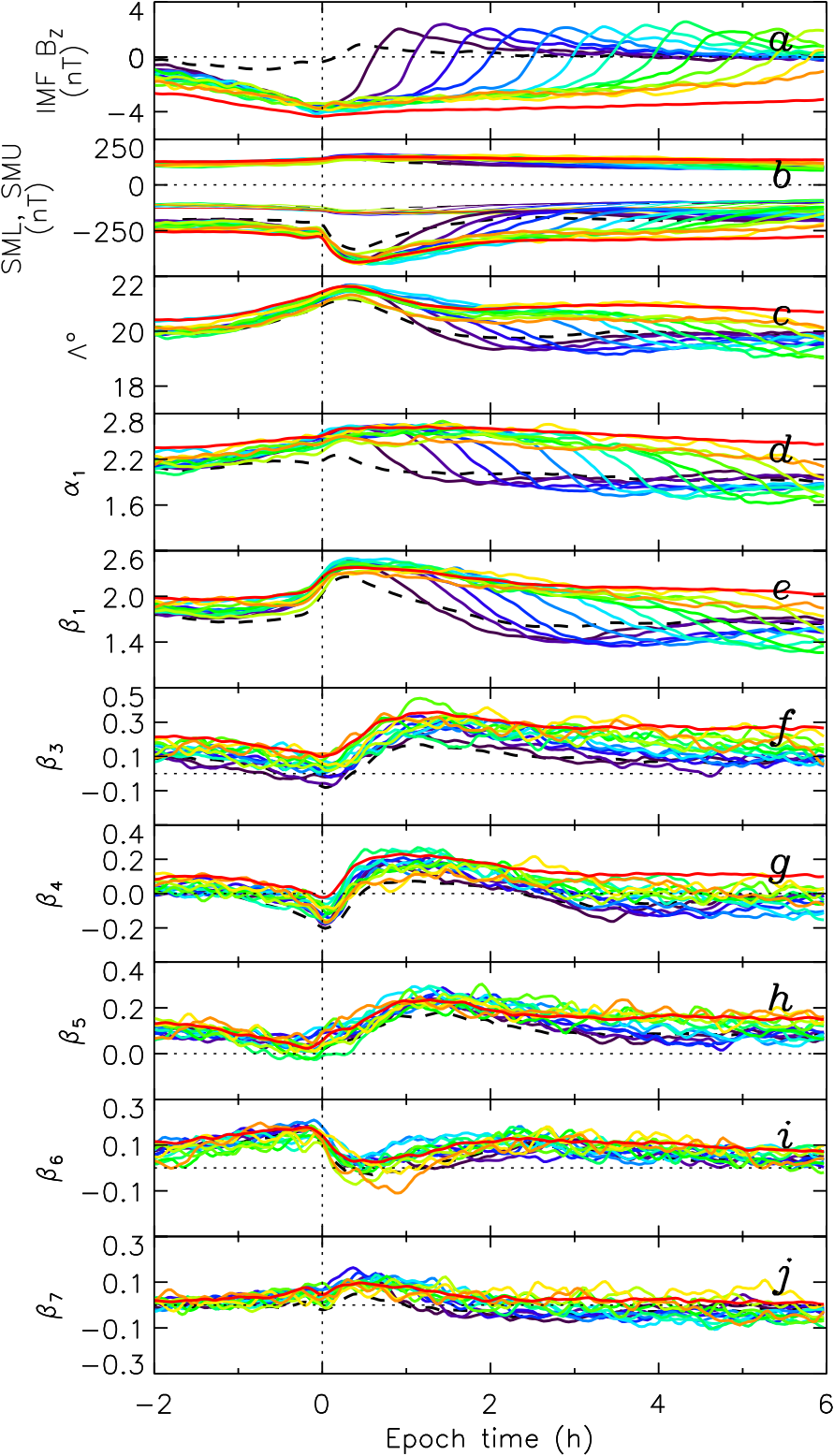


Figure 4.

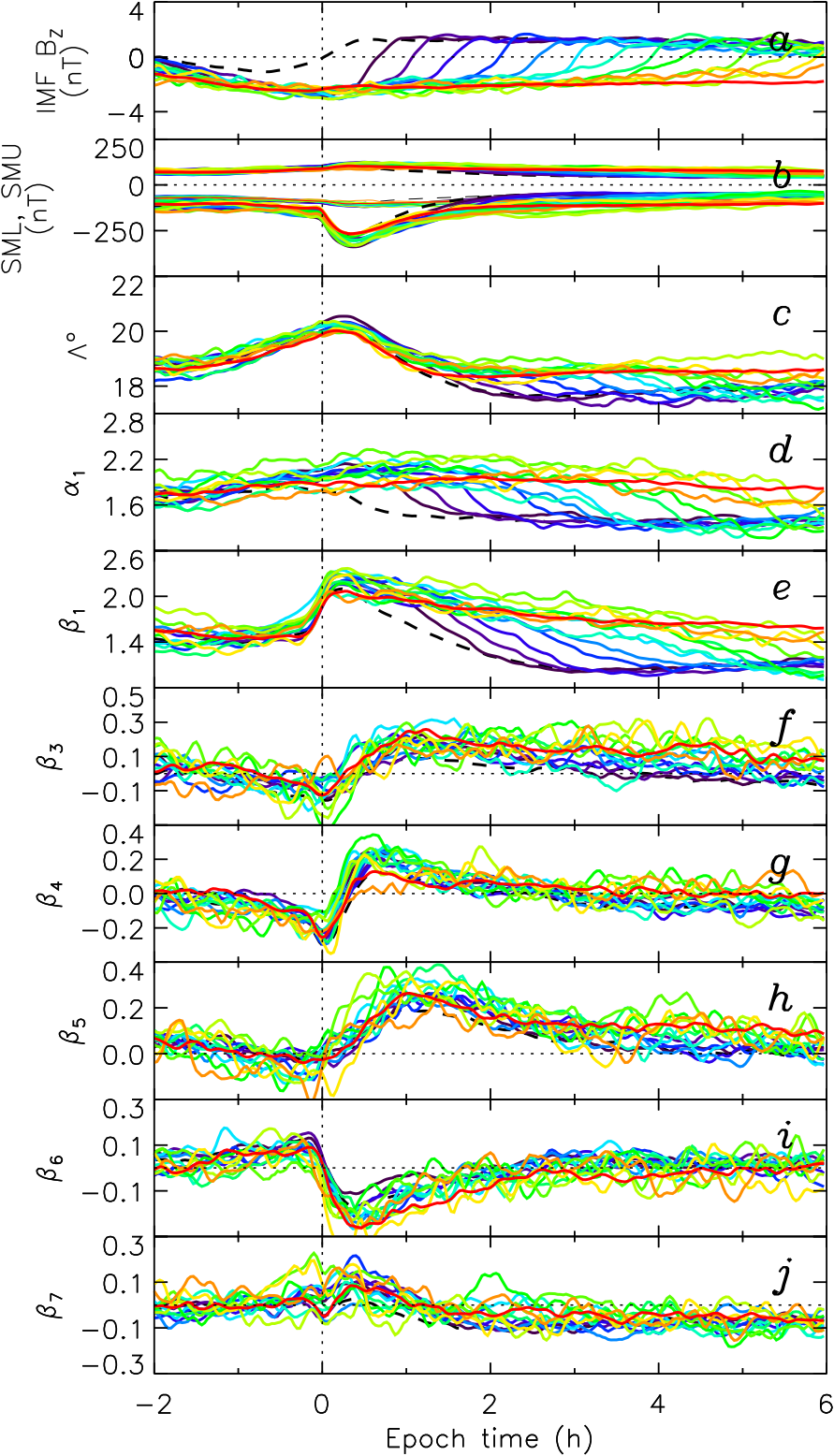
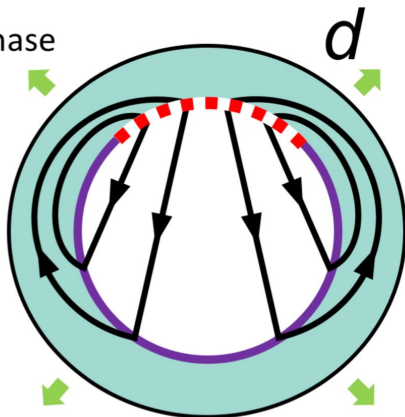
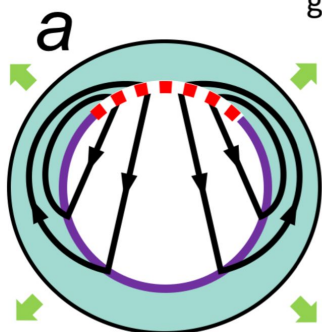


Figure 5.

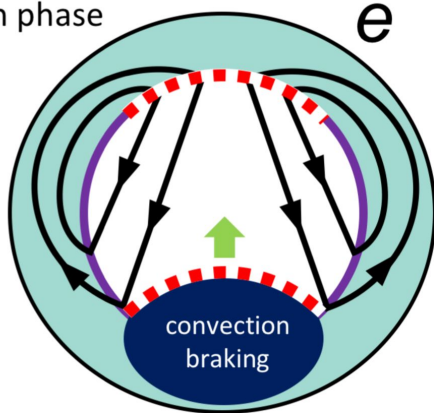
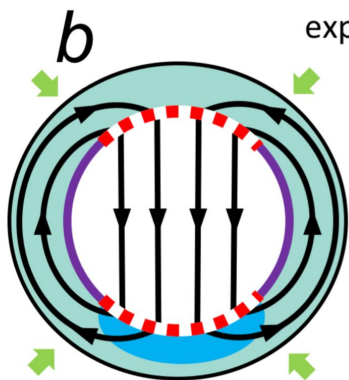
High latitude substorm

Low latitude substorm

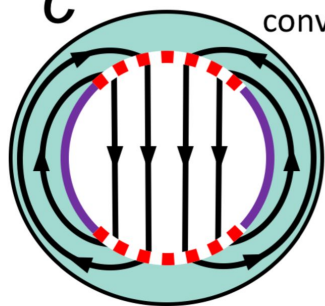
growth phase



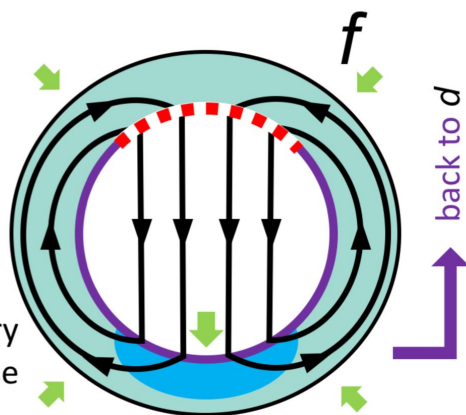
expansion phase



steady magnetospheric convection



recovery phase



back to d



Figure6.

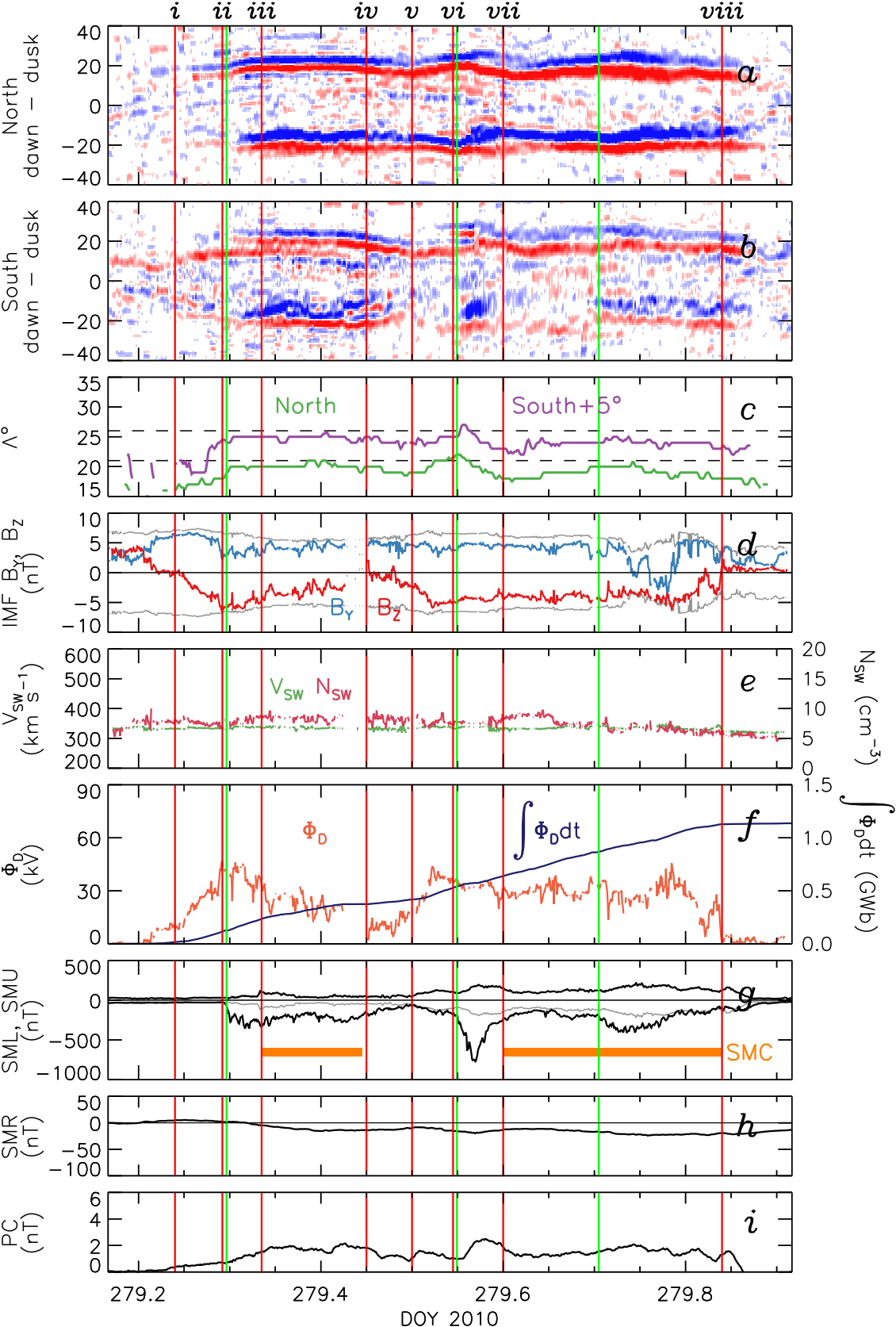


Figure 7.

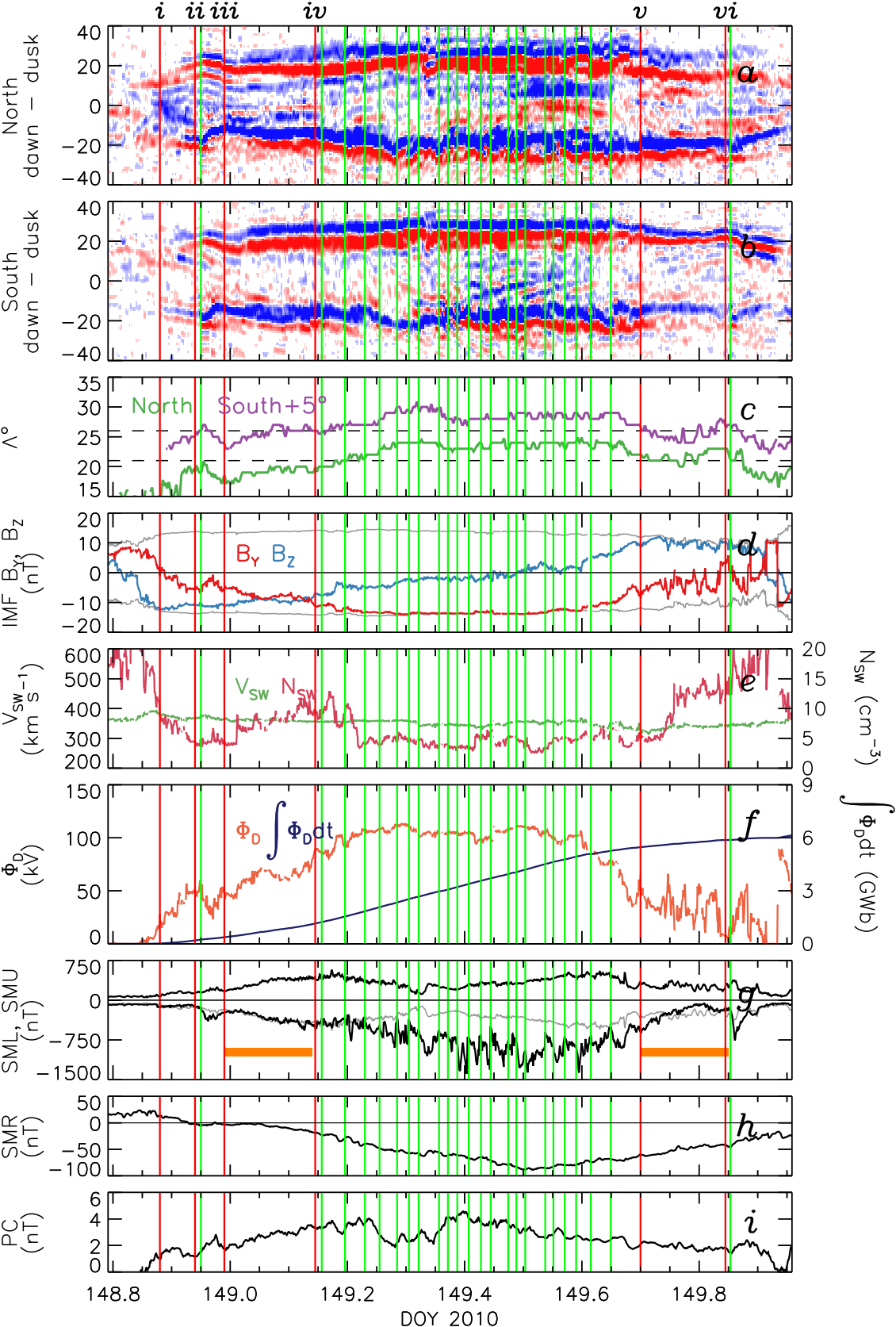


Figure 8.

